

WATER RESOURCES - RECONNAISSANCE SERIES

REPORT 49

WATER-RESOURCES APPRAISAL OF BUTTE VALLEY,
ELKO AND WHITE PINE COUNTIES, NEVADA

By

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Prepared cooperatively by the
Geological Survey, U.S. Department of the Interior

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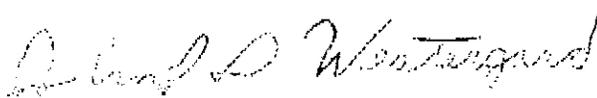
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FOREWORD

The program of reconnaissance water-resources studies was authorized by the 1960 Legislature to be carried on by the Department of Conservation and Natural Resources in cooperation with the U.S. Geological Survey.

This report is the 49th report prepared by the staff of the Nevada District of the U.S. Geological Survey. These 49 reports describe the hydrology of 148 valleys.

The reconnaissance surveys make available pertinent information of great and immediate value to many State and Federal agencies, the State cooperating agency, and the public. As development takes place in any area, demands for more detailed information will arise, and studies to supply such information will be undertaken. In the meantime, these reconnaissance-type studies are timely and adequately meet the immediate needs for information on the water resources of the areas covered by the reports.



Roland D. Westergard
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Division of Water
Resources

July 1968

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WATER-RESOURCES APPRAISAL OF BUTTE VALLEY,

ELKO AND WHITE PINE COUNTIES, NEVADA

By Patrick A. Glancy

SUMMARY

Butte Valley includes two topographically closed valleys in east-central Nevada. The valleys, referred to in this report as northern Butte Valley and southern Butte Valley, cover a total area of about 1,000 square miles. Principal hydrologic facts and estimates resulting from this reconnaissance are summarized in table 1.

Precipitation within the area is assumed to be the main source of water supply to the valley-fill reservoir. However, carbonate rocks, which are prevalent in the region, may be highly transmissive locally; therefore, the precise location of recharge boundaries is unknown. For purposes of reconnaissance estimates and calculations, recharge boundaries in most places are assumed coincident with topographic boundaries. However, regional intervalley flow through carbonate rocks may occur.

The principal known aquifers occur in the valley fill at generally shallow depths. However, only about 20 wells have been drilled, and no data are available regarding deeper parts of the valley fill. The carbonate rocks of the area also constitute an unexplored but probably significant ground-water system.

Natural ground-water discharge in the area is mainly by evapotranspiration, which totals almost 20,000 acre-feet per year.

Only five significant perennial streams occur in the area. Characteristics of numerous ephemeral stream channels and the nature of the geologic terrane suggest that mean annual runoff yields of the area are unusually low with regard to precipitation input when compared with most Nevada drainage basins.

Chemical analyses of water from 21 well, spring, and stream sources show very good quality of water in the area. All constituents in the water tested are acceptable for most common purposes requiring good quality of water, based on results of the

Table 1.--Hydrologic summary

/Estimates are in acre-feet per year, except where noted/

	Northern Butte Valley	Southern Butte Valley	Butte Valley Total
Approximate valley area (square miles)	270	730	1,000
Probable hydrologic closure	(a)	(b)	(a,b)
Surface-water runoff from mountains	2,700	9,400	12,000
Ground-water recharge	3,000	15,000	19,000
Water consumed by crops	1,100	1,200	2,300
Evapotranspiration of ground water	6,900	11,000	18,000
Ground-water discharge	8,700	12,000	21,000
Perennial yield	6,000	14,000	20,000
Transitional storage reserve ^{1/}	300,000	600,000	900,000

1. Total quantity available, in acre-feet.

a. May be some subsurface inflow through carbonate rocks from adjacent mountains beyond surficial watershed. An estimated outflow of 800 acre-feet per year to Ruby Valley occurs through valley fill.

b. May be some subsurface outflow through carbonate rocks to adjacent valleys.

constituents investigated; however, detrimental concentrations and constituents may occur that were not determined by the analyses.

Streamflow is used for irrigation but ground-water resources of the area are mainly undeveloped. Surface water and springflow are currently used mainly to irrigate about 2,000 acres of native pasture and alfalfa. Ground-water pumpage for domestic and stock use probably did not exceed 35 acre-feet in 1967. A few springs are used for stock-watering and domestic purposes. Future development may depend principally on whether the quantity of available water is adequate for the intended use and the economic limitations governing the extraction of ground water.

INTRODUCTION

Purpose and Scope of the Investigation

Nevada is currently experiencing a rapid growth in population and associated development that began more than a decade ago. Increased water requirements for domestic, industrial, agricultural, and recreation uses have accompanied this growth. Anticipating these increasing water needs, the Nevada State Legislature enacted legislation (Chapter 181, Statutes, 1960) authorizing an expansion of the established program of hydrologic investigations which were being conducted by the U.S. Geological Survey in cooperation with the Nevada Department of Conservation and Natural Resources. The legislation has provided financing, in the form of matching funds with the Federal Government, to conduct appraisals of the water resources of the State. The appraisals are being made as a series of reconnaissance investigations of individual or groups of areas. This investigation is the 49th in the series.

The objectives of this study, which are reconnaissance in scope and depth of treatment, are to (1) describe the general geology as it relates to the water resources, (2) appraise the source, occurrence, movement, and general chemical quality of water in the area, (3) estimate average annual recharge to and discharge from the ground-water reservoir, (4) evaluate the surface-water resources in the valleys, and (5) provide preliminary estimates of the perennial yield and transitional storage reserve.

This investigation was made under the general supervision of G. F. Worts, Jr., district chief in charge of hydrologic studies by the Geological Survey in Nevada. Field work was begun by J. L. Hughes, engineer, in the summer of 1965 and completed by the author in August 1967.

Location and General Geographic Features

Butte Valley, Nevada, is enclosed by lat $39^{\circ}27'$ and $40^{\circ}33'$ N., and long $114^{\circ}49'$ and $115^{\circ}17'$ W. (fig. 1). The area is in the north-central part of White Pine County and the south-central part of Elko County and includes about 1,000 square miles. It has a north-south length of about 76 miles and an east-west width of about 17 miles. It contains two surface-drainage basins separated by a gently sloping alluvial divide. These valleys are shown on plate 1 as Butte Valley - northern part and Butte Valley - southern part; throughout the text they are generally referred to as northern Butte Valley and Southern Butte Valley. Northern Butte Valley has an area of about 270 square miles; southern Butte Valley, about 730 square miles.

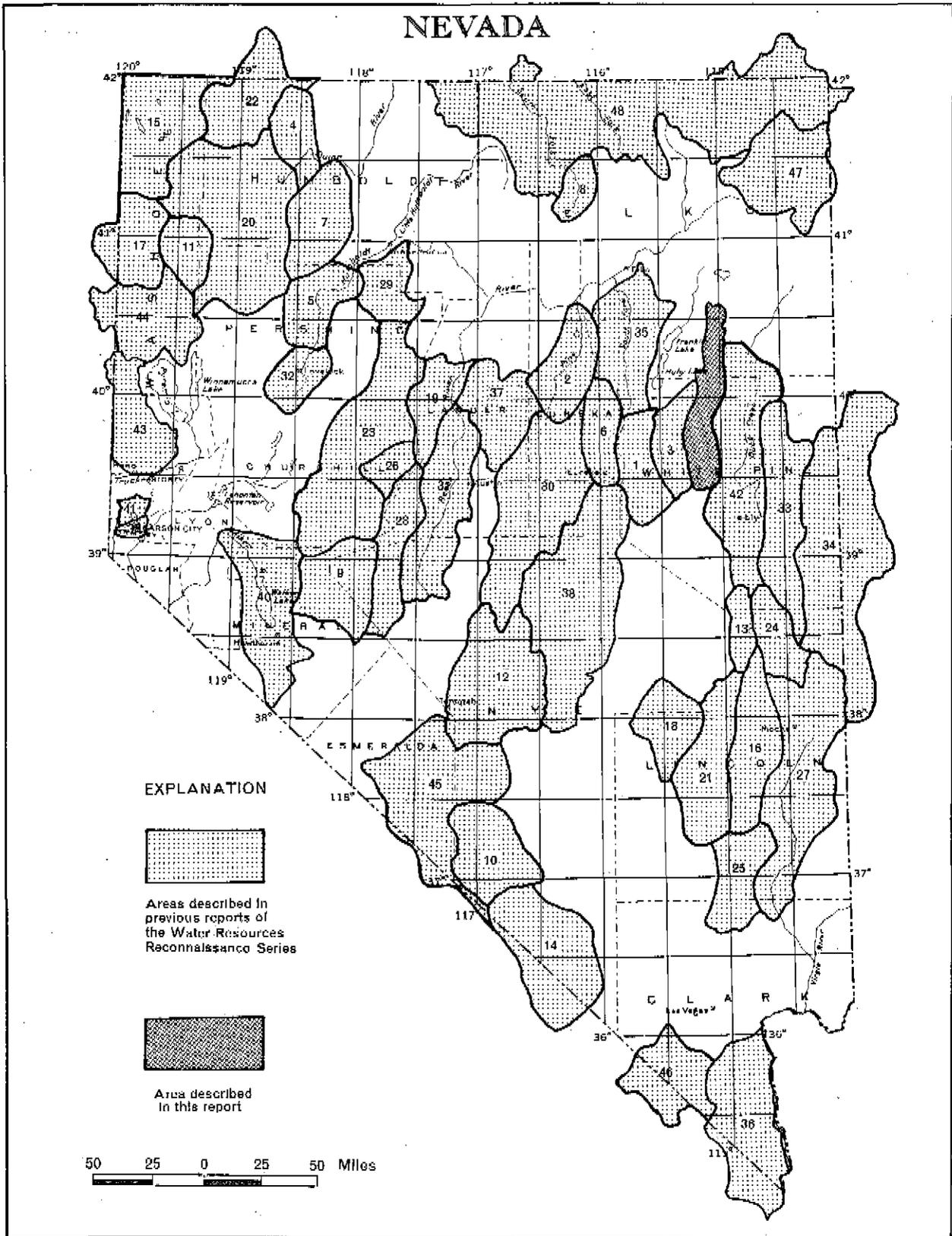


Figure 1.—Areas in Nevada described in previous reports of the Water Resources Reconnaissance Series and the area described in this report

Butte Valley is very sparsely populated; altogether, the total permanent residents probably number less than 30. Southern Butte Valley contains two ranches, the Uhalde Ranch in the southwestern part and Paris Ranch in the northern part. The Robert Healy Ranch, in the southwestern part of northern Butte Valley, on the Odgers Ranch Indian Reservation (pl. 1), currently houses the only permanent residents of that valley.

Ranching is the sole occupation of the area's residents and historically has been the main occupation. In 1967, the Bear Creek Mining Corporation reportedly was conducting exploration in and near T. 22-N., R. 61-E., in southern Butte Valley, and although their findings may influence the future economy of the area, no permanent residents are presently involved. In addition, some oil exploration has been done in the area.

The nearest towns are to the east and southeast in adjacent Steptoe Valley and include the small settlements of Currie and Cherry Creek, and the towns of Ely and McGill. The small town of Currie is about 8 miles east of northern Butte Valley.

Early Development

The earliest use of water resources by white men in Butte Valley was by explorers in the mid-1800's. Springs around the edges of the valley probably were used by Major George Chorpenning's overland mail carriers. Chorpenning was a trailmaker and Indian fighter and also the contractor in charge of the first overland mail service during the early 1850's (Chapman, 1932, p. 41). Much of the route his mail service followed was originally traveled by Major Howard Egan, a scout and captain under Brigham Young. The route and some of the watering places established in the Butte Valley area by Chorpenning's ventures were in general also utilized by Capt. J. H. Simpson of the Corps of Topographical Engineers during his exploration of the Great Basin. Capt. Simpson, also employing Major Howard Egan as a guide, had as his objective the search for a direct wagon route between Camp Floyd, Utah, and the new settlement at Genoa in Carson Valley of western Nevada. He traveled up Egan Canyon, over the Cherry Creek Mountains, and across Butte Valley on May 15, 1859. (Simpson, 1876, p. 60-62).

The famous Pony Express of 1860-61 utilized Simpson's route through this part of the State and was also dependent for its water supply on the springs discharging from the consolidated rocks bordering the valley. In the period between the termination of the Pony Express and the influx of permanent settlers, the valley was intermittently used as a haven for horse thieves and their stolen livestock (Mr. B. Paris, oral commun., 1967).

According to Mr. Paris, Richard Stratton was the first permanent settler. He arrived by wagon about 1880, and established a ranch with dwellings for his family in sec. 22, T. 26 N., R. 62 E. The general area of the ranch is still occupied today by the Paris family and their employees who along with residents of the Uhalde Ranch in T. 20 N., R. 60 E., and members of the Robert Healy family on the Odgers Ranch Indian Reservation in T. 28 N., R. 61 E., are the only permanent residents of Butte Valley.

The major development of the area has evolved around the raising of livestock. Livestock grazing utilizes seasonal forage and native meadowgrasses. At least two large areas have been seeded to crested wheat grass, one in T. 19 N., R. 61 E., and the other along the west flank of the Cherry Creek Mountains between Paris Creek and the Stratton Ranch.

Previous Work

An interesting historical account of early exploration in the valley is that by Simpson (1876). Facts directly concerned with the geology of the area are found in the following reports: Emmons (1878), Snelson (1955), Harlow (1956), Nelson (1956), Fritz (1957 and 1961), Adair and Stringham (1960), Douglas (1960), Misch (1960), Threet (1960), and Stevens (1965). Although the above list is not intended to be exhaustive, it references many of the important detailed geologic investigations of the report area.

Hydrology of the area with reference to livestock-watering development was reported by Snyder (1963). The presence of Pleistocene lakes in the area is documented by Snyder, Hardman, and Zdenek (1964). Hydrology of adjacent areas has been described in reports by Maxey and Eakin (1949), Eakin and others (1951), and Eakin (1961, 1966, and 1967).

Acknowledgments

The writer is grateful to the residents of the valley who provided valuable information about water use and development. Mr. Paris was particularly informative about the early settlement of the valley, and his generous donation of time and information is gratefully appreciated. All landowners are cordially thanked for granting permission of access to their property. Dr. Peter Misch, University of Washington, and Dr. W. H. Fritz, Geological Survey of Canada, provided exceptionally well-timed aid by assisting in the procurement of a geologic map of the northern Egan Range. The map was originally prepared by Dr. Fritz as part of his doctoral thesis at the University of Washington. The Bureau of Land Management and Robert Millard of Millard-Spink Associates, Inc. furnished hydrologic information on southern Butte Valley.

GENERAL HYDROLOGIC ENVIRONMENT

Physiographic Features

The report area is a small topographically closed segment of the Great Basin. It consists of an elongate structural depression nearly surrounded by mountains. The valley actually comprises two separate drainage basins, previously referred to as northern and southern Butte Valleys. They are separated by a narrow alluvial divide of very gentle relief compared to that generally separating Butte Valley from adjacent valleys.

Altitudes in northern Butte Valley range from about 5,990 feet on the playa to 9,498 feet at Mt. Taylor in the Cherry Creek Mountains, and therefore, the maximum relief is about 3,500 feet. The altitudes of southern Butte Valley range from about 6,160 feet on the valley floor to 10,600 feet in the Cherry Creek Mountains, and the maximum relief is about 4,400 feet. The Cherry Creek Mountains, forming much of the eastern boundary of the area, are also the highest range in the area. Maximum altitude of the Butte Mountains is about 9,030 feet and that of Medicine Range is about the same. The highest part of Spruce Mountain within the area has an altitude of about 10,080 feet. West Buttes, Palomino Ridge, Valley Mountain, and Delcer Buttes have maximum altitudes of 7,654, 7,383, about 7,200, and 6,908 feet, respectively.

Observation of the terrane discloses the following topographic characteristics: (1) relatively flat valley floors; (2) a system of smoothly coalescing alluvial fans joining the bases of the mountain ranges with the alluvial valley floors; (3) the alluvial fans are generally typified by surfaces that appear smooth and undissected compared to many alluvial fans in the basin and range physiographic province; and (4) generally rugged mountain masses that commonly rise abruptly and steeply above the heads of the alluvial fans.

According to relic shorelines visible on aerial photos, the maximum stillstand altitudes of Pleistocene lakes were as follows: northern Butte Valley, about 6,050 feet; southern Butte Valley, about 6,300 feet. The lake in southern Butte Valley may have spilled northward through the narrow alluvial gap into northern Butte Valley during its highest stand, and that in northern Butte Valley probably spilled to or was connected with the lake in Ruby Valley to the west.

Geologic Units

A generalized geologic map of the area is shown on plate 1, and a summary of geologic (lithologic) units and their character is included in table 2. Differentiation of four geologic units was based mainly on their hydrologic properties. For this reconnaissance, the alluvial deposits were grouped into older and younger alluvium. Criteria for separation of the units are: (1) alluvial areas where erosion and deformation appears to have been recently dominant over deposition are classified as older alluvium, (2) alluvial areas where deposition has been the recent dominant process are classed as younger alluvium, (3) alluvial deposits below the maximum stillstand of Pleistocene lakes, determined mainly by aerial photos, are considered younger alluvium, and (4) younger alluvium is assumed to have been deposited mainly during relatively recent geologic time and is generally less than 100 feet thick.

The consolidated rocks were grouped into noncarbonate and carbonate rocks (pl. 1). The carbonate rocks contain solution cavities or enlarged joints which, where interconnected, readily convey ground water. The noncarbonate rocks generally are poorly permeable. Structural deformation may have increased or decreased the transmissivity of the consolidated rocks, depending on a variety of factors and conditions; therefore, the present hydrologic characteristics of the rocks differ in varying degrees from those common to the rocks at the time of their emplacement. This reconnaissance suggests that structural deformation is mainly restricted to the consolidated rocks; although deformation of the older alluvium is not as readily obvious and extensive, it may have altered locally the transmissive character of this unit.

Table 2.—Generalized geologic units

	Geologic age	Geologic unit	Thickness (feet)	General character and extent	Water-bearing properties
QUATERNARY	Holocene and Pleistocene	Younger alluvium	0-100±	Unconsolidated lenses of gravel, sand, silt, and clay comprising fluvial, lacustrine, and eolian deposits; fluvial deposits commonly contain large-size gravel and boulders; lacustrine deposits range from clay-size particles to sand and gravel bar deposits; eolian deposits mainly silt-size material with some sand; detritus composed mainly of material derived from bordering upland consolidated rocks and reworking of older alluvium; unit probably thickest in valley troughs; probably mantles older alluvium over much of its extent	Yields water to domestic and stock wells where saturated; yields are variable, depending on character of deposits encountered by wells, and range from about 5 gpm to possibly several hundred gpm
TERTIARY(?) AND QUATERNARY	Pleistocene and older(?)	Older alluvium	0-several thousand(?)	Unconsolidated to semiconsolidated deposits of boulders, gravel, sand, silt, and clay exposed around margins of valley and buried at generally shallow depth beneath younger alluvium; composed mainly of debris derived from bordering consolidated rock areas; detrital constituents may be somewhat altered from their original character at time of deposition; alteration of unit may include decrease in transmissivity caused by cementation; unit thinnest in upland areas and thickest beneath valley troughs; mantles consolidated rock; probably present nearly everywhere beneath younger alluvium	May yield water to some stock wells; yields probably vary depending on character of deposits and may be several hundred gpm or more
CAMBRIAN TO TRIASSIC	Triassic to Middle Cambrian	Carbonate rocks ^{2/}	0-several thousand	Mainly limestone and dolomite; unit may contain minor strata of noncarbonate rocks; structurally deformed; general extent shown on plate 1; unit probably occurs extensively beneath valley-fill deposits	Transmits water through fracture and solution cavities; inter-basin subsurface flow may occur through these conduits; unit untested by wells; yields water to springs and perennial streams
PRECAMBRIAN TO TERTIARY	Tertiary to Late Pre-Cambrian	Non-carbonate rocks ^{2/}	—	Igneous, metamorphic, and sedimentary rocks; igneous rocks are mainly volcanics including basalt and ignimbrite; some small intrusives also present, mainly in northern Butte Valley; metamorphic rocks are mainly argillite and quartzite; sedimentary rocks include sandstone, siltstone, and shale; all rock types have been structurally deformed; general extent of unit shown on plate 1	Unit generally untested by wells; yields minor amounts of water to springs; might yield small amounts of water to wells locally from fracture zones; considered poorest water-yielding unit in area

1. Carbonate and noncarbonate rocks may be interbedded and are not discrete geologic time units.
2. Synthesized from various reports, as credited on plate 1.

VALLEY-FILL RESERVOIR

Extent and Boundaries

The valley-fill reservoir is formed by the older and younger alluvium, which extends continuously from the south end to the north end of Butte Valley, as shown on plate 1. However, the topographic divide between the northern and southern parts of Butte Valley also nearly coincides with the ground-water divide. Therefore, the valley-fill reservoir in northern Butte Valley is considered contiguous to but separate from the reservoir in southern Butte Valley. The ground-water divide, which could become transient if substantial ground-water development occurred, forms the hydraulic divide between the two areas.

The sides of the valley-fill reservoir are formed by the consolidated rocks, where they are present. The western side of northern Butte Valley locally is valley fill, which extends continuously into Ruby Valley. The alluvium at the north end of the valley extends continuously beneath the topographic divide into Clover Valley, but the divide area probably is underlain at shallow depth by consolidated rocks.

Occurrence and Movement of Ground Water

The known depth below land surface to the zone of saturation in the valley-fill reservoir ranges from about land surface in spring areas to about 150 feet in well 20/61-6d1. Throughout most of the area, shallow depths to water exist near the valley floors and generally increase toward the mountains. Exceptions to this generality occur in the northern part of southern Butte Valley and the southern part of northern Butte Valley where springs discharge from the alluvium at altitudes above the valley floor. This occurrence and the north-trending alignment of the springs strongly suggest structural control of the spring orifices. Westward ground-water flow toward the valley floor is being impeded by decreased permeability, causing the spring discharge. The movement in northern Butte Valley is generally toward the playa in the northwestern part of the valley. Water-level altitudes indicate that subsurface flow continues westward to Ruby Valley beneath the alluvial divide. (See section, "Subsurface outflow to adjacent valleys.")

In southern Butte Valley, ground water in the valley-fill reservoir moves toward the playa in the southern part of the valley where the depth to water is at least 50 feet (water-level altitude about 6,100 feet). Whether the sparse phreatophytes in the playa area consume sufficient ground water to keep the water level depressed to this depth is not known. The possibility of leakage into the underlying carbonate rocks and outflow to adjacent valleys is discussed in the section, "Subsurface outflow to adjacent valleys."

CARBONATE-ROCK RESERVOIR

Carbonate rocks locally may form a storage and transmission medium for ground water when fracture systems are enlarged by percolating waters that form solution cavities. Carbonate rocks, as shown on plate 1, make up a considerable part of the exposed, and probably also the buried, consolidated rocks of the area. Structural deformation of these rocks is intense and widespread. Field examination of exposed formation disclosed that some of the carbonate rocks are riddled by solution cavities. Limestone collapse breccias were also observed in the carbonate-rock areas. Assuming that these characteristics are also common in the subsurface, the carbonate rocks of the area are probably capable of containing and transmitting appreciable quantities of water.

No known water wells have been drilled very deeply into the carbonate rocks of the area; therefore, their water-yielding capabilities remain untested.

INFLOW TO THE VALLEY-FILL RESERVOIR

Precipitation

The general climate of east-central Nevada is arid to semiarid. The valley floors are arid and most precipitation falls on the surrounding mountains. The higher peaks are occasionally wet enough to be considered subhumid. Although no weather-recording stations are in Butte Valley, several are in nearby areas (fig. 2).

Precipitation is the source of virtually all water entering the hydrologic system of the report area. Throughout the area annual precipitation probably ranges from about 6 inches on the valley floor to more than 20 inches on the tops of higher mountains. This assumption is based on selected recorded data for adjacent areas (table 3), the area's topographic characteristics, and a precipitation map of Nevada (Hardman, 1965). Some precipitation occurs as rain, but much of it falls as snow during the winter and early spring, particularly at higher altitudes. Most regional precipitation occurs during late autumn, winter, and spring, as shown by data in table 4. Regional precipitation is generally lowest and least expected during the summer. However, high intensity local thundershowers occur unpredictably during the summer.

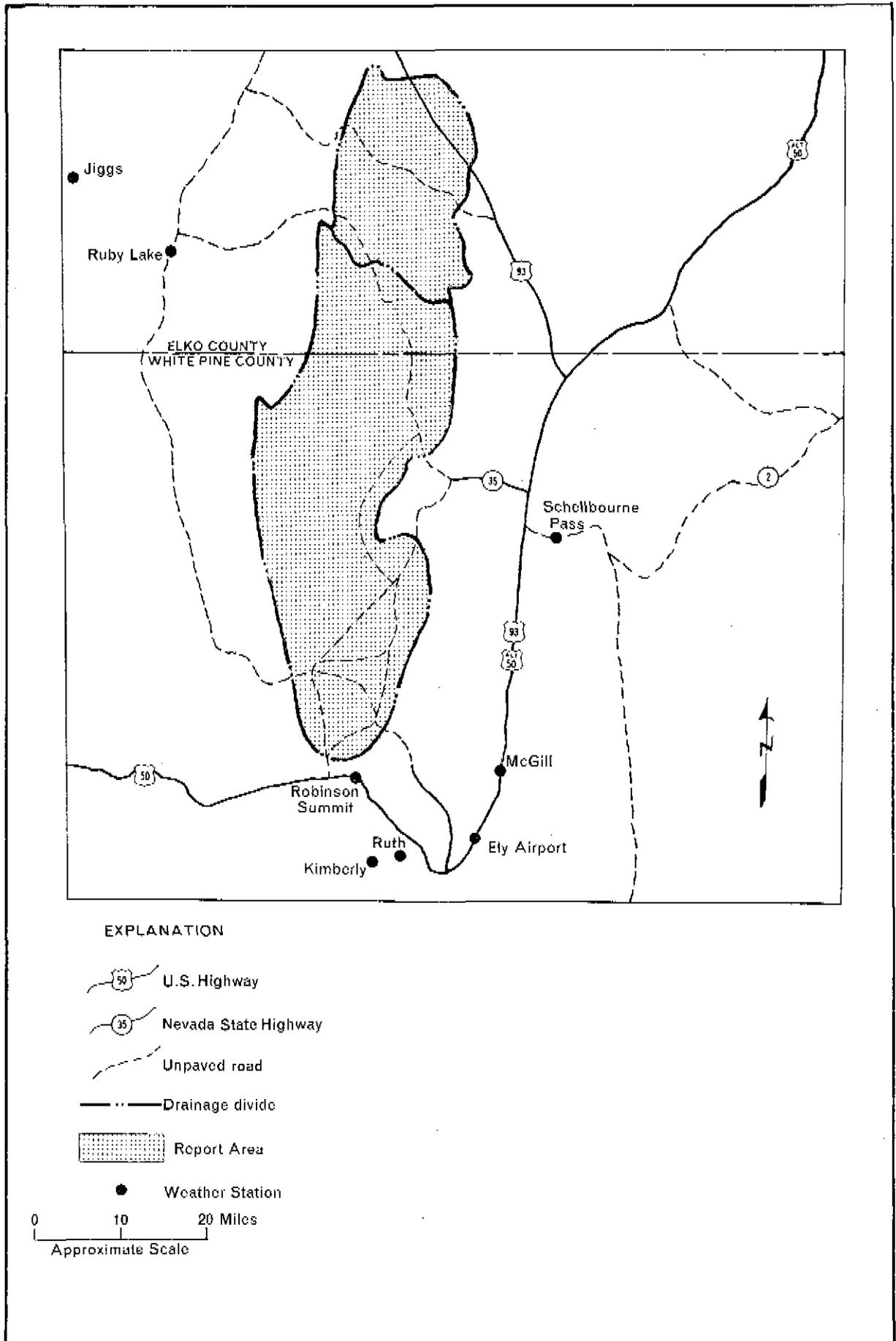


Figure 2.— Location of roads and weather stations.

Table 3.--Summary of average annual precipitation at selected stations

[Summarized from published records of the U.S. Weather Bureau/

Station	Location ^{1/}	Altitude (feet)	Period of record (years)	Average annual precipitation (inches)	Remarks
Ely Airport	17/63-35	6,257	28 years; 1939-66	8.34	
Jiggs	30/56-34	5,450	55 years; 1910-42, 1945-66	12.10	
Kimberly	16/62-8	7,250	29 years; 1929-57	13.24	
McGill	18/64-28	6,340	54 years; 1913-66	8.93	
Robinson Summit	18/61-23	7,630	13 years; Oct. 1953- Sept. 1966	9.94	Storage gage
Ruby Lake	27/58-19	6,012	24 years; 1940-43, 1945-50, 1952-55, 1957-66	12.49	
Ruth	16/62-3or4	6,832	8 years; 1959-66	10.76	
Schellbourne Pass	22/65-8	8,150	10 years; 1955-64	11.20	Storage gage

1. Station locations shown in figure 2.

Table 4.--Average monthly and annual precipitation,
in inches, at selected stations

From published records of the U.S. Weather Bureau

Station ^{1/}	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.	Oct.	Nov.	Dec.	Annual
Ely Airport	.63	.58	.89	1.01	.86	.78	.54	.50	.67	.65	.59	.64	8.34
Jiggs	1.12	1.05	1.19	1.48	1.51	.87	.55	.52	.59	1.04	.94	1.23	12.10
Kimberly	1.50	1.45	1.57	1.28	1.07	.67	.86	.90	.69	.88	.82	1.41	13.24
McGill	.81	.81	.98	1.03	1.15	.66	.63	.94	.62	.72	.63	.78	8.93
Ruby Lake	1.09	1.24	1.44	1.24	1.14	.85	.56	.53	.58	.81	1.20	1.55	12.49
Ruth	.75	1.30	.82	1.25	.77	1.24	.50	.88	.76	.45	1.01	1.02	10.76
Schellbourne Pass	.98	1.10	1.30	1.39	1.44	.64	.47	.88	.73	.56	.73	.97	11.20

1. See table 3 and figure 2 for station locations and period of record.

Surface Water

By. D. O. Moore

General Conditions

Northern and southern Butte Valleys are topographically closed basins and have no well-defined axial stream channels on their valley floors. Most of the runoff derived in the mountain blocks drain to the valley floors in ephemeral channels. Only five significant perennial streams occur in the area: Snow, Taylor Canyon, and Paris Creeks on the west flank of the Cherry Creek Range, and two small unnamed creeks, one each in northern and southern Butte Valleys, are springfed during low-flow periods. The unnamed creek of northern Butte Valley collects springflow in the western part of T. 27 N., R. 62 E., and that of southern Butte Valley in the southwestern part of T. 19 N., R. 62 E., also probably flows perennially from springflow.

In addition to streamflow from the mountain blocks, occasional flow may occur locally on the alluvial fans and lowlands in response to heavy precipitation from thunderstorms. Generally, this type of streamflow is so erratic in frequency and duration that it has little value to economic development.

No records of streamflow have been obtained on the streams in Butte Valley. For the purpose of this study, miscellaneous discharge measurements were made on Snow, Taylor Canyon, and Paris Creeks. Estimates of flow were also made for the two unnamed creeks. The data obtained are shown in table 5.

Streamflow in Butte Valley is variable with respect to the time of the year and from year to year. As there is no record of gaged streamflow in the area, records of a nearby gaged stream were used to show the variability of streamflow. This stream, Overland Creek near Ruby Valley (not shown on pl. 1), which is to the northwest of Butte Valley, is assumed to have the same general flow characteristics as the streamflow in Butte Valley. The graph in figure 3 shows the annual streamflow pattern for Overland Creek at the gaging station. The monthly plot points are shown as a percentage of the average annual streamflow. The middle lines of the graphs represent the median distribution of the monthly mean discharge for each month; that is, for each month, 50 percent of the monthly flows of record were less than and 50 percent were more than

Table 5.--Instantaneous discharge of perennial streams

Stream data point number on plate 1	Creek	Date	Discharge (cfs)
1	Unnamed Creek (T. 27 N., R. 62 E.)	8-19-67	a 2
2	Taylor Canyon Creek	10- 5-65	1.04
3	Snow Creek	10- 5-65	1.46
4	Paris Creek	10- 5-65	1.77
5	Unnamed creek (T. 19 N., R. 62 E.)	8-15-67	a 0.1

a. Estimated.

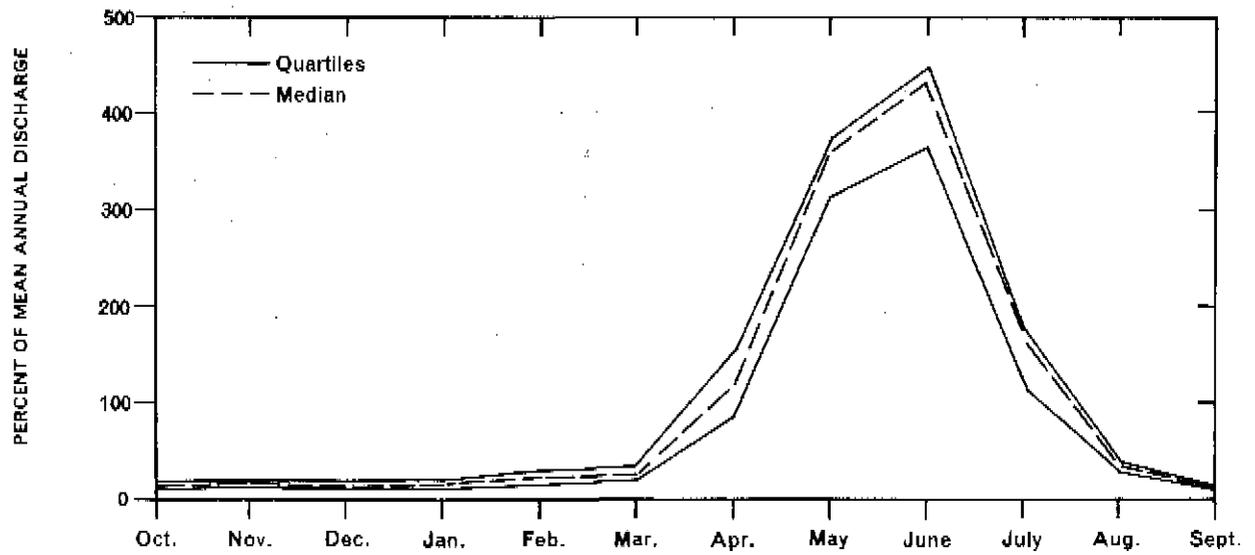


Figure 3.—Monthly discharge in percent of mean annual discharge of Overland Creek near Ruby Valley, Nevada (1960-65).

the proportional amount shown by the graph. The upper line of the graph, the upper quartile, represents a plot of the proportional monthly flow for which only 25 percent of the monthly flows of record were higher than and 75 percent were less than the proportions indicated. The lower line of the graph, the lower quartile represents a plot of the monthly proportion of annual flow for which 75 percent of the monthly flows of record were greater than and 25 percent were less than the proportion indicated. It can be noted from figure 3 that the annual hydrograph of flow for Overland Creek reflects a dominant influence of snowmelt runoff. This is also assumed true for streams in Butte Valley.

Estimated Average Annual Runoff

The amount of runoff that reaches the alluvial fans from the mountain blocks cannot be computed directly because of a scarcity of available streamflow data. Methods have been devised recently to estimate runoff in Nevada, particularly for application to areas where few or no streamflow records are available. These methods were described in detail by Moore (1968). Using the drainage areas supplying the natural flow to streams where gaging stations had been or are operated, recorded runoff is prorated by altitude zones (1,000-foot intervals) with due regard to the proportional areas of the several zones and with increasing unit values of runoff for increasing altitudes. As the different physical characteristics such as vegetation, geology, types of soil, and amounts of precipitation vary locally within large areas, the runoff for each altitude increment is adjusted accordingly. The adjustment of the runoff coefficients for local conditions are based on measurements of streamflow and channel geometry.

The estimated average annual runoff in Butte Valley, summarized in table 6, totals about 12,000 acre-feet per year. The runoff is not evenly distributed throughout the valley. It is estimated that about 80 percent occurs in the mountains on the eastern side and the remainder on the western side. On the eastern side, that part of the Cherry Creek Range north of the road between Cherry Creek and Butte Valley, though comprising less than 30 percent of the total runoff area, yields about 50 percent of the total runoff.

Ground-Water Recharge

Ground-water recharge in the Butte Valley area is derived mainly from precipitation within the area. A method described by Eakin and others (1951, p. 79-81) is used to estimate recharge in this report. The method assumes that a percentage

Table 3.--Estimated average annual runoff.

Mountain segment	Location	Runoff area (acres)	Estimated runoff (acre-feet per year)
NORTHERN BUTTE VALLEY			
Cherry Creek Range	North and west flanks of mountains above 7,000 feet	18,000	1,900
Spruce Mountain	Southwest flank of mountain above 7,000 feet	4,600	600
Medicine Range and Valley Mountain	Northeastern flank of Medicine Range and southeastern flank of Valley Mountain above 7,000 feet	5,400	100
Subtotal (rounded)		28,000	2,700
SOUTHERN BUTTE VALLEY			
Cherry Creek Range	West flank of mountains above 7,000 feet north of road connecting Butte Valley and Cherry Creek	25,000	4,000
Cherry Creek Range	West flank of mountains above 7,000 feet south of road connecting Butte Valley and Cherry Creek	20,000	1,700
Egan Range	West flank of mountains above 7,000 feet	27,000	1,600
Butte Mountains and Medicine Range	East flank of mountains and southeast flank of range above 7,000 feet	47,000	2,100
Subtotal (rounded)		127,000	9,400

of the average annual precipitation recharges the ground-water reservoir. Hardman (1965) showed that in gross aspect the average annual precipitation in Nevada is related closely to altitude and that it can be estimated with a reasonable degree of accuracy by assigning precipitation rates to various altitude zones. Estimates of recharge are shown in table 7. The estimated average annual precipitation on Butte Valley is about 560,000 acre-feet, and the estimated average annual recharge is about 19,000 acre-feet. Thus, about 3.4 percent of the total precipitation is computed to reach the ground-water reservoir.

Much of the recharge probably occurs by seepage loss as the streams cross the alluvial fans; however, the estimated mean annual runoff at the valley fill-consolidated rock contact of 12,000 acre-feet is considerably less than the estimated average annual recharge of 19,000 acre-feet. Much of the recharge reaching the valley floor may occur in the mountains by infiltration of precipitation and runoff into the carbonate rocks. The highly transmissive and structurally deformed character of the carbonate rocks can strongly affect the magnitude and direction of ground-water flow through these rocks. Therefore, the recharge boundaries, arbitrarily chosen to be coincident to surficial drainage boundaries for the compilation of table 7, may not be correct. However, because of the reconnaissance nature of the study and the lack of conclusive data that would permit a more accurate determination of recharge boundary locations, the surficial drainage boundaries were utilized for computation purposes. Some of the recharge in the mountains of southern Butte Valley may actually be moving as underflow through the carbonate rocks to northern Butte Valley or to adjacent valleys.

Table 7.--Estimated average annual precipitation and ground-water recharge

Altitude zone (feet)	Estimated annual precipitation			Estimated recharge	
	Area (acres)	Range (inches)	Average (feet)	Assumed percentage of precipitation	(acre-feet per year)
<u>NORTHERN BUTTE VALLEY</u>					
Above 9,000	1,100	20+	1.8	25	500
8,000-9,000	6,900	15-20	1.5	15	1,500
7,000-8,000	19,500	12-15	1.1	7	1,500
6,000-7,000	141,000	8-12	.8	3	360
Below 6,000	2,000	<8	.4	--	--
Subtotal (rounded)	170,000				3,900
<u>SOUTHERN BUTTE VALLEY</u>					
Above 9,000	3,200	20+	1.8	25	1,400
8,000-9,000	21,600	15-20	1.5	15	4,800
7,000-8,000	95,100	12-15	1.1	7	7,000
Below 7,000	345,000	<12	.8	3	1,700
Subtotal (rounded)	465,000				15,000
Total (rounded)	635,000				19,000

a. Most of this area is underlain by alluvium where the estimated 8 to 12 inches of precipitation probably supplies negligible recharge. About 15,000 acres in northern Butte Valley and about 70,000 acres in southern Butte Valley are assumed to be effective for estimating recharge by this method.

OUTFLOW FROM THE VALLEY-FILL RESERVOIR

Irrigation Use

Farming in the valley is mainly restricted to a small amount of irrigation of alfalfa and native pasture. Growing-season length is one of the main factors that has influenced agricultural development of the valley. Temperature data, useful in determining and predicting the growing season length, are unavailable for Butte Valley proper.

Temperature data are recorded at a few nearby stations outside the report area, as shown in figure 2. Since 1949, the U.S. Weather Bureau has published freeze data for many of their temperature-recording stations; these data for five stations in adjacent areas are summarized in table 8. Because killing frosts vary with the type of crop, temperatures of 32°F, 28°F, and 24°F are used to determine the number of days between the last spring minimum (prior to July 1) and the first fall minimum (after July 1). The temperature stations at Ely Airport and Ruby Lake are on or near the floors of their respective valleys, and their altitudes are similar to those of the actual and potential agricultural areas of Butte Valley. However, their respective growing seasons are somewhat dissimilar (table 8). The length of growing season on the floor of Butte Valley may be between those of the Ely Airport and Ruby Lake temperature stations. Thus, the growing season for crops experiencing a killing frost of 28°F would probably average between 104 and 130 days, based on data collected between 1950 and 1966 (table 8). In any event, a somewhat shorter growing season would be expected on the valley floor than on the alluvial slopes because of the tendency for cooler air to accumulate at lower altitudes in most closed or partially closed Nevada valleys.

Irrigation of alfalfa and native pasture has been limited in part to diversion of perennial streams and springs and in part by natural subirrigation in areas of shallow ground water. The springfed streams and the shallow depth of ground water in irrigated areas complicates a clear-cut separation of surface-water and ground-water sources. Nevertheless, rough estimates have been made as to source of supply. Table 9 shows the estimated consumptive use of stream, spring, and shallow ground water in the valley.

Evapotranspiration

Evapotranspiration from the valley-fill reservoir is greatest in the northern half of southern Butte Valley and in the southern half of northern Butte Valley (pl. 1). These areas are not at the lowest altitudes of their respective valley segments; rather, they

Table 3.--Length of growing season between killing frosts

that occur at 32°F, 28°F, and 24°F

[Summarized from published records of the U.S. Weather Bureau]

Station/ :(feet)	Period of record :(years)	Minimum recorded :(days)			Maximum recorded :(days)			Average :(days)		
		32°F :28°F	24°F	32°F	28°F	24°F	32°F			
Ely Airport 6,257	1950-66	7	63	96	125	153	188	77	104	141
Kimberly 7,250	1950-57	7	31	114	119	197	207	90	124	151
McGill 6,340	1950-66	50	93	130	150	204	216	118	147	170
Ruby Lake 6,012	1950-66	63	91	133	142	170	189	105	130	155
Ruth 6,032	1950-66	10	20	28	90	150	167	30	85	97

1. See table 3 and figure 2 for location.

Table 9.--Estimated consumption of water by

irrigated and subirrigated crops

Ranch	Crop	Approximate irrigated area (acres)	Estimated rate (feet per year)	Estimated consumptive use (acre-feet per year)		
				Diversion of streams and springs	Shallow ground water ¹	Total
<u>NORTHERN BUTTE VALLEY</u>						
Healy Ranch	Alfalfa	100	a 2	b 200	--	200
Do.	Do.	60	a 2	120	--	120
Do.	Native pasture	640	1	b 320	320	640
Ranch (owner unknown)	Do.	140	1	b 70	70	140
Total		940		710	390	1,100
<u>SOUTHERN BUTTE VALLEY</u>						
Pazis Ranch	Alfalfa	150	a 2	300	--	300
	Crested wheat grass	30	1	30	--	30
Stratton Ranch	Alfalfa	60	a 2	b 60	c 60	120
	Native pasture	780	1	b 50	c 730	780
Total		1,020		440	790	1,200

1. Depth to water generally less than 5 feet.

a. Alfalfa use estimated at 1 foot per cutting; alfalfa reportedly yields two cuttings per year.

b. Supply mostly from springs.

c. Supplemented by diversion of Snow Creek when it is flowing.

straddle the topographic divide between the two areas and are the places where most spring discharge and the shallowest depths to ground water generally occur.

Evapotranspiration occurs chiefly by phreatophyte consumption and some evaporation from free water surfaces. Greasewood, rabbitbrush, and native meadowgrass are the most prevalent phreatophytes, areally; however, cottonwood trees, wild rose, willows, and rushes occasionally occur around springs and perennial stream channels. A few small storage reservoirs, ponds, and perennial springfed streams lose water by evaporation. Average annual evaporation rate of free water surfaces in the valley is about 45 inches (Kohler and others, 1959, pl. 2). The estimated evapotranspiration of ground water in nonirrigated areas is shown in table 10.

Springs

Springs discharge from both the valley fill and the consolidated rocks. A summary of springs visited during this investigation is shown in table 11. Those discharging from the valley fill are generally north of Paris Creek in southern Butte Valley and south of West Buttes in northern Butte Valley, as shown on plate 1. Their discharges range from seeps of less than 1 gpm (gallon per minute) to about 400 gpm (spring 28/61-11d).

Springs also discharge from both the carbonate and noncarbonate rocks. Those visited during the course of this investigation are shown on plate 1. Noncarbonate-rock springs visited have discharges ranging from less than 5 gpm to about 50 gpm. Most of the discharge is dissipated by evapotranspiration or infiltration near the discharge area. However, numerous small springs occur in the southwestern part of T. 19 N., R. 62 E., and maintained a small, generally perennial stream; this stream was observed near its lower extremity, and its flow was estimated at 0.1 cfs (table 5 and pl. 1). A number of springs issue from the alluvium in the northwestern part of T. 27 N., R. 62 E. An estimate of about 2 cfs (cubic feet per second) was made in the unnamed creek that drains their composite flow (table 5 and pl. 1). The larger springs in the carbonate rocks probably are those that maintain the low flows of Paris, Taylor Canyon, and Snow Creeks.

Because springflow is diverted to crop use (table 9) or is accounted for as phreatophyte discharge (table 10), no attempt was made to measure the discharge of many of the springs.

Table 10.--Estimated average annual evapotranspiration of
ground water in nonirrigated areas

Phreatophyte or type of discharge	Approximate area (acres)	Depth to water range (feet)	Estimated water use rate (feet per year)	Estimated ground- water discharge (acre-feet per year)
<u>NORTHERN BUTTE VALLEY</u>				
Evaporation from ponds, reser- voirs, and streams	10	--	3.5	35
Mostly meadow grass	1,400	2-5	0.75	1,000
Do.	4,000	0-10	0.5	2,000
Rabbitbrush and greasewood	13,000	10-35	0.2	2,600
Greasewood	13,000	30-50+	0.1	1,300
Total (rounded)	31,400			6,900
<u>SOUTHERN BUTTE VALLEY</u>				
Evaporation from ponds, reser- voirs, and streams	10	--	3.5	35
Saltgrass	9,400	0-10	0.5	4,700
Rabbitbrush and greasewood	22,000	10-35	0.2	4,400
Greasewood	14,000	30-50+	0.1	1,400
Total (rounded)	45,400			11,000

Table 11.--Records of selected springs

Use: D, domestic; I, irrigation;
S, stock; U, unused

Spring number	Name and (or) owner	Estimated discharge (gpm)	Use	Temperature (°F)	Temperature (°C)
19/60-4b1	Uhalde Ranch	10	S	--	--
19/62-9c1	Gulch Spring	10-20	U	--	--
19/62-32c1	Summit Springs	<25	U	--	--
20/60-33d1	Thirty mile Spring - Uhalde Ranch	40-50	D,I,S	43	9
22/62-21d1	Nine Mile Spring	<10	S	51	11
26/62-15c1	Stratton Spring - Paris Ranch	250	I,S	57	14
26/62-33d1	Owens Springs - Paris Ranch	50-100	I,S	50	10
27/62-33c1	Twin Springs	10-25	S	57	14
28/61-2d1	Quilici Spring	<50	S	52	14
28/61-11d1	Healy Ranch	400	I,S	--	--
28/61-26d1	--	<5	S	56	13
28/62-9c1	--	5	S	57	14
29/62-23d1	--	350+	I,S	--	--

1. Chemical analyses available shown in table 13 for all springs except

19/60-4b1, 19/62-9c1, and 19/62-32c1.

Subsurface Outflow to Adjacent Valleys

Ground water moving northward in northern Butte Valley is consumed principally by evapotranspiration; the residual flows westward to Ruby Valley mainly through older alluvium. Some outflow may also occur through the underlying carbonate rocks. Underflow can be computed by means of a form of Darcy's law

$$Q = 0.00112 \cdot T \cdot I \cdot W$$

in which Q is the quantity of flow, in acre-feet per year; T is the coefficient of transmissibility, in gallons per day per foot; I is the hydraulic gradient, in feet per mile; W is the effective width of the underflow section, in miles; and 0.00112 is a factor for converting gallons per day to acre-feet per year.

The effective width along the alluvial divide between northern Butte Valley and Ruby Valley probably is about 2 miles. The coefficient of transmissibility of the older alluvium is assumed to be 50,000 gallons per day per foot. The hydraulic gradient, based on wells 30/62-18c1 and 30/61-7d1 (water-level altitudes 5,935 and 5,895 feet, respectively), is 40 feet in 5.5 miles, or about 7 feet per mile. The estimated outflow, using the above equation and foregoing values, is computed to be about 800 acre-feet per year.

Pumpage

All wells drilled in the report area are used for domestic or stockwatering purposes. Only about 16 wells were known to be in use in 1967; these are listed in table 15. In northern Butte Valley, 200 head of cattle each using about 15 gpd (gallons of water per day) would consume about 3 acre-feet per year; 5 people, using about 100 gpd, would consume about 0.5 acre-foot per year; total domestic and stock consumption in that valley probably is less than 10 acre-feet per year. In southern Butte Valley, 7,000 sheep each using about 2 gpd would consume about 15 acre-feet per year; 500 cattle, about 7 acre-feet; and 25 people, about 2.5 acre-feet; total domestic and stock consumption in that valley probably is about 25 acre-feet per year. Therefore, total domestic and stock consumption in Butte Valley probably is less than 35 acre-feet per year.

GROUND-WATER BUDGET FOR NEAR-NATURAL CONDITIONS

Over the long term and for natural conditions, recharge to and discharge from an area are equal, provided that (1) the ground-water system is in a state of natural equilibrium, and (2) the long-term climatic regimen remains nearly constant. A ground-water budget for the state of nature expresses the total quantity of water flowing in the system under equilibrium conditions. Accordingly, the purpose of preparing a budget is to compare the estimates of recharge and discharge for each valley, to determine the magnitude of the difference between estimates, and to select a reconnaissance value that may reasonably represent both recharge and discharge.

Because of the small development in Butte Valley, near-natural equilibrium conditions are assumed to prevail. Table 12 summarizes the several estimates of recharge and discharge made in the preceding sections of the report, shows the water balances achieved, and the reconnaissance value selected for both recharge and discharge.

The table also shows that in northern and southern Butte Valley the imbalances are 4,800 and 3,000 acre-feet, respectively, which are significant parts of the total quantities involved. However, for Butte Valley as a whole, the imbalance is only about 10 percent of the total recharge or discharge. The imbalances could be caused by regional intervalley flow through the carbonate rocks, or by errors in the assumptions made in deriving the estimates of recharge and discharge. Because of these uncertainties, the reconnaissance values selected for both recharge and discharge are taken as the average of the two values.

Table 12.--Preliminary ground-water budget for near-natural conditions

(All estimates in acre-feet per year)

Budget elements	Northern Butte Valley (1)	Southern Butte Valley (2)	Butte Valley total (1)+(2)
<u>RECHARGE:</u>			
From precipitation (table 7)	3,900	15,000	19,000
Subsurface inflow	--	--	--
Total (rounded): (1)	3,900	15,000	19,000
<u>DISCHARGE:</u>			
Irrigation use (table 9) ^{1/}	980	900	1,900
Evapotranspiration (table 10) ^{2/}	6,900	11,000	18,000
Subsurface outflow (p. 30)	300	--	800+
Pumpage (p. 30)	10	25	35
Total (rounded): (2)	8,700	12,000	21,000
<u>IMBALANCE:</u> (1) - (2)	-4,800	3,000	2,000
<u>RECONNAISSANCE VALUE SELECTED</u>			
<u>FOR RECHARGE AND DISCHARGE</u>	6,300	14,000	20,000

1. Includes estimated ground-water discharge only from springs and shallow water table; excludes stream diversions.

2. Evapotranspiration in nonirrigated areas includes discharge from springs and shallow water table.

CHEMICAL QUALITY OF WATER

General Characteristics

Analytical results for 21 water samples collected during the study are shown in table 13. The water analyses for this report are somewhat areally biased, because water-quality sampling was generally restricted to local areas of development. Moreover, most water sampled was from comparatively shallow aquifers; however, deeper aquifers or other locations may contain water of considerably different quality.

The dissolved-solids content of water is indicated in a general way by the measurement of specific conductance. The dissolved-solids content, in ppm (parts per million), commonly equals about two-thirds the specific conductance (micromhos per centimeter at 25°C; abbreviated, micromhos). Specific conductance of sampled water in the area ranges from 200 to 629 micromhos; therefore, dissolved-solids content probably ranges from about 130 to more than 400 mg/l (milligrams per liter, which is about equal to parts per million; see footnote 1, table 13).

The analytical results of table 13 permit the following generalizations about the chemical quality of water in the area: with regard to the dissolved-solids content, water quality is exceptionally good compared to many areas of Nevada; variability in overall water quality apparently is not systematically related to geographical location; however, the analysis of water from well 21/61-6c1 suggests by the higher dissolved-solids, sodium, and chloride content that a slight degradation of quality may be occurring near the topographically low part of southern Butte Valley. As expected, the general quality of surface water is somewhat better than the general quality of spring and well water. Water sampled ranges from moderately hard to very hard and chemically does not appear to be affected by its proximity to areas of heaviest evapotranspiration.

Unverified reports of deterioration of ground-water quality associated with high pumping rates and (or) prolonged periods of pumping were filed with the Bureau of Land Management by local ranchers in 1959 and 1960. These reports are part of records involving a decision in 1960 regarding applications for homesteads under the Desert Land Act. However, this present reconnaissance did not allow testing to determine the effects of high pumping rates and prolonged pumping on water quality.

Table 13.—Partial chemical analyses of water from wells, springs, and streams

[Field-office analyses by the U.S. Geological Survey]

Location	Date sampled	Tem- per- ature °F °C	Milligrams per liter (upper number) and milliequivalents per liter (lower number) ^{1/}								Specific conduct- ance (micro- mhos per cm at 25°C)	pH (lab. deter- mina- tion)	Factors affecting suitability for irrigation ^{2/}		
			Cal- cium (Ca)	Mag- ne- sium (Mg)	Sod- ium (Na) (K) ^{3/}	Chlo- ride (Cl)	Sul- fate (SO ₄) ^{4/}	Bicar- bonate (HCO ₃) ^{5/}	Hard- ness as CaCO ₃	Salinity hazard			Sodium adsorp- tion ratio (SAR)	Sodium hazard	Residual sodium carbonate (RSC)
WELLS															
21/61-6c1	9-22-65	-- --	45	28	53	122	47	138	279	629	8.0	Medium	1.5	Low	8(0.00)
			2.25	2.33	2.29	2.00	0.98	3.09	4.58						
22/61-6c1	10- 5-65	48 9	28	18	15	154	32	11	143	298	8.2	Medium	.5	Low	8(0.00)
			1.40	1.46	0.64	2.52	0.67	0.31	2.86						
23/61-7d1	9-22-65	47 8	25	20	37	201	28	11	144	373	8.4	Medium	1.3	Low	8(.71)
			1.25	1.63	1.60	3.29	0.58	0.31	2.88						
24/61-14c1	9-22-65	56 13	37	29	32	159	64	58	210	534	8.1	Medium	1.0	Low	8(0.00)
			1.85	2.35	1.38	2.61	1.33	1.64	4.20						
25/62-17b1	8-10-67	54 12	51	18	12	240	20	7.9	203	410	8.0	Medium	.4	Low	8(0.00)
			2.54	1.52	0.51	3.93	0.42	0.22	4.06						
26/62-22a1	8-13-67	-- --	44	18	8.0	222	9.6	4.2	184	350	8.3	Medium	.3	Low	8(.03)
			2.20	1.48	0.35	3.64	0.20	0.12	3.68						
SPRINGS															
20/60-33d1	8-15-67	48 9	16	5.1	16	174	8.7	6.9	86	230	7.7	Low	.7	Low	8(.31)
			1.30	0.42	0.68	2.03	0.18	0.19	1.72						
22/62-21d1	8-21-67	51 10	58	6.4	22	210	24	16	171	420	7.6	Medium	.7	Low	8(.02)
			2.89	0.53	0.97	3.44	0.50	0.45	3.42						
26/22-15a1	8-18-67	57 14	40	19	7.4	208	14	6.5	178	350	8.0	Medium	.2	Low	8(0.00)
			2.00	1.56	0.32	3.41	0.29	0.18	3.56						
26/62-33d1	8-18-67	50 10	42	19	9.0	220	12	6.2	182	350	8.0	Medium	.3	Low	8(0.00)
			2.10	1.54	0.39	3.61	0.25	0.17	3.66						
a 27/62-23c1	8-19-67	-- --	44	21	3.0	222	13	6.9	196	360	8.2	Medium	.1	Low	8(0.00)
			2.20	1.72	0.13	3.64	0.27	0.16	3.92						
a 28/61-2d1	8-19-67	58 14	42	17	10	199	21	7.2	173	350	7.9	Medium	.3	Low	8(0.00)
			2.10	1.36	0.44	3.26	0.44	0.20	3.46						
a 28/61-11d1	8-19-67	-- --	37	16	11	183	19	7.9	157	330	8.0	Medium	.4	Low	8(0.00)
			1.85	1.29	0.48	3.00	0.40	0.22	3.14						
a 28/61-26d1	8-19-67	56 13	39	19	11	195	26	10	177	360	8.1	Medium	.4	Low	8(0.00)
			1.95	1.59	0.48	3.20	0.54	0.28	3.54						
a 28/62-0a1	8-20-67	-- --	63	27	21	332	30	8.6	269	590	7.8	Medium	.6	Low	8(.07)
			3.14	2.73	0.93	5.44	0.62	0.24	5.37						
a 29/62-23d1	8-19-67	67 19	62	28	21	288	62	10	270	540	8.0	Medium	.5	Low	8(0.00)
			3.09	2.30	0.90	4.72	1.29	0.28	5.39						
STREAMS															
Unnamed creek	8-15-67	65 18	39	7.9	24	178	19	17	130	340	7.9	Medium	.9	Low	8(.32)
19/62-30b1			1.95	0.65	1.06	2.92	0.40	0.34	2.60						
Paris Creek	10- 5-65	50 10	21	23	12	158	24	5.4	147	269	8.4	Medium	.4	Low	8(0.00)
25/62-21			1.05	1.89	0.53	2.59	0.50	0.19	2.84						
Snow Creek	10- 5-65	51 10	27	10	8.7	126	18	4.8	110	200	8.1	Low	.4	Low	8(0.00)
26/62-35			1.35	0.85	0.38	2.07	0.37	0.14	2.20						
Unnamed creek	8-19-67	70 21	24	20	8.5	171	13	4.2	141	290	8.2	Medium	.3	Low	8(0.00)
a 27/62-8a1			1.20	1.62	0.37	2.80	0.27	0.12	2.82						
Taylor Creek	10- 5-65	54 12	25	15	13	146	24	6.8	125	230	8.2	Low	.5	Low	8(0.00)
a 27/62-12			1.25	1.25	0.58	2.39	0.50	0.19	2.50						

1. Milligrams per liter and milliequivalents per liter are metric units of measure that are virtually identical to parts per million and equivalents per million, respectively, for all waters having a specific conductance less than about 10,000 micromhos. The metric system of measurement is receiving increased use throughout the United States because of its value as an international form of scientific communication. Therefore, the U.S. Geological Survey recently has adopted the system for reporting all water-quality data.
2. **Salinity hazard** is based on specific conductance (in micromhos) as follows: low, 0-250; medium, 251-750; high, 751-2,250; very high, >2,250. **Sodium-adsorption ratio (SAR)** provides an indication of what effect an irrigation water will have on soil-drainage characteristics. SAR is calculated as follows, using milliequivalents per liter: $SAR = Na / \sqrt{(Ca + Mg)/2}$. **Sodium hazard** is based on an empirical relation between salinity hazard and sodium-adsorption ratio. **Residual sodium carbonate** (expressed in milliequivalents per liter) is tentatively related to suitability for irrigation as follows: safe (S), 0-1.25; marginal (M), 1.26-2.50; unsuitable (U), >2.50. The several factors should be used as general indicators only, because the suitability of a water for irrigation also depends on climate, type of soil, drainage characteristics, plant type, and amount of water applied. These and other aspects of water quality for irrigation are discussed by the U.S. Salinity Laboratory Staff (1954).
3. Computed as the milliequivalent-per-liter difference between the determined negative and positive ions, expressed as sodium. Computation assumes that concentrations of undetermined ions—especially nitrate—are small.
4. All carbonate (CO₃) values 0 mg/l except: 23/61-7d1, 9 mg/l (0.30 meq/l); 26/62-22a1, 2 mg/l (0.07 meq/l); 25/62-21, 7 mg/l (0.23 meq/l).
5. In northern Butte Valley. ⁷⁵ Not footnoted, in southern Butte Valley.

Suitability of Water for Various Uses

Factors affecting the suitability of Butte Valley waters for agriculture are shown in table 13. The factors are based on consideration of several of the chemical characteristics that were determined for the sampled waters. The presence of boron, which was not determined in the analyses, is also an important factor that determines suitability of water for agriculture. Boron, which is necessary in small quantities for healthy plant growth, is toxic to plants when present in water in quantities only slightly exceeding the desirable amount.

Increased agricultural development in the study area may ultimately be influenced as much by the chemical quality as by the quantity of available water. Evaluation of prospective agricultural development warrants careful consideration of the chemical quality of the water and the chemical and physical character of the lowland soils. This would help ensure compatibility of soil and irrigation water with the type of crops planned.

Future recycling of ground water for intensive irrigation around the lowland areas probably would degrade the chemical quality of the water, particularly if commercial fertilizers are used in substantial quantities and if the soils contain appreciable quantities of leachable salts.

Drinking-water standards recommended by the U.S. Public Health Service (1962) commonly are cited as limits for domestic use. Several of these standards are as follows:

<u>Chemical constituent</u>	<u>Recommended maximum concentration (mg/l)</u>
Sulfate	250
Chloride	250
Total dissolved solids	500

As indicated by table 13 and the above tabulation, all waters analyzed are acceptable for human consumption, based on their sulfate, chloride, and total dissolved-solids content.

Hardness of water, which is mainly caused by calcium and magnesium, adversely affects suitability of water for domestic use, especially for cooking and washing, and may also be detrimental to certain industrial uses. The U.S. Geological Survey uses the following classification of water hardness:

<u>Hardness range</u> (mg/l)	<u>Classification</u>
0-60	Soft
61-120	Moderately hard
121-180	Hard
Greater than 180	Very hard

Water analyzed from the study area generally ranges from moderately hard to very hard.

Not all developed water in the area was analyzed, and the analytical procedures employed were not intended to determine the presence of many elements known to be toxic or undesirable in excessive concentrations; also, bacteriological tests were not performed on any samples. Therefore, to assure that the water is safe and acceptable for human consumption, questionable water should be submitted to a reliable laboratory or to the Nevada Bureau of Environmental Health for analysis when domestic or municipal development is planned for a particular water supply.

The suitability of water for industrial use depends on the quality requirements of the industry. Water considered unacceptable for internal consumption by humans and animals might be desirable for industrial use.

THE AVAILABLE WATER SUPPLY

The available ground-water supplies of the two valleys of the area consists of two interrelated entities: (1) perennial yield, and (2) transitional storage reserve, which are described below.

Perennial Yield

Perennial yield of a ground-water reservoir may be defined as the maximum amount of water of usable chemical quality that can be withdrawn and consumed economically each year for an indefinite period of time. If the perennial yield is continually exceeded, water levels will decline until the ground-water reservoir is depleted of water of usable quality or until the pumping lifts become uneconomical to maintain. Perennial yield cannot exceed the natural recharge to an area and ultimately is limited to the maximum amount of natural discharge that can be salvaged for beneficial use. Salvage of natural discharge implies diversion of ground water presently destined for areas of natural discharge to areas of pumping. A method of accomplishing the diversion is the lowering of water levels in and near areas of natural discharge by scheduled depletion of storage according to the concept of transitional storage reserve (defined below).

Table 12 shows that the reconnaissance values selected for recharge and discharge for northern and southern Butte Valleys are 6,300 and 14,000 acre-feet per year, respectively. In northern Butte Valley, where about 980 acre-feet per year has been salvaged for irrigation use, most of the selected natural discharge probably could be salvaged for useful purposes. Thus, the estimated perennial yield is about 6,000 acre-feet.

In southern Butte Valley, where about 900 acre-feet per year has been salvaged for irrigation use, most of the selected natural discharge probably also could be salvaged for beneficial use. Therefore, the estimated perennial yield is about 14,000 acre-feet. The perennial yield for Butte Valley as a whole, then, is about 20,000 acre-feet.

When considering the location of wells or other development to salvage the natural discharge for beneficial use, table 14 shows the distribution and approximate percentage of natural evapotranspiration in the two valleys. The table does not include the 1967 irrigation use.

Table 14.--Areal distribution of ground-water discharge
by natural evapotranspiration

(Estimates from table 10).

<u>Location</u>	<u>Approximate percentage of discharge</u>
<u>NORTHERN BUTTE VALLEY</u>	
(Estimated discharge, 6,900 acre-feet per year)	
Tps. 27 and 28 N.	50
Tps. 29 and 30 N.	50
<u>SOUTHERN BUTTE VALLEY</u>	
(Estimated discharge, 11,000 acre-feet per year)	
Tps. 19 and 20 N.	77
Tps. 21 and 22 N.	15
Tps. 23 and 24 N.	25
Tps. 25 and 26 N.	60

Transitional Storage Reserve

Transitional storage reserve has been defined by Worts (1967) as the quantity of water in storage in a particular ground-water reservoir that can be extracted and beneficially used during the transition period between equilibrium conditions in a state of nature and new equilibrium conditions under the perennial yield concept of ground-water development. In the arid environment of the Great Basin, the transitional storage reserve of such a reservoir is the amount of stored water available for withdrawal by pumping during the nonequilibrium period of development or period of lowering water levels. Therefore, transitional storage reserve is a specific part of the water available from ground water in storage; it is a quantity additional to that of perennial yield, but can be withdrawn on a once-only basis.

Ground-water development inherently involves storage depletion; the magnitude of depletion is commensurate with the amount of pumpage, the hydraulic characteristics of the aquifer, and locations of wells with respect to recharge and discharge boundaries. Desert valleys often have well-defined discharge boundaries, such as areas of evapotranspiration, but recharge boundaries, such as live streams or lakes, are uncommon.

Computation of transitional storage reserve for valleys of the report area was based on the following assumptions: (1) Development wells would be strategically located in, near, and around the areas of natural discharge so that any subsurface outflow losses could be reduced and any evapotranspiration losses stopped with a minimum of water-level drawdown in the pumped wells; (2) in general, water levels would be lowered to and stabilized at a minimum depth of 50 feet below the land surface in areas of phreatophyte growth, which would curtail virtually all evapotranspiration losses from the ground-water reservoir; (3) long-term pumping would cause a moderately uniform depletion of storage throughout the valley-fill reservoir, except possibly in the very fine-grained playa deposits where transmissibility and storage coefficients are small and therefore storage depletion also would be small or occur over a very long period of time; (4) the specific yield of the valley fill is 10 percent; (5) water levels are within the range of economic pumping lift for the intended use; (6) the pumping development causes little or no effect on adjacent valleys and only small quantities of water are withdrawn from the adjacent consolidated-rock mountain masses; and (7) the water is of suitable chemical quality for the desired use.

Preliminary estimates of the transitional storage reserves of the valley-fill reservoirs are computed as follows: (1) northern Butte Valley--the estimated effective area (100,000 acres) times the assumed specific yield (10 percent) times the selected saturated thickness (30 feet), or about 300,000 acre-feet; and (2) southern Butte Valley--the estimated effective area (200,000 acres) times the assumed specific yield (10 percent) times the selected saturated thickness (30 feet), or about 600,000 acre-feet.

The manner in which transitional storage reserve augments the perennial yield has been described by Worts (1967) and in its simplified form is shown by the following equation:

$$Q = \frac{\text{Transitional storage reserve}}{t} + \frac{\text{Perennial yield}}{2}$$

in which Q is the pumping rate, in acre-feet per year, and t is the time, in years, to exhaust the storage reserve. This basic equation, of course, could be modified to allow for changing rates of storage depletion and salvage of natural discharge. The equation, however, is not valid for pumping rates less than the perennial yield.

Using the above equation and the estimates for northern Butte Valley as an example (transitional storage reserve 300,000 acre-feet; perennial yield 6,000 acre-feet) and using a pumping rate (Q) equal in quantity to the perennial yield, in accordance with the general intent of Nevada water law, the time (t) to deplete the transitional storage reserve is computed to be about 100 years. Similarly, depletion of the storage reserve in southern Butte Valley would take about 85 years. At the end of those respective times, the transitional storage reserves of the two valleys would be exhausted, subject to the assumptions previously described.

What is not shown by the example is that in the first year virtually all the pumpage would be supplied from storage, and very little, if any, would be derived by salvage of natural discharge. On the other hand, during the last year of the period nearly all pumpage would be derived from the salvage of natural discharge and virtually none from the storage reserve.

During the period of depletion the ground-water flow net would be substantially modified. The recharge that originally flowed from around the sides of the valleys to areas of natural discharge would ultimately flow directly to the pumping wells.

To meet the needs of an emergency or other special purpose requiring ground-water pumpage in excess of the perennial yield for specified periods of time, the transitional storage reserves would be depleted at a more rapid rate than in the example given. The above equation can be used to compute the time required to exhaust the storage reserve for any selected pumping rate in excess of the perennial yield. However, once the transitional storage reserve was exhausted, the pumping rate should be reduced to the perennial yield as soon thereafter as possible. Pumpage in excess of the perennial yield would result in an overdraft, and pumping lifts would continue to increase and stored water would continue to be depleted until some undesired result occurred.

NUMBERING SYSTEM FOR WELLS AND SPRINGS

Numbers assigned to wells and springs in this report are based on the rectangular subdivisions of the public lands referenced to the Mount Diablo base line and meridian. Each number consists of three units, the first is the township north of the base line. The second unit, separated from the first by a slant, is the range east of the meridian. The third unit, separated from the second by a dash, designates the section number, which in turn is followed by a letter that indicates the quarter section. The quarter sections are designated counterclockwise in sequence beginning with "a" for the northeast quarter section. Following the letter, a number indicates the order in which the well or spring was recorded within the 160-acre tract. For example, well 31/62-34b1 is the first well recorded in the NW $\frac{1}{4}$ sec. 34, T. 31 N., R. 62 E., Mount Diablo base line and meridian.

Because of limitation of space, wells and springs are identified on plate 1 only by the section number, quarter-section letter, and the number indicating the order in which the well or spring was located. Township and range numbers are shown along the margins of the plate.

Table 15. Records of selected wells

Owner or name: BLM, U.S. Bureau of Land Management
 Use: D, domestic S, stock U, unused
 State log number: Log number in the files of the
 Nevada State Engineer

Well number	Owner or name	Year drilled	Depth (feet)	Diameter (inches)	Use	Yield (gpm) and drawdown (feet)	Altitude of water level measurement		Date	Depth to water table (feet) below land surface)	State log number
							(feet above mean sea level)	(feet below surface)			
19/61-30b1	BLM	1966	270	6	S	---	6,800	199.10	8-15-67	199.10	9030
20/61-6d1	Uthalde White Sage well	---	---	6	S	---	6,295	150.92	8-15-67	150.92	---
20/61-14d1	Gulf Oil Corp.	1965	105	6	U	---	6,240	a 65	12--65	a 65	8020 8804
21/61-6c1	B. Paris	---	---	6	S	---	6,170	66.75	8-17-67	66.75	---
22/60-26a1	B. Paris	1925+	---	6	S	---	6,190	65.16	8-17-67	65.16	---
22/61-6c1	B. Paris	---	185	6-8	S	---	6,190	38.54	6-26-58	38.54	---
22/61-15	B. Paris	---	36	---	U	---	7,400	32.0	6-30-58	32.0	---
22/61-33	B. Paris	---	12.3	---	S	---	6,690	10.5	7-9-58	10.5	---
23/60-22b1	BLM - West Butte well	---	105	6	S	25/---	6,255	54.81	8-17-67	54.81	9242
23/61-7d1	B. Paris	---	40	6-8	S	---	6,280	27.36	8-16-67	27.36	---
23/61-13	B. Paris	---	10	---	U	---	7,615	a 10	---	a 10	---
24/60-33b1	BLM - Marren Robison well	1966	420	---	U	Dry hole	---	---	8--66	---	9258
24/61-14c1	B. Paris	---	---	4-6	S	---	6,320	53.65	8-16-67	53.65	---
25/62-17b1	B. Paris	---	---	---	S	---	6,300	9.15	8-16-67	9.15	---
26/62-22a1	B. Paris	---	---	---	D,S	---	6,390	15.25	8-16-67	15.25	---
29/62-23b1	Don Christianson	1950	54	6	D	200+/-	6,190	43.10	8-19-67	43.10	1279

Table 15.---Continued

Well number	Owner or name	Year drilled	Depth (feet)	Diameter (inches)	Use	Yield (gpm) and drawdown (feet)	Altitude of land surface		Date	Depth to water Nevada State log number	
							(feet above mean sea level)	(feet above surface)		table (feet below land surface)	number
30/62-18c1	Robison and Sorenson	1949	110	6	S	--	6,010	75.40	8-20-67	1171	--
30/62-33c1	---	---	---	6	S	---	6,090	35.49	8-20-67	---	---
30/63-31b1	Lloyd Sorenson	1962	124	6	S	30/---	6,080	45.21	8-20-67	6515	---
31/62-34a1	RLM - Liza Jane well	---	---	6	S	---	6,105	109.70	8-20-67	---	---

a. Reported.

b. Not visited during this study; data from Snyder (1963).

Table 16. -- Available drillers' logs of wells

Material	Thick- ness (feet)	Depth (feet)	Material	Thick- ness (feet)	Depth (feet)
<u>19/61-30b1</u>			<u>29/62-23b1</u>		
Clay, brown	50	50	Gravel and soil	4	4
Rock, volcanic	40	90	Rock and yellow clay	16	20
Sand, gray	20	110	Rock, loose, and dry gravel	4	24
Sand and clay	20	130	Rock, lime	12	36
Gravel, water-bearing	140	270	Gravel, loose	2	38
<u>20/61-14d1</u>			Quartz	12	50
Topsoil	12	12	Gravel, water-bearing	2	52
Sand and gravel	4	16	Quartz	2	54
Clay	34	50	<u>30/62-18c1</u>		
Cemented	5	55	Gravel, cemented	90	90
Clay	30	85	Clay, sandy	5	95
Gravel, water-bearing	20	105	Sand and gravel, water-bearing	15	110
<u>23/60-22b1</u>			<u>30/63-31b1</u>		
Unknown	35	35	Clay, yellow	42	42
Sand	5	90	Gravel, sand, and clay	23	65
Gravel, water-bearing	15	105	Gravel and sand, water-bearing	59	124
<u>24/60-33b1 (dry hole)</u>					
Clay and gravel	40	40			
Limestone	380	420			

REFERENCES CITED

- Adair, D. H., and Stringham, B. F., 1960, Intrusive igneous rocks of east-central Nevada, in Geology of east-central Nevada: Intermountain Assoc. Petroleum Geologists, 11th Ann. Field Conf., 1960 Guidebook, p. 229-231.
- Chapman, Arthur, 1932, the Pony Express: New York-Chicago, A. L. Burt Co., 319 p.
- Douglas, W. B., Jr., 1960, Geology of the Southern Butte Mountains, White Pine County, Nevada, in Geology of east-central Nevada: Intermountain Assoc. Petroleum Geologists, 11th Ann. Field Conf., 1960 Guidebook, p. 181-185.
- Eakin, T. E., and others, 1951, Contributions to the hydrology of eastern Nevada: Nevada State Engineer Water Resources Bull. 12, 171 p.
- Eakin, T. E., 1961, Ground-water appraisal of Long Valley, White Pine and Elko Counties, Nevada: Nevada Dept. Conserv. and Nat. Resources, Ground-Water Resources - Reconnaissance Ser., Rept. 3, 35 p.
- _____, 1966, A regional interbasin ground-water system in the White River Area, Southeastern Nevada: Nevada Dept. Conserv. and Nat. Resources, Water Resources Bull. 33, p. 251-271.
- _____, 1967, Water-resources appraisal of Steptoe Valley, White Pine and Elko Counties, Nevada: Nevada Dept. Conserv. and Nat. Resources, Water-Resources Recon. Ser., Rept. 42, 48 p.
- Emmons, S. F., 1878, Geology of Franklin Buttes: U.S. Geol. Explor. 40th Parallel (King), v. 2, no. 18, p. 491-493.
- Fritz, W. H., 1957, Structure and stratigraphy of Telegraph Canyon area, northern Egan Range, east-central Nevada: Univ. of Washington, unpubl. thesis (PhD).
- _____, 1961, Structure and stratigraphy of the northern Egan Range, White Pine County, Nevada [abs.]: Dissert. Abs., v. 21, no. 12, p. 3748-3749.
- Hardman, George, 1965, Nevada precipitation map, adapted from map prepared by George Hardman and others, 1936: Nevada Univ. Agr. Expt. Sta. Bull. 183, 57 p.

- Harlow, George, 1956, Stratigraphy and structure of the Spruce Mountain area, Elko County, Nevada: Univ. of Washington, unpubl. thesis (PhD).
- Kohler, M. A., Nordenson, T. J., and Baker, D. R., 1959, Evaporation maps for the United States: U.S. Weather Bur. Tech. Pub. 37.
- Maxey, G. B., and Eakin, T. E., 1949, Ground water in White River Valley, White Pine, Nye, and Lincoln Counties, Nevada: Nevada State Engineer Water Resources Bull. 8, 59 p.
- Misch, Peter, 1960, Regional structural reconnaissance in central-northeast Nevada and some adjacent areas-- Observations and interpretations, in Geology of east-central Nevada: Intermountain Assoc. Petroleum Geologists, 11th Ann. Field Conf., 1960 Guidebook, p. 17-42.
- Moore, D. O., 1968, Estimating mean runoff in ungaged semiarid areas, in Internat. Assoc. Sci. Hydrology Bull. (in press).
- Nelson, R. B., 1956, The stratigraphy and structure of the region surrounding Currie, Elko County, Nevada: Univ. of Washington, unpubl. thesis (M.S.).
- Simpson, J. H., 1876, Explorations across the Great Basin of the Territory of Utah in 1859: Washington, U.S. Army Eng. Dept., 518 p.
- Snelson, Sigmund, 1955, Geology of the Southern Pequop Mountains, Elko County, N.E. Nevada: Univ. of Washington, unpubl. thesis (M.S.).
- Snyder, C. T., 1963, Hydrology of stock-watering development in the Ely Grazing District, Nevada: U.S. Geol. Survey Water-Supply Paper 1475-L, p. 383-441.
- Snyder, C. T., Hardman, George, and Zdenek, F. F., 1964, Pleistocene lakes in the Great Basin: U.S. Geol. Survey, Misc. Geol. Inv. Map I-416.
- Stevens, C. H., 1965, Pre-Kaibab Permian stratigraphy and history of Butte Basin, Nevada and Utah: American Assoc. Petroleum Geologists Bull., v. 49, no. 2, p. 139-156.
- Tagg, K. M., and others, compilers, 1964, Geologic map of Nevada, in Mineral and water resources of Nevada: Nevada Bur. Mines Bull. 65, fig. 3, opp. p. 12.

Threet, R. L., 1960, Geomorphology of east-central Nevada, in Geology of east-central Nevada: Intermountain Assoc. Petroleum Geologists, 11th Ann. Field Conf., 1960 Guidebook, p. 7-10.

U.S. Department of Agriculture, 1954, Diagnosis and improvement of saline and alkali soils: Agriculture Handbook no. 60, 160 p.

U.S. Public Health Service, 1962, Public Health Service drinking water standards 1962: Public Health Service Pub. no. 956, 61 p.

Worts, G. F., Jr., 1967, The available water supply, in Rush, F. E., and Glancy, P. A., Water-resources appraisal of the Warm Springs - Lemmon Valley area, Washoe County, Nevada: Nevada Dept. Conserv. and Nat. Resources. Water Resources Recon. Ser. Rept. 43, p. 48-53.

LIST OF PREVIOUSLY PUBLISHED REPORTS IN THIS SERIES

Report No.	Valley	Report No.	Valley
1	Newark (out of print)	28	Smith Creek and Ione
2	Pine (out of print)	29	Grass (near Winnemucca)
3	Long (out of print)	30	Monitor, Antelope, Kobeh
4	Pine Forest (out of print)	31	Upper Reese
5	Imlay area (out of print)	32	Lovelock
6	Diamond (out of print)	33	Spring (near Ely) (out of print)
7	Desert	34	Snake
8	Independence		Hamlin
9	Gabbs		Antelope
10	Sarcobatus and Oasis		Pleasant
11	Hualapai Flat		Ferguson Desert (out of print)
12	Ralston and Stonecabin	35	Huntington
13	Cave		Dixie Flat
14	Amargosa		Whitesage Flat (out of print)
15	Long Surprise	36	Eldorado - Piute Valley (Nevada and California)
	Massacre Lake Coleman	37	Grass and Carico Lake (Lander and Eureka Counties)
	Mosquito Guano	38	Hot Creek
	Boulder		Little Smoky
16	Dry Lake and Delamar		Little Fish Lake
17	Duck Lake	39	Eagle (Ormsby County)
18	Garden and Coal	40	Walker Lake
19	Middle Reese and Antelope		Rawhide Flats
20	Black Rock Desert		Whiskey Flat
	Granite Basin	41	Washoe Valley
	High Rock Lake	42	Steptoe Valley
	Summit Lake	43	Honey Lake Warm Springs
21	Pahranagat and Pahroc		Newcomb Lake Cold Spring
22	Pueblo Continental Lake		Dry Lemmon
	Virgin Gridley Lake		Red Rock Spanish Springs
23	Dixie Stingaree		Bedell Flat Sun
	Fairview Pleasant		Antelope
	Eastgate Jersey	44	Smoke Creek Desert
	Cowkick		San Emidio Desert
24	Lake		Pilgrim Flat
25	Coyote Spring		Painters Flat
	Kane Spring		Skedaddle Creek
	Muddy River Springs		Dry (near Sand Pass)
26	Edwards Creek		Sano
27	Lower Meadow Patterson		
	Spring (near Panaca)		
	Panaca Eagle		
	Clover Dry		

LIST OF PREVIOUSLY PUBLISHED REPORTS IN THIS SERIES--cont.

Report No.	Valley	Report No.	Valley
45	Clayton Alkali Spring Lida Stonewall Flat Oriental Wash Grapevine Canyon		
46	Mesquite Ivanpah Jean Lake Hidden		
47	Thousand Springs		
48	Goose Creek Salmon Falls Creek Bruneau River Jarbidge River Owyhee River (Duck) South Fork Owyhee Independence East Little Owyhee River		

EXPLANATION

UNCONSOLIDATED ROCKS

QUATERNARY

Pleistocene and Holocene

 Younger alluvium
 Unconsolidated gravel, sand, silt, and clay. Coarse material yields small to large quantities of water to wells where saturated

Pleistocene and older(?)

 Older alluvium
 Unconsolidated to semiconsolidated boulders, gravel, sand, silt, and clay. Coarse material yields small to large quantities of water where saturated

CONSOLIDATED ROCKS

CAMBRIAN TO TRIASSIC

Middle Cambrian to Triassic

 Carbonate rocks
 Limestone and dolomite. May yield small to large quantities of water from fracture and solution cavities

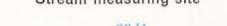
PRECAMBRIAN TO TERTIARY


 Noncarbonate rocks
 Igneous, sedimentary, and metamorphic rocks. May yield small amounts of water from fracture zones

EXPLANATION


 Drainage divide


 Geologic contact


 Stream measuring site

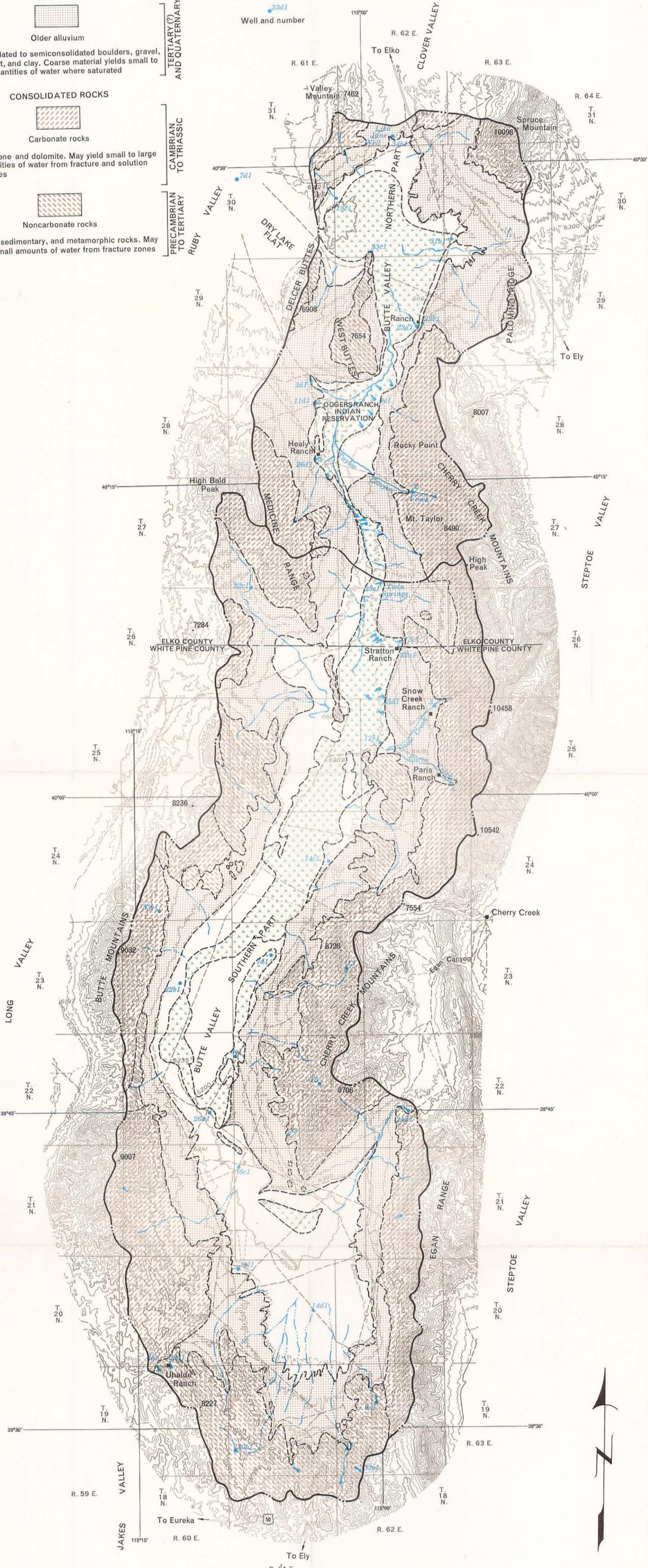

 Well and number


 Spring and number


 Phreatophytes


 Scale
 0 1 2 3 4 5 Miles

Contour Interval 200 feet
 Datum is mean sea level.



Base From Army Map Service 1:250,000 Scale
 Series: Elko (1955) and Ely (1956)

Hydrology by Patrick A. Glancy 1968. Geology partly synthesized and modified from Harlow (1956), Fritz (1960), and Tagg and others (1964).

PLATE 1.—GENERALIZED HYDROGEOLOGIC MAP OF BUTTE VALLEY, ELKO AND WHITE PINE COUNTIES, NEVADA