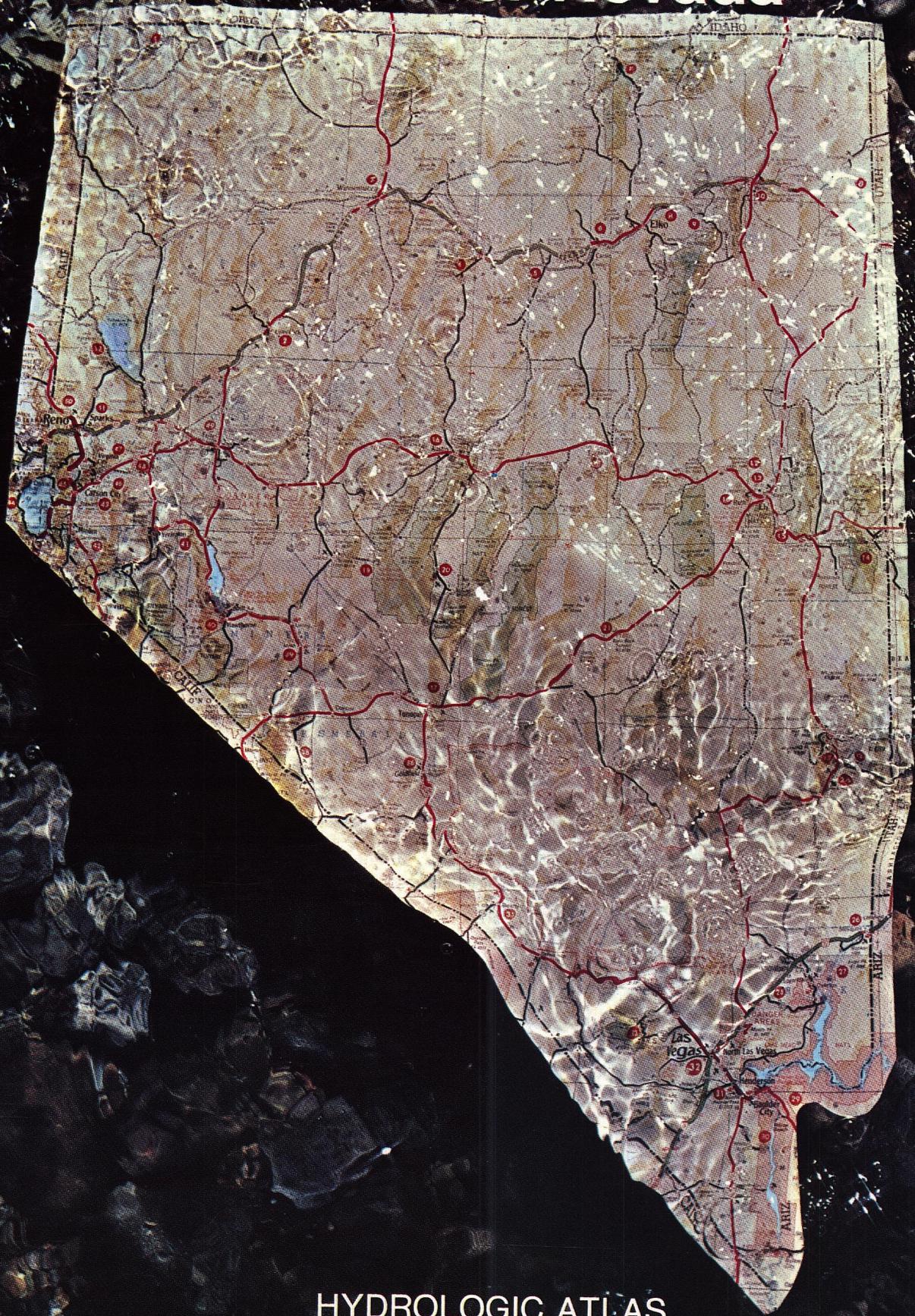


Water for Nevada



HYDROLOGIC ATLAS

ELMO J. DeRICCO
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STATE OF NEVADA

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DEPARTMENT OF CONSERVATION AND NATURAL RESOURCES

DIVISION OF WATER RESOURCES

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In reply refer to
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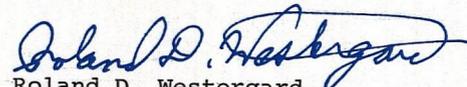
This Hydrologic Atlas is one of a continuing series of publications developed by the Division of Water Resources as a part of the comprehensive State Water Plan. It is the result of the efforts of many people. B. J. Vasey, Planning Engineer for the Division of Water Resources conceived this atlas and supervised its production. Bruce Scott, a member of the Division planning staff, was instrumental in the preparation of many of the maps. He was responsible for the coordination of the various maps and publication of this atlas. Lawrence Roach, Jr., also of the Division of Water Resources assisted in compilation and did the cartography on most of the maps. F. Eugene Rush of the U. S. Geological Survey in Carson City and the Branch of Technical Illustrations of the same agency in Menlo Park, California were also major contributors.

Several maps in this atlas are based on State Federal Comprehensive framework studies. Many of the others are results of the cooperative program between the State of Nevada and the U. S. Geological Survey.

Maps showing such information as average annual precipitation, runoff and evaporation, designated ground water areas, vegetal cover, land suitability and existing lakes and reservoirs provide general information for the state as a whole. Bathymetric surveys for some of Nevada's major lakes and reservoirs, and a map of water resources and interbasin flows, give more detailed information concerning specific areas of the state.

This atlas constitutes another portion of the inventory phase in the development of the State Water Plan. Periodically, new maps will be added to the series, both to update earlier maps and to provide additional information.

Respectfully,


Roland D. Westergard
State Engineer

WATER FOR NEVADA HYDROLOGIC ATLAS

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(June 1972)

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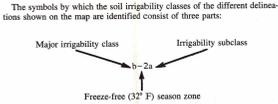
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IRRIGABILITY CLASSIFICATION OF SOILS

This classification is a rating system which indicates the potential usefulness of different kinds of soil for irrigated agriculture in Nevada. It is based on both judgments of the feasibility of initial improvements such as drainage, removing stones, leveling, ripping hardpan, or removing soluble salts and alkali, and on continuing limitations for use requiring special management, such as control of salt buildup. It does not consider the present availability or lack of irrigation water. The classification and map are intended for general planning, rather than for location of specific fields for agriculture or for identification of specific management needs. But, it does indicate major limitations for initial improvements and major hazards to continued use.

Five major classes of potential usefulness for irrigation are presented here; these major classes are divided into subclasses to indicate the nature of the most outstanding limitations. Since length of growing season is so important in determining the kinds of crops which will be grown, and their economic value, the irrigability classes are subdivided again according to the freeze-free (32° F) season zone in which the various soils occur (Sakamoto et al., 1972).



The five major soil irrigability classes A, B, C, D, and E are defined in Table 1 in general terms of decreasing potential usefulness and increasing limitations or hazards. Class A soils are the most valuable; class E soils have such serious limitations they can not be used for irrigated agriculture. The freeze-free (32° F) season zone in which the various soils occur (Sakamoto et al., 1972).

Each major class in Table 1, and a brief description of the kind of outstanding limitation is given there. The number of days' length of the five freeze-free (32° F) season zones of Nevada are also given in Table 1.

Each soil irrigability class includes many different kinds of soils with sometimes rather different properties. However, all the soils in a class offer about the same potential usefulness or degree of limitation. The soils of any one class should respond in similar manner for irrigation, but they might require somewhat different management. They are apt to be differently suitable for other uses, such as for roads, foundations, septic tank filter fields, or recreational areas. The criteria by which the soil irrigability classes were assigned to soils were adapted with limited modification from Pacific Southwest Interagency Committee criteria. The soil irrigability subclasses are defined only to reflect soil and landscape conditions in Nevada. All criteria were developed jointly by the Division of Water Resources, Nevada Department of Conservation and Natural Resources, the Agricultural Experiment Station, University of Nevada Reno, and the Soil Conservation Service, U.S. Department of Agriculture.

References:
Sakamoto, C. M., F. P. Peterson, E. A. Naphan, H. P. Cebal, R. B. Gardner, and R. O. Gilford. 1972. Freeze-free (32° F) seasons of the major basins and plateaus of Nevada. Water for Nevada, Hydrology Atlas, State of Nevada Water Planning Report. Division of Water Resources, Nevada Department of Conservation and Natural Resources, Carson City, Nevada.

*Criteria for soil irrigability classes, Nevada Reclamation Soil Survey (Class. 372). Division of Water Resources, Nevada Department of Conservation and Natural Resources, Agricultural Experiment Station, University of Nevada, Reno, and Soil Conservation Service, U.S. Department of Agriculture.

TABLE I
SOIL IRRIGABILITY CLASSES AND SUBCLASSES
(Legend for use with the explanatory map of the irrigable soils of Nevada)

Irrigability class and subclasses ¹	Freeze-free (32° F) season zones (days)				
	> 200	130-200	100-130	70-100	< 70
CLASS A — Soils that have slight or few limitations that restrict their use for irrigated agriculture	A-1	A-2	A-3	A-4	A-5
CLASS B — Soils that have moderate limitations that reduce choice of crops or require moderate conservation practices	B-1a	B-2a	B-3a	B-4a	B-5a
CLASS C — Soils that have severe limitations that reduce choice of crops or require special conservation practices or both	C-1a	C-2a	C-3a	C-4a	C-5a
CLASS D — Soils that have very low available waterholding capacity	D-1a	D-2a	D-3a	D-4a	D-5a
CLASS E — Soils having very low available waterholding capacity and 4 to 15 percent slopes	E	E	E	E	E

¹The specific class and subclass criteria are: "Criteria for Soil Irrigability Classes," Nevada Reclamation Soil Survey (Class. 372), Agricultural Experiment Station, University of Nevada Reno, Division of Water Resources, Nevada Department of Conservation and Natural Resources, Carson City, Nevada.

IRRIGABILITY CLASSIFICATION OF SOILS AND SOIL RESOURCE GROUPS

Economic projections and evaluations require a system of land classification. The Economic Research Service, Soil Conservation Service, and Forest Service have established a uniform classification system termed Soil Resource Groups (SRGs). An SRG is defined as a group of land capability units having similar cropping patterns, yield characteristics, responses to fertilizers, and management and land treatment measures. They were developed primarily

for use in USDA River basin planning. SRG data is widely used in river basin studies and provides a comparable base for USDA agencies for evaluating agricultural production, employment and income impacts; for comparing alternative projects and programs; and for displaying location of resource base proposals.

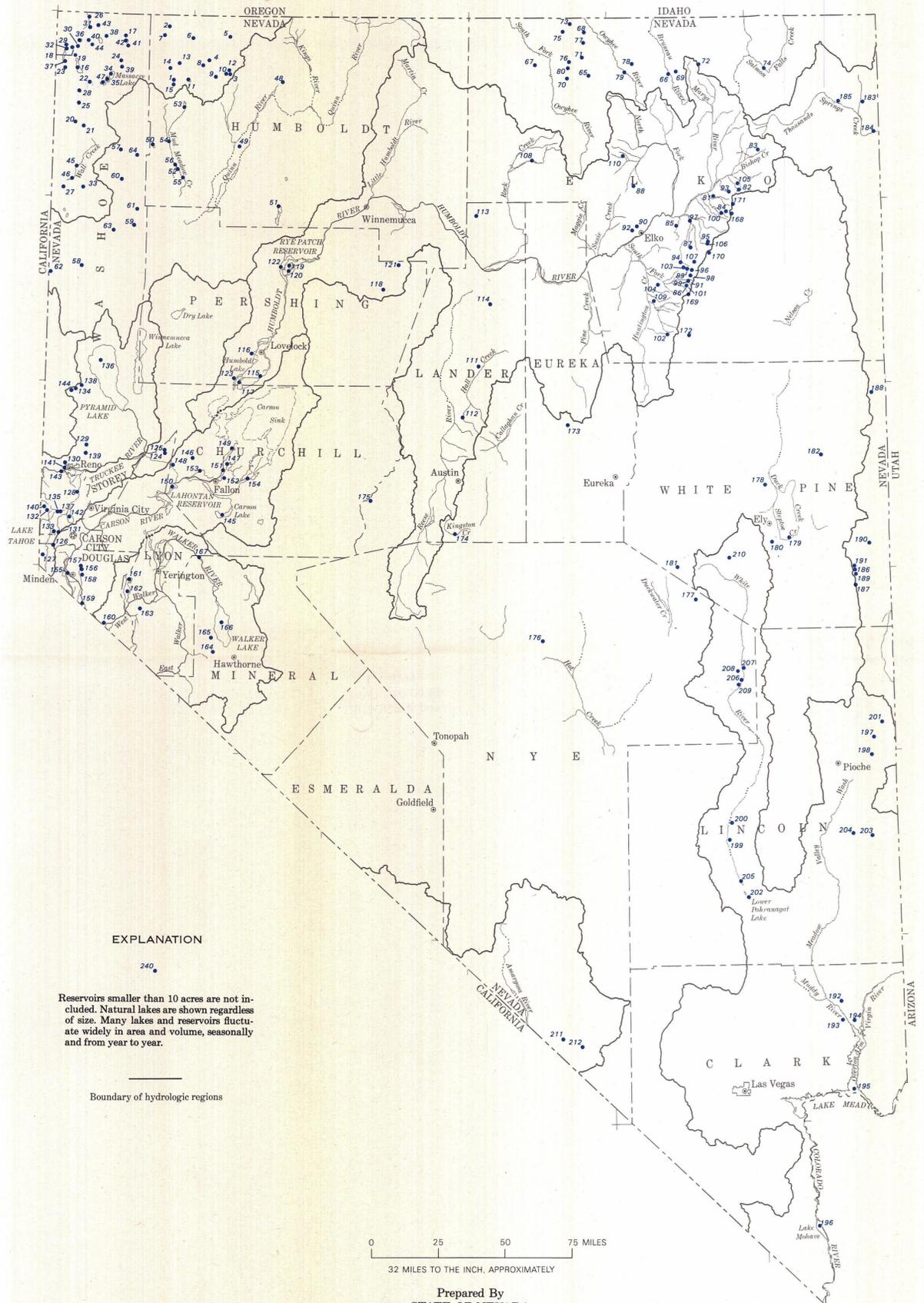
The IRRIGABILITY CLASSIFICATION OF SOILS can be converted to SRGs for this purpose by utilizing the following conversion table.

TABLE II
CONVERSION OF NEVADA IRRIGABILITY SUBCLASS (IS) TO SOIL RESOURCE GROUP (SRG), JANUARY 1, 1973

Irrigability Subclass	Soil Resource Group								
A-1	1	B-2w	92A	B-5a	17A	C-2w	88A1	C-3c	17A
A-2	4A	B-2c	67A	B-5b	96	C-2w	88A2	C-3c	17A
A-3	4B	B-3a	131B	B-5c	10	C-2w	92A	C-3c	17A
A-4	7	B-3b	67B	B-5d	96	C-2w	67A	C-3c	17A
A-5	7	B-3c	131B	B-5e	10	C-2w	14B	C-3c	17A
B-1a	25	B-3d	31B	B-5f	12A	C-2w	14B	C-3c	17A
B-1b	25	B-3e	31B	B-5g	12A	C-2w	14B	C-3c	17A
B-1c	25	B-3f	31B	B-5h	12A	C-2w	14B	C-3c	17A
B-1d	25	B-3g	31B	B-5i	12A	C-2w	14B	C-3c	17A
B-1e	25	B-3h	31B	B-5j	12A	C-2w	14B	C-3c	17A
B-1f	25	B-3i	31B	B-5k	12A	C-2w	14B	C-3c	17A
B-1g	25	B-3j	31B	B-5l	12A	C-2w	14B	C-3c	17A
B-1h	25	B-3k	31B	B-5m	12A	C-2w	14B	C-3c	17A
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B-1v	25	B-3y	31B	B-5aa	12A	C-2w	14B	C-3c	17A
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B-1z	25	B-3ac	31B	B-5ae	12A	C-2w	14B	C-3c	17A
B-2a	25	B-3ad	31B	B-5af	12A	C-2w	14B	C-3c	17A
B-2b	25	B-3ae	31B	B-5ag	12A	C-2w	14B	C-3c	17A
B-2c	25	B-3af	31B	B-5ah	12A	C-2w	14B	C-3c	17A
B-2d	25	B-3ag	31B	B-5ai	12A	C-2w	14B	C-3c	17A
B-2e	25	B-3ah	31B	B-5aj	12A	C-2w	14B	C-3c	17A
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B-2q	25	B-3at	31B	B-5av	12A	C-2w	14B	C-3c	17A
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LAKES & RESERVOIRS

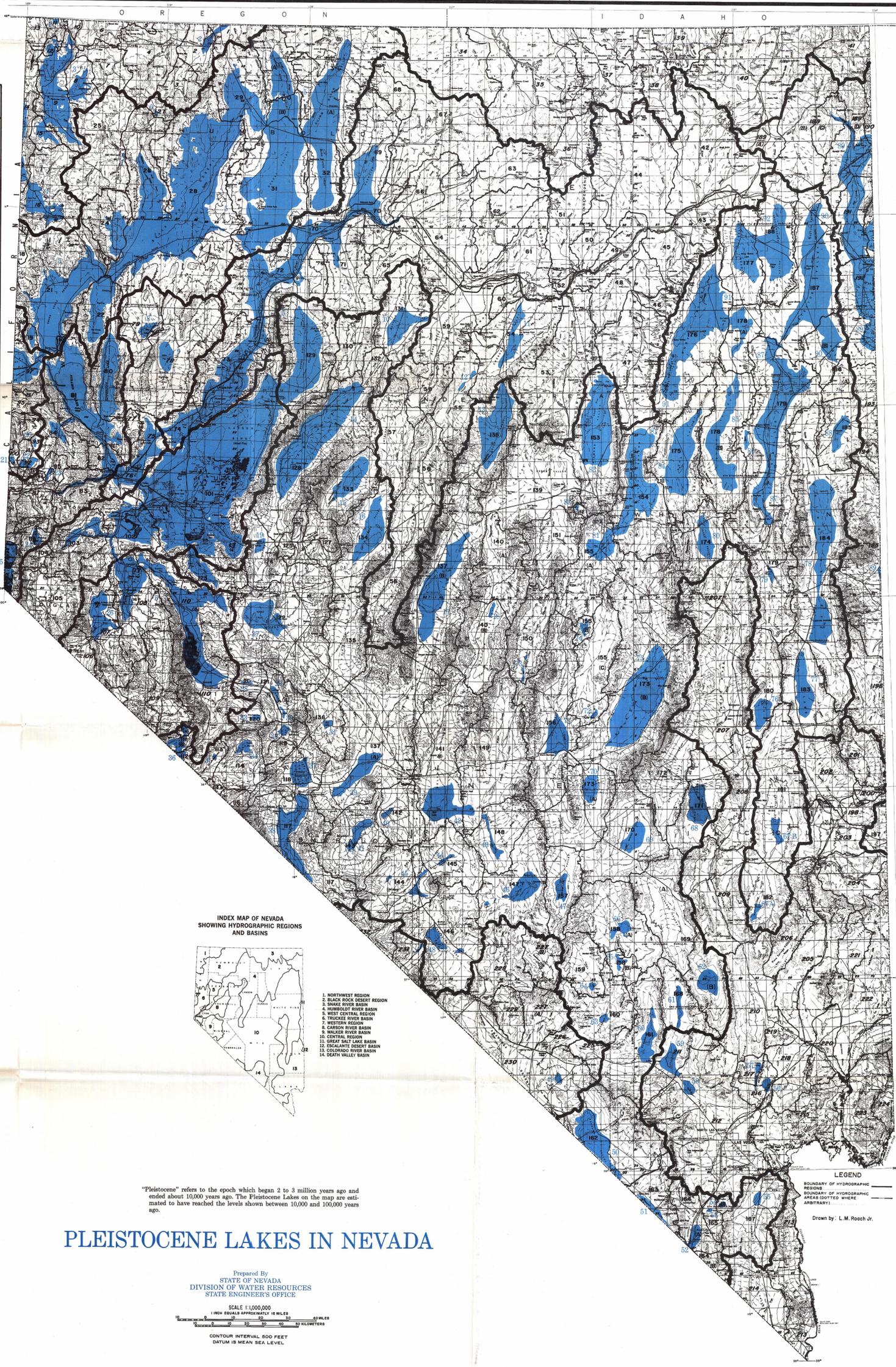
Name of Reservoir	Map Number	Surface Area (Ac.) ¹	Capacity (Ac. Ft.) ²	Name of Reservoir	Map Number	Surface Area (Ac.) ¹	Capacity (Ac. Ft.) ²
NORTHWEST REGION				WESTERN CENTRAL REGION			
Humboldt County				Lyon County			
Alkali Reservoir	1	97	1,233	Fernley Dam #1	124	276	910
Big Spring Reservoir	2	480	1,680	Fernley Dam #3	125	95	476
Blue Lakes	3	20	(120)	TRUCKEE RIVER BASIN			
Bog Hot Reservoir	4	38	154	Douglas County			
Continental Lake	5	500	(4,250)	Sqooner Lake	126	69	(400)
Dufurrena Ponds	6	25	150	Lake Tahoe	127	121,000 ³	745,000 ⁴
Gooch Lake	7	154	(154)	Washoe County			
Gridley Lake	8	320	(320)	Lake Alexander	128	58	(250)
Knott Creek Reservoir	9	98	1,620	Gasperi	129	30	90
Little Onion Reservoir	10	30	325	Highland	130	10	(54)
Onion Lake	11	774	(774)	Hobart Creek Reservoir	131	10	110
Onion Valley Reservoir	12	101	1,500	Incline Lake	132	30	157
Rock Springs Table Reservoir	13	40	500	Marlette Lake	133	350	10,400
Smith Lake	14	5	(10)	Milk Ranch Reservoir	134	23	252
Virgin Valley Reservoir	15	20	50	Price Lake	135	10	(54)
Washoe County				Pyramid Lake			
Alkali Lake	16	3750	(3,750)	Rock Lake	136	108,000	20,510,000
Bald Mountain Lake	17	216	(216)	Spanish Flat Reservoir	137	20	(105)
Big Holy Lake	18	500	(1000)	Spanish Springs	138	236	1,000
Broyles Reservoir	19	78	(510)	Tamarack Lake	139	30	(185)
Boulder Lake	20	15	(78)	Virginia Lake	140	10	(54)
Boulder Reservoir	21	10	40	Washoe Lakes	141	24	140
Cap Johnson Reservoir	22	160	1,500	Wheeler Reservoir (Evans Creek)	142	4,100	20,400
Carter Reservoir	23	222	935	Winnemucca Ranch Reservoir	143	46	948
Catnip Reservoir	24	55	220	CARSON RIVER BASIN			
Central Lake	25	321	(150)	Churchill County			
Coleman Reservoir	26	40	500	Carson Lake	145	2,000	(4,000)
Duck Lake	27	3000	(3,000)	Desert Gun Club Reservoirs	146	100	500
Forty-nine Lake	28	352	(200)	Harmon Lake	147	200	400
Frog Pond Dam	29	92	395	Hazen Reservoir	148	10	20
Hill Dam	30	77	396	Indian Lakes	149	400	(3,500)
Little Valley Reservoir	31	80	400	Lahontan Reservoir	150	10,000	273,600
Little Holy Lake	32	30	(30)	Old River Reservoir	151	270	500
Lost Creek Reservoir	33	12	98	Ollies Pond "S"	152	350	700
Massacre Lake	34	2,532	(3,000)	Soda Lake	153	600	(5,300)
Middle Lake	35	1,198	(900)	Stillwater Point Reservoir	154	1,900	19,000
Mosquito Lake	36	935	(400)	Douglas County			
New Years Lake	37	1,500	6,000	Bose Reservoir	155	30	90
Racetrack Reservoir	38	32	75	Dangberg Reservoir #1 & 2	156	45	375
Rye Cross Reservoir	39	200	498	Dangberg Reservoir #3	157	80	500
Mud Lake	40	900	(1,200)	Dangberg Reservoir #4	158	150	1,000
Swan Lake	41	130	(300)	Mud Lake	159	300	1,800
Swan Lake Reservoir	42	500	1,000	WALKER RIVER BASIN			
Toney Reservoir	43	15	(85)	Douglas County			
Wimer Reservoir	44	70	(350)	Topaz Lake	160	2,300 ⁵	59,400 ⁶
Wall Creek Dam	45	182	400	Lyon County			
Wall Creek Dam #2	46	133	2,200	Artesia Lake	161	1,000	1,000
West Lake	47	1,248	(900)	Beaman Lake	162	80	480
BLACK ROCK DESERT REGION				Nuti Reservoir			
Humboldt County				Mineral County			
Bilk Creek Reservoir	48	110	800	Cat Creek Reservoir	164	25	1,155
Delong Reservoir	49	500	2,275	Rose Creek Reservoir	165	32	656
High Rock Lake	50	650	(500)	Walker Lake	166	38,800	2,990,000
Jungo Flat Lake	51	10	25	Weber Reservoir	167	950	(13,000)
Mud Meadows Reservoir	52	80	215	CENTRAL REGION			
Summit Lake	53	560	(5,000)	Elko County			
Van Vleck Reservoir	54	250	2,750	Steele Lake/Gibbs Lake	168	6	(30)
Weiss & Vogel Reservoir	55	150	450	Overland Lake	169	20	(120)
Wheeler Reservoir (Donnley Creek)	56	154	1,100	Robinson Lake	170	17	(95)
Washoe County				Winchell Lake			
Denio Camp Reservoir	57	30	110	Ruby Lake	172	9,000	(30,000)
Dewey Parker Reservoir	58	156	428	Eureka County			
Fly Reservoir	59	40	350	Roberts Creek Reservoir			
Grass Valley Reservoir	60	10	50	Lander County			
Negro Creek Dam	61	50	497	Grove Lake	174	17	190
Smoke Creek Reservoir ⁷	62	90	1200	Smith Creek Reservoir	175	25	350
Squaw Valley Reservoir	63	47	1,200	Nye County			
Woodruff Reservoir	64	128	(500)	Fish Lake	176	80	(160)
SNAKE RIVER BASIN				Manzonie Reservoir			
Elko County				White Pine County			
Bull Run Reservoir	65	106	1,246	Bassett Lake	178	120	(1,300)
Charleston Reservoir	66	40	200	Cave Creek	179	32	784
Chimney Creek Reservoir	67	928	9,950	Comins Lake	180	40	290
Coyote Hole Reservoir	68	18	(36)	Bull Creek #2	181	10	51
Coyote Lake	69	25	(50)	Spring Valley Wash Dam	182	64	121
Deep Creek Reservoir	70	136	1,410	GREAT SALT LAKE BASIN			
Dry Creek Reservoir	71	110	1,910	Elko County			
Emerald Lake	72	1	(4)	Crittenden Reservoir	183	230	4,300
Groundhog Reservoir	73	16	(32)	Daek Reservoir	184	2,909	5,340
Jakes Creek Reservoir	74	62	472	23 Mile Reservoir	185	652	7,457
Josephine Reservoir	75	250	1,250	White Pine County			
Rawhide Reservoir	76	147	1,540	Baker Lake	186	10	(50)
Sheep Creek Reservoir	77	850	7,500	Dead Lake	187	3	(10)
Sunflower Reservoir	78	60	(120)	Goshute Reservoir	188	200	300
Wildhorse Reservoir	79	3,000	72,000	Johnson Lake	189	5	(25)
Wilson Reservoir	80	828	10,468	Silver Creek Reservoir	190	13	200
HUMBOLDT RIVER BASIN				Stella Lakes			
Elko County				COLORADO RIVER BASIN			
Ackler Lake	81	10	(54)	Clark County			
Angel Lake	82	13	(70)	Bowman Reservoir	192	165	4,000
Bishop Creek Reservoir	83	782	30,000	Glassand Pond	193	16	(53)
Boulder Lake	84	6	(30)	Honey Bee Pond	194	32	(100)
Boyd Reservoir	85	120	(85)	Lake Mead	195	164,000	29,700,000
Castle Lake	86	9	(48)	Lake Mohave	196	28,200	1,820,000
Cold Lake	87	6	(30)	Lincoln County			
Dorsey Creek Reservoir	88	14	150	Eagle Valley Reservoir	197	59	640
Echo Lake	89	29	(175)	Echo Reservoir	198	64	1,400
Eight Mile Creek Reservoir	90	45	944	Frenchy Lake	199	74	(150)
Favre Lake	91	19	(110)	Hiko Lake	200	246	(500)
Fifth St. Wash Reservoir	92	10	94	Hollinger Debris Basin	201	90	640
Greys Lake	93	5	(25)	Lower Pahrangat Lake	202	583	(1000)
Griswold Lake	94	15	(85)	Mathews Canyon Reservoir	203	420	12,420
Hidden Lake	95	9	(48)	Pine Canyon Reservoir	204	354	12,470
Island Lake	96	7	(35)	Upper Pahrangat Lake	205	370	3,580
John Day Reservoir	97	127	561	Nye County			
Lamoille Lake	98	13	(70)	Dacey Reservoir	206	215	784
Liberty Lake	99	21	(125)	Hay Meadow Reservoir	207	203	1,120
Lost Lake	100	3	(14)	Sunnyside Reservoir	208	791	3,330
North Furlong Lake	101	8	(40)	Tule Field Reservoir	209	218	507
Pearl Lake	102	5	(25)	White Pine County			
Seitz Lake	103	3	(14)	Preston Reservoir	210	109	1,271
Sleeman Ponds	104	12	20	DEATH VALLEY BASIN			
Smith Lake	105	3	(14)	Nye County			
Soldier Lake	106	6	(30)	Lake #1	211	69	243
Verdi Lake	107	5	(25)	Lake "C"	212	70	618
Willow Creek Reservoir	108	761	18,064	¹ Values in parentheses are estimates			
Zamino Reservoir	109	30	(180)	² More than 90% in California			
Saval Reservoir	110	10	15	³ Total Area—36,400 Acres in Nevada, remainder in California			
Lander County				⁴ Total within the 6-foot operating range.			
Carico Lake	111	1,032	1,550	⁵ Total Area and Capacity—1250 acres in Nevada, remainder in California			
Iowa Canyon Reservoir	112	28	437				
Izzenhood Ranch Reservoir	113	10	50				
Nelson Reservoir	114	13	100				
Pershing County							
Big Five Dam	115	787	1,720				
Graveyard Slough	116	80	100				
Humboldt Lake	117	4,200	(46,000)				
Mud Springs Dam	118	70	(490)				
Upper Pitt-Taylor	120	(1,700)	20,200				
Lower Pitt-Taylor	121	(1,700)	20,200				
Pumpernickel Reservoir	122	37	236				
Rye Patch	122	10,800	179,000				
Toulon Lake	123	3,500	(38,000)				



EXISTING LAKES AND RESERVOIRS

Lake Name	Lake Area (sq. mi.)	Drainage Basin Area (sq. mi.)	Max. Depth of Water (Feet)	Spilled	Spilled To	Valleys Covered by Lake Area
1. Warner	459*	2,548*	250	No		11. Coleman V.
2. Alvord	491*	2,242*	200	No		1. Pueblo V.
3. Meiner	742	300-		Yes	Warner V.	2. Continental Lake V.
4. Surprise	506*	1,461*	650	No		8. Mono Lake V.
5. Lahontan	8,665*	42,322*		No		9. Long V.
6. Unnamed	4	38		Yes	L. Lahontan	10. Macy Flat
7. High Rock	11	700		Yes	L. Lahontan	12. Mosquito V.
8. Kumiva	22	354		No		15. Warner V.
9. Granite Spgs.	41	256	50	No	L. Lahontan	14. Surprise V.
10. Buffalo	55	509		Yes	L. Lahontan	16. Duck Lake V.
12. Carlo	6	373		Yes	L. Lahontan	19. Dry V.
13. Gilbert	209	543	250	Yes	L. Lahontan	20. Sassa V.
14. Dixie	120	2,438	255	No		21. Smoke Creek Desert
15. Edwards	98	396	150	No		22. San Emilio Desert
16. Desatoya	140	590		No		24. Hardsnap Flat
17. Toyabe	250	1,295	170	No		26. Mud Meadow
18. Diana	6	627		Yes	Diamond V.	28. Black Rock Desert
19. Laboo	15	300		Yes	Lake Dixie	29. Pine Forest V.
20. Fred	4	22		No		30. Kings River V.
21. Loughton	5	50		No		(A) Es King Subarea
22. Lemmon	15	93	80+	Yes	L. Lahontan	(B) Sol House Subarea
23. Unnamed	7	72		Yes	L. Lahontan	31. Desert V.
24. Washoe	25	97		Yes	L. Lahontan	(A) Owens Subarea
25. Tahoe	211*	551*	300	Yes	do	(B) McDermitt Subarea
26. Wellington	90	920		Yes	do	69. Paradise V.
27. Gabbs	120	1,280		No	do	70. Winnemucca Segment
28. Acme	170	Shallow		No	do	71. Grass V.
29. Luning	10	160		No	do	72. Limy Area
30. Mina	2	30		No	do	73. Lovelock V.
31. Monte Cristo	8	227		No	do	(A) Owens Subarea
32. Rhodes	15	170		No	do	74. White Plains
33. Garfield	3	93		No	do	75. Brady's Hot Springs Area
34. Teal	18	278		No	do	76. Fernley Area
35. Hammon	3	146		No	do	77. Winnemucca Lake V.
36. Russell	287*	794*	750	Yes	Owens River	78. Pyramid Lake V.
37. Columbus	60	1,400		No		79. Kuniya V.
38. White Mt.	185*	1,072*	80	Yes	Columbus Salt Marsh	80. Winnemucca Lake V.
39. Tonopah	95	3,085		No		81. Pyramid Lake V.
40. Goldfield	7	300		No		82. Dodge Flat
41. Clayton	39	547	Shallow	No		83. Tracy Segment
42. Mud	165	2,945		No		84. Spanish Springs V.
43. Cactus Flat	9	358	Shallow	No		85. Honey Lake V.
44. Stonewall	9	358		Yes	Sarcobatus	101. Carson Desert
45. Gold Flat	31	982	Shallow	No		102. Churchill V.
46. Lida	30	335		Yes		103. Dayton V.
47. Kawich	33	303		No		104. Mason V.
48. Bonno Caire	57	1,485	Shallow	No		105. Walker Lake V.
49. Ash Meadows	4	292		Yes	Death Valley	(A) Schurz Subarea
50. Pahump	242*	992*		No	do	(B) Lake Subarea
51. Mesquite	54*	446*		No	do	(C) Whiskey Flat-Hawthorne Subarea
52. Ivanpah	67*	888*		No	do	123. Hawside Flat
53. Jean	2	94		No		124. Buena Vista V.
54. East Jean	3	84		No		125. Summit Lake V.
55. Eldorado	11	323		No		27. Summit Lake V.
56. Apex	25	215		Yes	Apex	28. High Rock Lake V.
57. Apex North	20	110		Yes	Apex	29. Kuniya V.
58. Corn Creek	43	625		Yes	Colorado R. Lake Corn Cr.	78. Granite Spgs. V.
59. East	62	641		No		101. Crescent V.7
60. Indian Springs	21	380		Yes	Lake East	53. Carson Lake V.
61. Norman	51	974		No		138. Grass V.
62. Unnamed	9	455		No		139. Dixie V.
63. Frenchman	7	200		No		140. Nevada Creek V.
64. Yuca	9	160	Shallow	Yes	Possibly into Groom Lake	141. Smith Creek V.
65. Papoose	25	790		Yes		142. Big Smoky V.
66. Groom	10	355		No		143. Soda Spring V.
67A. Delamar	25	1,180		No		(B) Western Part
67B. Bristol	83	981	40	No		(A) Eastern Part
68. Coal	60	663		No		85. Spanish Springs V.
69. Penoyer	50	2,188		Yes	Lake Railroad	86. Washoe V.
70. Reveille	88	1,487		Yes	Railroad V.	87. Lake Tahoe Basin
71. Hot Creek	4	491	Shallow	Yes	Hot Creek V.	107. Smith V.
72. Unnamed	10	69		Yes		122. Gabbs V.
73. Unnamed	525	3,565		No		121. Soda Spring V.
74. Unnamed	51	356		No		(A) Eastern Part
75. Carpenter	107	533		No		122. Gabbs V.
76. Spring	332	1,641		Yes	White River	123. Soda Spring V.
77. Unnamed	10	338		Yes		(B) Southern Part
78. Jakes	67	429		Yes	Lake Steptoe	124. Fairview V.
81. Newark	357	1,886		Yes		93. Antelope V.
82. Yahoo	3	19		Yes	Lake Diamond	100. Coal Spring V.
83. Diamond	294	3,153	120	Yes	Lake Lahontan	92. Lemmon Valley
84. Hubbs	205	635	250	Yes		(A) Western Part
85. Gale	181	745	150	Yes	Ruby Valley (Lake Franklin)	(B) Eastern Part
86. Unnamed	19	62		Yes	Steptoe V. (Lake Franklin)	88. Spanish Springs V.
87. Steptoe	459	1,785	350	Yes	Goehute V.	89. Washoe V.
88. Antelope	47	335	75	Yes	Snake River	90. Lake Tahoe Basin
89. Bonneville	19,940*	53,325*		Yes		107. Smith V.
90. Waring	478	3,290	175	No		122. Gabbs V.
91. Franklin	471	1,897	210	Yes	Lake Clover	121. Soda Spring V.
92. Clover	342	1,016		No		(A) Eastern Part

*Lakes whose areas and/or drainage areas are only partially in Nevada. Adapted from "Pleistocene Lakes in the Great Basin," by Snyder, Hartman and Zdenek.



INDEX MAP OF NEVADA SHOWING HYDROGRAPHIC REGIONS AND BASINS

- NORTHWEST REGION
- BLACK ROCK DESERT REGION
- SHAKE RIVER BASIN
- HUMBOLDT RIVER BASIN
- WEST CENTRAL REGION
- TRUCKEE RIVER BASIN
- WESTERN REGION
- CARSON RIVER BASIN
- WALKER RIVER BASIN
- CENTRAL REGION
- GREAT SALT LAKE BASIN
- COLORED RIVER BASIN
- DEATH VALLEY BASIN

LEGEND
BOUNDARY OF HYDROGRAPHIC REGIONS
BOUNDARY OF HYDROGRAPHIC AREAS (DOTTED LINE)
ARBITRARY
Drawn by: L. M. Roach Jr.

"Pleistocene" refers to the epoch which began 2 to 3 million years ago and ended about 10,000 years ago. The Pleistocene Lakes on the map are estimated to have reached the levels shown between 10,000 and 100,000 years ago.

PLEISTOCENE LAKES IN NEVADA

Prepared By
STATE OF NEVADA
DIVISION OF WATER RESOURCES
STATE ENGINEER'S OFFICE
SCALE 1:1,000,000
1 INCH EQUALS APPROXIMATELY 16 MILES
CONTOUR INTERVAL 500 FEET
DATUM IS MEAN SEA LEVEL

RECONNAISSANCE BATHYMETRIC MAP AND GENERAL HYDROLOGY OF LAHONTAN RESERVOIR, NEVADA

By T. L. Katzer

Introduction

Lahontan Reservoir is about 35 miles east of Carson City and 15 miles west of Fallon, Nev. (fig. 1). This report redefines the altitude-area-capacity relations of Lahontan Reservoir, and describes the general hydrology of the reservoir. The field work was done under the supervision of E. E. Harris, U.S. Geological Survey, and in cooperation with the U.S. Bureau of Reclamation and the Nevada Department of Fish and Game.

Historical Sketch

The Federal Reclamation Act of 1902 authorized construction of irrigation projects in the 17 Western States (U.S. Bureau of Reclamation, 1961, p. 534-542). The Newlands Project, of which Lahontan Reservoir is a part, was authorized by Congress in 1903 and first phase construction began that same year. The project was originally named the Truckee-Carson Project, but was renamed the Newlands Project in 1919 in honor of the late U.S. Senator Francis G. Newlands of Nevada.

Construction on Lahontan Dam did not start until January 1911 and was completed in June 1915. The dam is a zoned, earthfill structure, 162 feet high, and has a volume of 733,000 cubic yards. The spillway-crest altitude is 4,162 feet above sea level (1917 datum). Operating criteria allow the provisional installation of 20-inch-high flashboards on the spillway crest, which brings the maximum storage altitude to about 4,164 feet. The altitude of the outlet invert is 4,070 feet.

Precautionary reservoir drawdowns, based on basin runoff forecasts, have limited the reservoir to several spills during the 54 years of operation. The minimum reservoir stage of 4,070.0 feet was reached in September 1929 during a prolonged drought. The lowest stage during recent years was 4,096.6 feet in October 1961. Figure 2 shows the maximum and minimum stages and volumes of stored water in Lahontan Reservoir for the period of record.

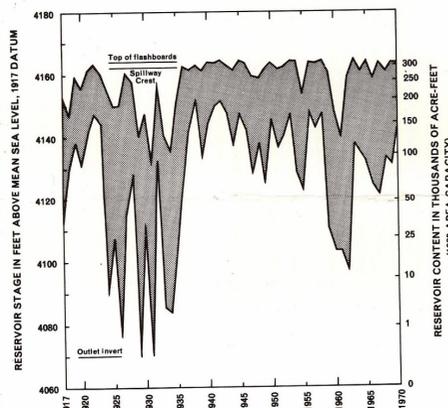


Figure 2.—Annual maximum and minimum stages and volume of stored water of Lahontan Reservoir, 1917-70 calendar years

The Truckee-Carson Irrigation District was formed on December 19, 1926 to be the legal operating agency for the project works, and on December 31, 1926, the Department of the Interior transferred the operation and maintenance of the project from the U.S. Bureau of Reclamation to the Irrigation District; however, the U.S. Bureau of Reclamation is still responsible for setting certain operational and maintenance standards.

Project features now (1971) include Lahontan Dam and Reservoir on the Carson River, a dam and outlet works on the Truckee River at Lake Tahoe (fig. 1), three diversion dams, and nearly 900 miles of canals, laterals, and drains. Also included are two hydroelectric power plants with accessory substations and transmission lines.

Water Sources

The principal sources of water for Lahontan Reservoir are the drainages of the Carson and Truckee Rivers (about 3,120 square miles), whose headwaters are in the high Sierra Nevada where precipitation is much greater than in the eastern parts of the basins. The waters of the two rivers are partly depleted prior to reaching Lahontan Reservoir principally by diversion for irrigation and public supply, seepage loss to ground-water systems, and evapotranspiration by native vegetation.

The Carson River system flows northeast from the Sierra Nevada to Lahontan Reservoir, irrigating about 51,000 acres upstream from the reservoir, and then through the Fallon farming area downstream from the reservoir and terminates in the Carson Sink about 20 miles northeast of Fallon. Several small surface-water reservoirs provide about 15,000 acre-feet of storage upstream from the reservoir. The river is gaged near Fort Churchill (fig. 1), approximately 10 miles upstream from the reservoir, and the flow has averaged about 264,000 acre-feet per year for the past 58 years. Figure 3 shows the monthly flow distribution at the Fort Churchill gage.

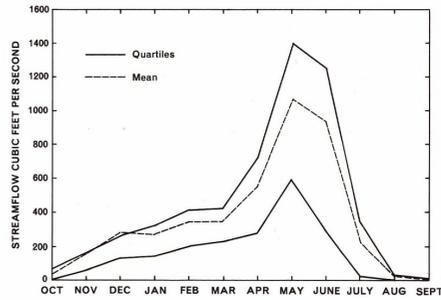


Figure 3.—Mean monthly flow distribution, Carson River near Fort Churchill, 1919-69 water years. Quartiles show 25 percent of the monthly flows were higher and lower than indicated.

Truckee River water is stored in six main reservoirs (fig. 1) upstream from the diversion to Lahontan Reservoir and is used to irrigate approximately 23,000 acres (U.S. Bur. Reclamation, oral commun., 1969). Water is transferred from the Truckee River basin to Lahontan Reservoir in the Carson River basin through the Truckee Canal, which heads at Derby Dam (fig. 1). The inflow to Lahontan Reservoir is measured at a gage on the Truckee Canal near Hazen. There is one diversion from the canal, the KX Lateral (fig. 1), between the gage near Hazen and Lahontan Reservoir.

Table 1 shows the total surface-water inflow to Lahontan Reservoir for 1964-69 water years.

Table 1.—Surface-water inflow to Lahontan Reservoir, 1964-69 (Inflow in acre-feet, rounded)

Water year ¹	Carson River ²	Truckee Canal ³	Total
1964	136,000	425,000	391,000
1965	382,000	244,000	626,000
1966	171,000	289,000	460,000
1967	449,000	209,000	658,000
1968	162,000	114,000	276,000
1969	561,000	107,000	668,000

Average, 1964-69	310,000	203,000	513,000
Average, 1919-69	252,000	518,000	5430,000

- October 1 to September 30.
- Excludes return flow from the Buckland Ditch.
- Excludes diversions to KX Lateral canal.
- First year of measured flow in Truckee canal below turnouts to the Truckee River.
- Estimated (Lamke, R. D., oral commun., 1970).

Stage data for Lahontan Reservoir have been collected since January 1917 by the U.S. Bureau of Reclamation and the Truckee-Carson Irrigation District and are published as follows:

Period	U.S. Geological Survey publication
1917-65	Water-Supply Paper 1734
1961-65	Water-Supply Paper 1927
1966-70	Water Resources Data for Nevada, 1965, 1966, 1967, 1968, 1969, 1970 (annual reports).

Summaries of flows for the Carson River and Buckland Ditch also are published in the above reports. Records for Truckee Canal near Hazen have been published by the U.S. Geological Survey since October 1966. Prior records are unpublished, but are available from the Carson City office of the U.S. Bureau of Reclamation. The KX Lateral records of diversions are also available from the same source.

The Buckland Ditch (fig. 1), which diverts flow from the Carson River just upstream from the Fort Churchill gage, provides irrigation water for about 2,100 acres. The average annual diversion by this ditch for the past 7 years has been 16,400 acre-feet. Return flow to the river from this diversion may average about 50 percent.

The water from Lahontan Reservoir is used to irrigate about 57,000 acres of farmland in the Fallon area (U.S. Bur. Reclamation, oral commun., 1969). Figure 4 shows the annual releases from the reservoir for the period of record. In addition to the irrigated land, approximately 25,000 acres in the Stillwater Wildlife Area and 28,000 acres in Carson Lake Pasture receive streamflow, return flow from irrigated lands, and flow from reservoir spills.

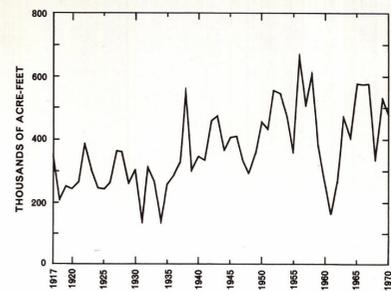


Figure 4.—Lahontan Reservoir releases to the Carson River, 1917-70 calendar years

General Reservoir Hydrology

The reservoir pool mainly overlies Quaternary lake deposits of ancient Lake Lahontan, younger fan gravels, and fluvial deposits. In the general vicinity of the reservoir, Quaternary basalt crops out and consists of thin lava flows with interbeds of scoriaceous basalt breccia and diatomaceous deposits. Also cropping out are older Tertiary andesite, dacite, and rhyolite interbedded with breccia and agglomerate (Morrison, 1964, and Speed and Willden, 1966).

Construction of Lahontan Dam changed the ground-water system of the surrounding area. As the reservoir filled, water seeped into the ground-water system, causing a rise in nearby ground-water levels. Some water probably is seeping from the reservoir through the volcanic rocks and sediments that are present in the eastern subsurface of the reservoir in the vicinity of the dam. Ground-water levels measured in June 1970 in the vicinity of the reservoir were all within a few feet of the reservoir surface. These high ground-water levels indicate that reservoir water recharges the local ground-water system. As Lahontan Reservoir changes stage, a corresponding change probably occurs in the surrounding ground-water reservoir. The magnitude of the response of ground-water levels to changes in reservoir stage varies inversely with distance from the reservoir.

The average annual precipitation on the reservoir is about 5 inches, and the average pan evaporation rate adjusted to Lahontan free water surface is about 54 inches (U.S. Bur. Reclamation, oral commun., 1970). At a spillway crest altitude of 4,162 feet, the reservoir has a surface area of about 11,200 acres (table 3). Thus, when the reservoir is full, the average inflow from precipitation is about 4,700 acre-feet, and the loss from evaporation is about 50,000 acre-feet. About 22,000 acres of phreatophytes surrounding the reservoir are estimated by the author and P. A. Glancy, U.S. Geological Survey, to use an average of about 8,000 acre-feet of ground water per year.

For irrigation use, the quality of surface water in Lahontan Reservoir has been generally satisfactory at the time of sampling. The abundance of the biologic nutrients, orthophosphate, nitrate, and the relatively low values for dissolved oxygen, are of primary concern to the esthetic and fishery quality of the reservoir. Table 2 summarizes the results of chemical analyses of water samples collected at four different locations on Lahontan Reservoir (fig. 1).

Sedimentation data are limited to one sample taken by the U.S. Geological Survey on the Carson River at Weeks (fig. 1), which is about 10 miles upstream from the reservoir. During the period May 11 to June 10, 1969, the mean daily flow was between 3,000 and 4,000 cfs (cubic feet per second), which represents a runoff value of about 210,000 acre-feet, or 37 percent of the flow for the 1969 water year. The sediment sample, collected on June 1 during the spring-flow recession, had a suspended-sediment concentration of 622 milligrams per liter. River flow at the Fort Churchill gage at this time was 3,400 cfs, when considered with the sediment concentration, represents a suspended-sediment load of about 5,700 tons per day moving past the sampling site.

Some material probably is being deposited between the sampling site and the reservoir, and some in the reservoir, slowly reducing its water-storage capacity. Some very fine suspended sediment moves out of the reservoir, but no data are available to determine this amount. Lack of extensive sedimentation in the reservoir downstream from the river delta is shown by the trace of the old river channel, labeled thalweg on the bathymetric map (fig. 5). The pre-reservoir channel is still prominent, showing only minimal effects of sedimentation.

Table 2.—Water-quality data for four sites in Lahontan Reservoir

Sampling sites, in downstream order (shown in fig. 1)	[Data from Nevada State Health Division] ¹			
	1	2	3	4
Temperature (°C)				
Maximum	27	24	24	22
Average	20	18	18	16
Minimum	10	12	10	11
pH				
Maximum	8.6	8.3	8.3	8.6
Average	7.8	8.0	8.1	7.9
Minimum	7.5	7.6	7.6	7.5
Dissolved oxygen (mg/l)				
Maximum	15	9.1	9.7	9.2
Average	8.2	7.7	7.8	7.7
Minimum	5.4	6.1	6.4	5.1
Orthophosphate (PO ₄ , in mg/l)				
Maximum	0.73	0.85	1.0	1.6
Average	.42	.49	.52	.86
Minimum	.28	.20	.20	.30
Nitrate (NO ₃ , in mg/l)				
Maximum	1.2	4.8	4.8	10
Average	.6	1.5	1.8	2.4
Minimum	.0	.0	.0	.0
Dissolved solids (residue on evaporation at 105°C, in mg/l)				
Maximum	200	233	238	183
Average	161	168	167	152
Minimum	121	129	116	119

¹ Based on samples collected during spring and summer, 1966-69. Samples were collected from a boat, at about 0-1 foot depth.

Reservoir Area and Capacity

Fathometer soundings, topographic maps, and a series of aerial photographs taken by the U.S. Bureau of Reclamation, showing the surface area of the reservoir at various stages, were used to develop the reconnaissance bathymetric map (fig. 5) and stage-area-capacity relations (table 3 and fig. 6). A more detailed stage-capacity table is available, and shows reservoir stage to 0.1 foot (U.S. Bur. Reclamation, Carson City project office, written commun., 1970).

Forty-nine continuously recording-fathometer traverses were made by boat across Lahontan Reservoir. Two transit teams on opposite shores were used to maintain directional control. Contour lines were fitted to the data. The areas between the contours were planimeted and the area and volume of the reservoir computed. The overall limit of error of the bathymetric data is considered to be less than 5 percent.

The 1929 datum is 3.73 feet higher than the 1917 datum, but for simplicity and continuity, the new area and capacity table is based on the 1917 datum, which is in common usage.

Prior to October 1970, the so-called 1917 area-capacity relations were in use (table 3); however, beginning in October 1970, all concerned water agencies started using the new (1969) area-capacity relations. The differences between the 1917 and 1969 area-capacity relations are variable, and at an altitude of 4,070 feet, the 1969 area-capacity table shows only about 24 percent of the volume of the 1917 area-capacity table. This percentage difference decreases at increasing reservoir stage until at altitude 4,089 feet the relation is virtually the same. Above this altitude, the 1969 capacity data are as much as 13 percent but generally 6 to 8 percent larger than the 1917 capacity data. The large shallow area in the southeast part of the reservoir, designated area A in figure 5, is now used for storage and was not included in the 1917 calculations. This area has a minimum altitude of 4,147 feet and contains about 10,600 acre-feet of water and has a surface area of about 3,740 acres when filled to altitude 4,164 feet. Even if area A is deleted from the 1969 capacity data, the reservoir contains 16,000 acre-feet more storage than was originally calculated in 1917.

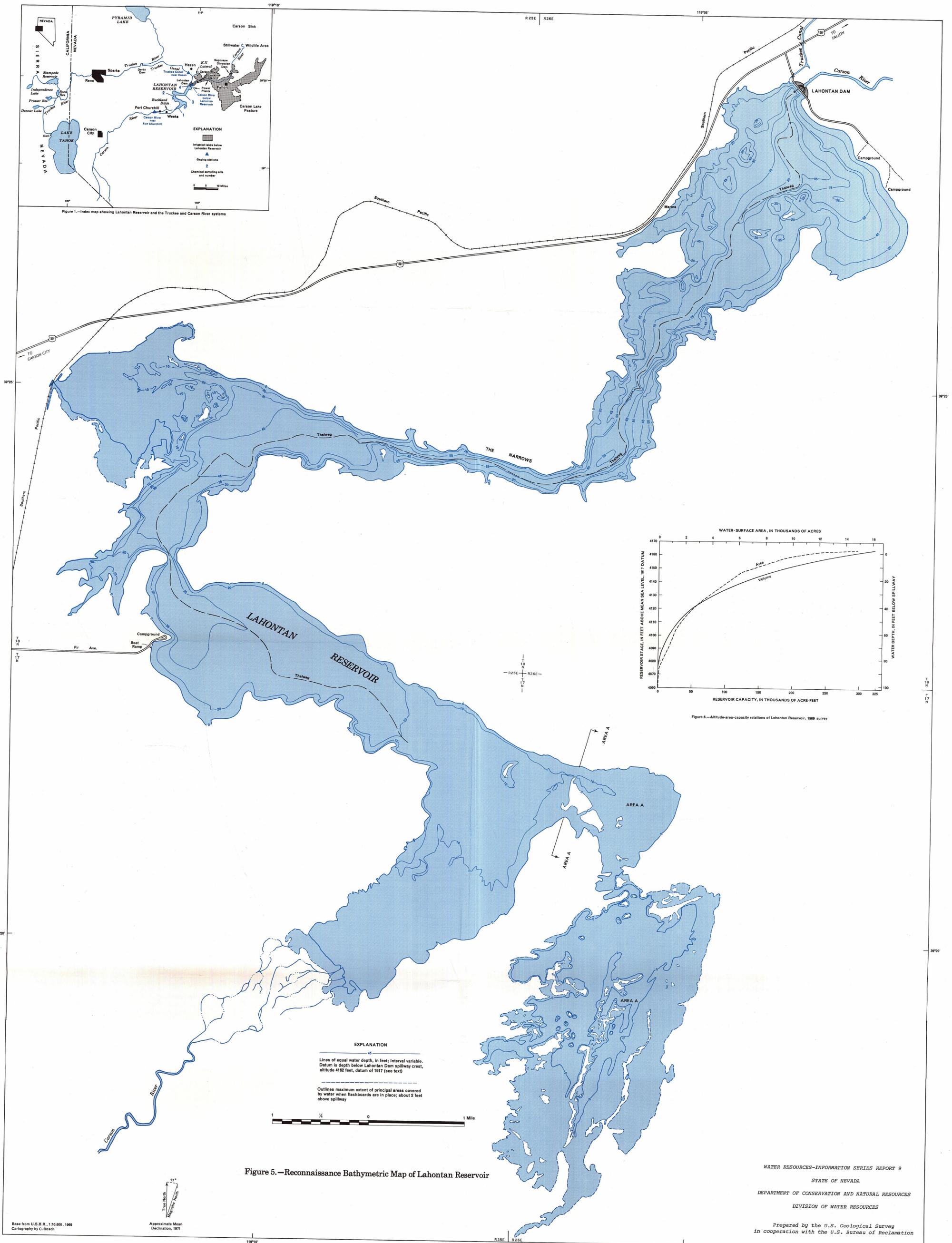
Table 3.—Lahontan Reservoir area and capacity data for 1917 and 1969

Reservoir stage in feet above sea level (1917 datum) ¹	Water-surface area (acres)		Water in storage (acre-feet, rounded)	
	1917 rating	1969 rating	1917 rating	1969 rating
4,060		0		0
4,065		7		17
4,070	100	30	385	91
4,075	145	94	960	368
4,080	255	263	1,990	1,170
4,085	420	505	3,640	3,080
4,090	555	743	6,070	6,210
4,095	755	950	9,440	10,500
4,100	985	1,130	13,900	15,700
4,105	1,240	1,360	19,600	21,800
4,110	1,500	1,720	26,500	29,500
4,115	1,840	2,140	34,900	39,100
4,120	2,420	2,610	45,800	50,900
4,125	3,000	3,170	59,800	65,300
4,130	3,600	3,820	76,400	82,700
4,135	4,500	4,510	97,300	104,000
4,140	4,800	5,190	121,000	128,000
4,145	5,200	5,880	146,000	155,000
4,150	6,400	6,960	175,000	187,000
4,155	7,700	8,470	211,000	226,000
4,160	9,200	10,700	254,000	273,000
4,162	9,800	12,100	274,000	295,000
4,164	10,600	14,800	294,000	5322,000

- The 1917 datum is 3.73 feet lower than the 1929 datum.
- Outlet invert altitude.
- Spillway crest altitude.
- Approximate altitude of top of flashboards.
- Includes extensive shallow-water area in southeast part of reservoir (area A in fig. 5).

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RECONNAISSANCE BATHYMETRIC MAP AND GENERAL HYDROLOGY OF LAHONTAN RESERVOIR, NEVADA

By
T. L. Katzer
1972

WATER RESOURCES—INFORMATION SERIES REPORT 9
STATE OF NEVADA
DEPARTMENT OF CONSERVATION AND NATURAL RESOURCES
DIVISION OF WATER RESOURCES
Prepared by the U.S. Geological Survey
in cooperation with the U.S. Bureau of Reclamation

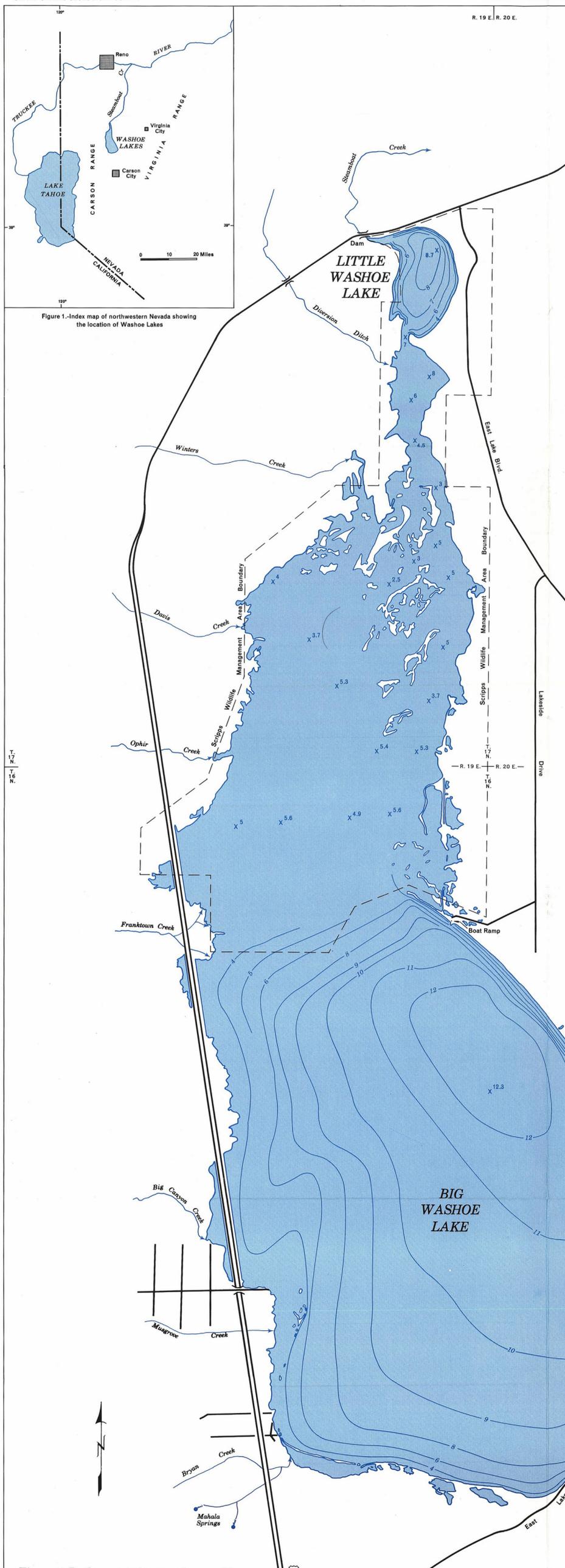


Figure 2.-Bathymetric Reconnaissance Map of Big and Little Washoe Lakes

EXPLANATION

Lines of equal water depth, in feet; interval variable. Datum is lake surface at a stage of 5029.8 feet above mean sea level

Control point
Number is depth of water, in feet



Base from aerial photographs, 1:19,000, 1971
Cartography by C. Bosch

R. 19 E., R. 20 E.

INTRODUCTION

Big and Little Washoe Lakes are 15 miles south of Reno, Nev., and 5 miles north of Carson City on the floor of Washoe Valley (fig. 1). Mountains of the Carson Range of the Sierra Nevada and the Virginia Range surround the valley, except for a narrow canyon to the north through which Steamboat Creek drains the lakes to the Truckee River. Washoe Valley has an area of about 84 square miles.

The principal source of water for the lake is snow-melt in the Carson Range. Small quantities of water are imported to the Washoe Lakes from Browns, Galena, and Third Creeks west and north of the basin. Some water that would normally flow to Big Washoe Lake is exported from the basin to Carson City and Virginia City through the Marlette Water System. A small dam (about 30 feet wide and less than 10 feet high) at the north end of Little Washoe Lake (fig. 2), having a narrow spillway (about 8 feet wide) at an altitude of 5,028.9 feet, is used in part to control the stage of the lakes and the outflow to Steamboat Creek for downstream irrigation.

During most days of the winter, Big Washoe Lake is not frozen over with ice.

Washoe Lake is used mostly for fishing and boating; however, those uses are slight, probably because the lake water is highly turbid (Hutchinson, 1937, p. 85). However, the beaches and sand dunes along the east side of Big Washoe Lake are picturesque. The area is being considered as a site for a State park.

The Washoe Lakes were at a stage of 5,029.8 feet above sea level during June 1971, when the bathymetric survey was made. The lakes usually are a single body of water, because of the flooding of the intervening marshy area between Franktown Creek and the diversion ditch shown in figure 2. Following unusually dry periods of several years, the lakes have gone dry. The last time was in 1934 (Rush, 1967, p. 11). Figure 3 is a summary of stage variations for the period 1963-71.

BATHYMETRY

A continuously recording, electronic fathometer was used to measure the depths of the Washoe Lakes. The author was assisted by his son, Steven, in making the 19 traverses of the two lakes. The results of the survey are summarized in figures 2 and 4. The dimensions of the lakes at the time the survey was made, for the extremes in stage for the period 1963-71, and at spillway stage are summarized in table 1.

Table 1.—Dimensions of the Washoe Lakes at various stages

Dimension	Big Washoe Lake	Little Washoe Lake	Total
JUNE 1971			
Stage, in feet	5,029.8 ± 0.1	5,029.9 ± 0.1	--
Area, in acres	5,800		5,900
Volume, in acre-feet	33,000		34,000
AT SPILLWAY STAGE			
Stage, in feet	5,028.9	5,028.9	--
Area, in acres	5,700	110	5,800
Volume, in acre-feet	30,000	500	31,000
MAXIMUM RECORDED STAGE, 1963-71 ¹			
Stage, in feet	5,030.6	5,030.6	--
Area, in acres	5,900	112	6,000
Volume, in acre-feet	36,000	850	37,000
MINIMUM RECORDED STAGE, 1963-71			
Stage, in feet	a5,023.5	b5,021.9	--
Area, in acres	3,160	16	--
Volume, in acre-feet	10,000	6	--

1. On February 12 and 24, 1970.
- a. On November 24, 1964.
- b. On October 28, 1964.

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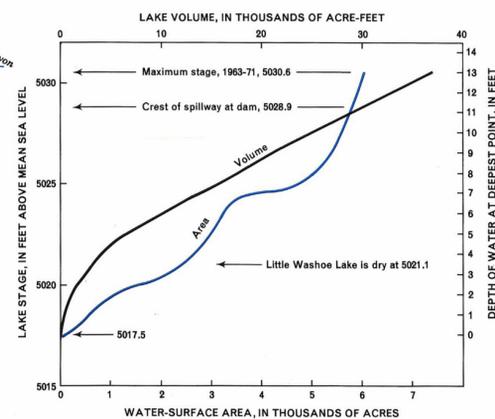


Figure 3.—Stage-area-volume relations for Big and Little Washoe Lakes

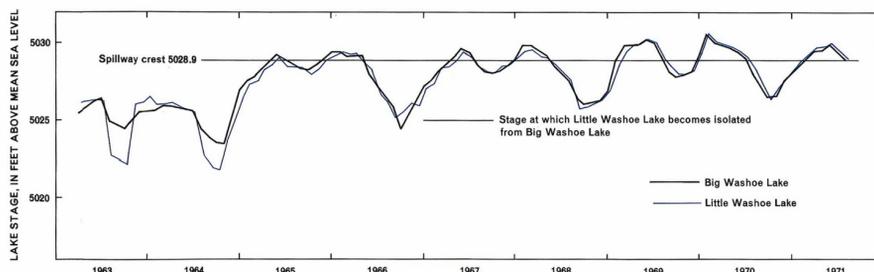


Figure 4.—Variations in lake stage, 1963-71

BATHYMETRIC RECONNAISSANCE OF BIG AND LITTLE WASHOE LAKES, WASHOE COUNTY, NEVADA

By
F. Eugene Rush
1972

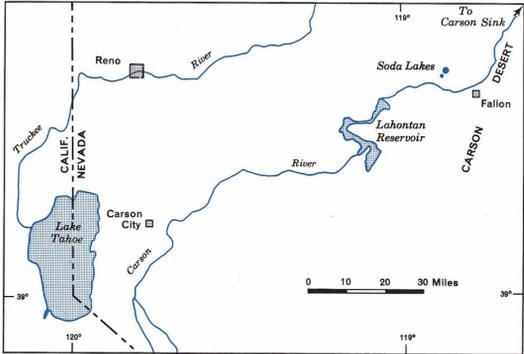


Figure 1.- Index map of the Reno-Fallon area of northwestern Nevada showing the location of the Soda Lakes

INTRODUCTION

Big and Little Soda Lakes are about 7 miles northwest of Fallon, Nev. (fig. 1), and about 50 miles east of Reno on the floor of Carson Desert. The two lakes occupy small craters composed of volcanic, basaltic sand (Morrison, 1964, p. 71 and Willden and Speed, 1968), surrounded by a broad expanse of alluvium. The form of the craters is shown in figure 2. On June 29, 1971, the stage of Big Soda Lake was at an altitude of 3,989.3 feet and Little Soda Lake, at 3,989.8 feet. The surrounding valley floor is generally at an altitude of about 4,000 feet. Regional ground-water flow is generally northeastward toward Carson Sink.

Prior to large-scale irrigation, fresh water entered the Soda Lakes mostly by subsurface percolation (Lee and Clark, 1916, p. 666) and numerous small springs (King, 1878, v. 1, p. 512 and Russell, 1885, pl. 16) flowing from the adjacent saturated alluvium about the lakes. In addition, some thermal water (temperature, 86°F or 30°C) may be entering Big Soda Lake near its center (Breese, 1968, p. 25). Evaporation, the principal form of discharge from the lakes under native conditions, offset the inflow. The lakes functioned much like large wells or sumps below the general ground-water level of the area. Under native conditions prior to 1906, the lake stages were maintained at a fairly constant low level, as shown in figure 3. In 1906, when extensive irrigation began in the area, the lake levels began to rise, continuing until about 1930. The total rise in stage for the period was about 60 feet. The principal cause of the rise was attributed to seepage losses from the T, U, and N canals (fig. 2), which carried water from the Carson River to fields in the Fallon area as part of the Newlands Project of the U.S. Bureau of Reclamation (Lee and Clark, 1916, p. 672-675).

Big Soda Lake is highly saline, as shown in table 1, and reportedly has no fish population (Breese, 1968, p. 2). Extraction of soda (essentially equal parts of sodium carbonate and sodium sulfate, according to Russell, 1885, p. 78-79) from Big Soda Lake began in 1875 (Russell, 1885, p. 79, and Morrison, 1964, p. 116) and continued until the facilities were flooded by rising water. Little Soda Lake is much less saline, but was used for soda extraction beginning in 1867 (Russell, 1885, p. 80). Little Soda Lake was reported (King, 1878, v. 1, p. 512) to have been nearly dry in 1867.

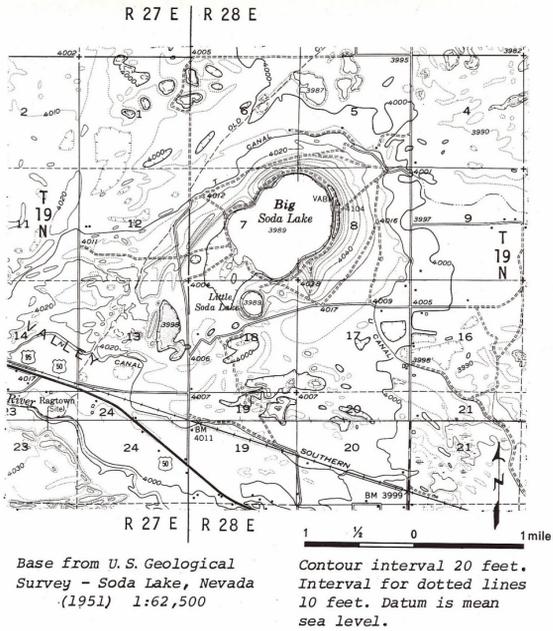


Figure 2.- Topography of the Soda Lakes area

BATHYMETRY

A continuously recording electronic fathometer was used to measure depths in the Soda Lakes in June 1971. Data were gathered along 17 traverse lines, utilizing 10 control points along the shore of Big Soda Lake. The author was assisted by his daughter, Teresa, in collecting the field data. The bathymetry, stage-area-volume relations, and a cross section are shown in figures 4, 5, 6, and 7. The deepest point surveyed in Big Soda Lake was 207 feet below the lake surface. An island that is shown on Russell's bathymetric map of 1882 (1885, pl. 16) was located at a depth of 58 feet, as shown in figures 4 and 6. Big Soda Lake (fig. 6) is within a compound crater. Below a depth of about 80 feet, a small, steep-sided crater occupies the center of the larger crater of topographic prominence, as shown in figures 2 and 6.

Little Soda Lake had a maximum depth of about 56 feet when this survey was made (fig. 4).

Lake stage, area, and volume for (1) equilibrium under near native conditions (prior to 1907) and for (2) equilibrium under irrigation conditions (since 1930) are summarized as follows:

	Near native conditions (Prior to 1907)	Irrigation conditions (Since 1930)
Big Soda Lake:		
Stage (feet)	3,930 ± 2	3,988 ± 2
Area (acres)	270 ± 5	385 ± 5
Volume (acre-feet)	15,500 ± 500	35,000 ± 500
Little Soda Lake:		
Stage (feet)	3,937 ± 4	3,989 ± 2
Area (acres)	76	20 ± 1
Volume (acre-feet)	770	770 ± 50

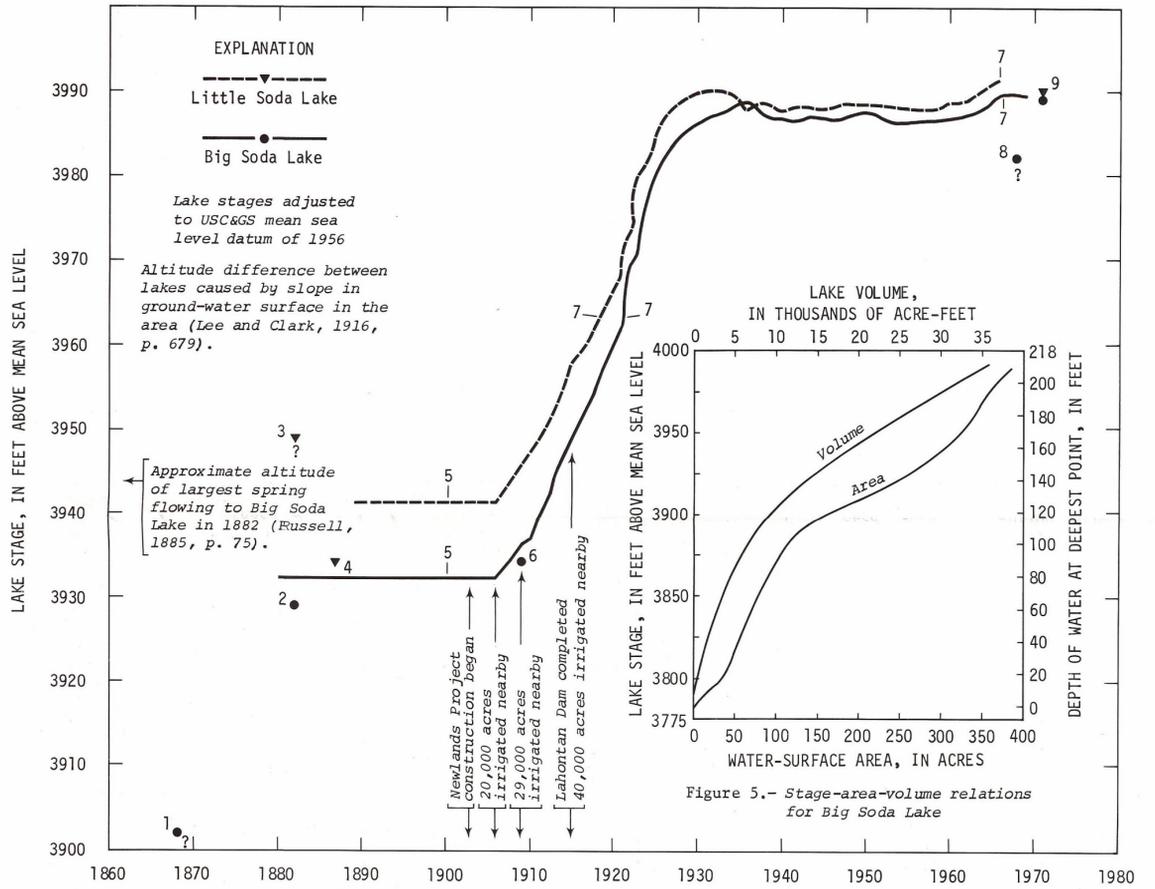


Figure 3.- Variations in the stages, areas and volumes of the Soda Lakes

FOOTNOTES FOR FIGURE 3

1. From King (1877, v. 4, p. 351); may be in error.
2. Based on lake area and water depth as reported by Russell (1885, p. 73).
3. Based on footnote by Russell (1885, p. 73). Probably not a natural stage.
4. As reported by Breese (1968, p. 8), Chatard (1890) observed Little Soda Lake to be dry. When first discovered, reported to have been dry (Russell, 1885, p. 79).
5. From Lee and Clark (1916, p. 705-706) for period 1880-1910. Data adjusted to U.S. Coast and Geodetic Survey mean sea level datum of 1956 by subtracting 3.7 feet from reported data.
6. From Strahorn and van Duyne (1911), as reported by Breese (1968, p. 13).
7. Data for period 1911-69 from U.S. Bureau of Reclamation.
8. From Breese (1968, p. 14). Probably in error.
9. Determined as part of present study in June 1971: Big Soda Lake, 3,989.3 feet; Little Soda Lake, 3,989.8 feet.

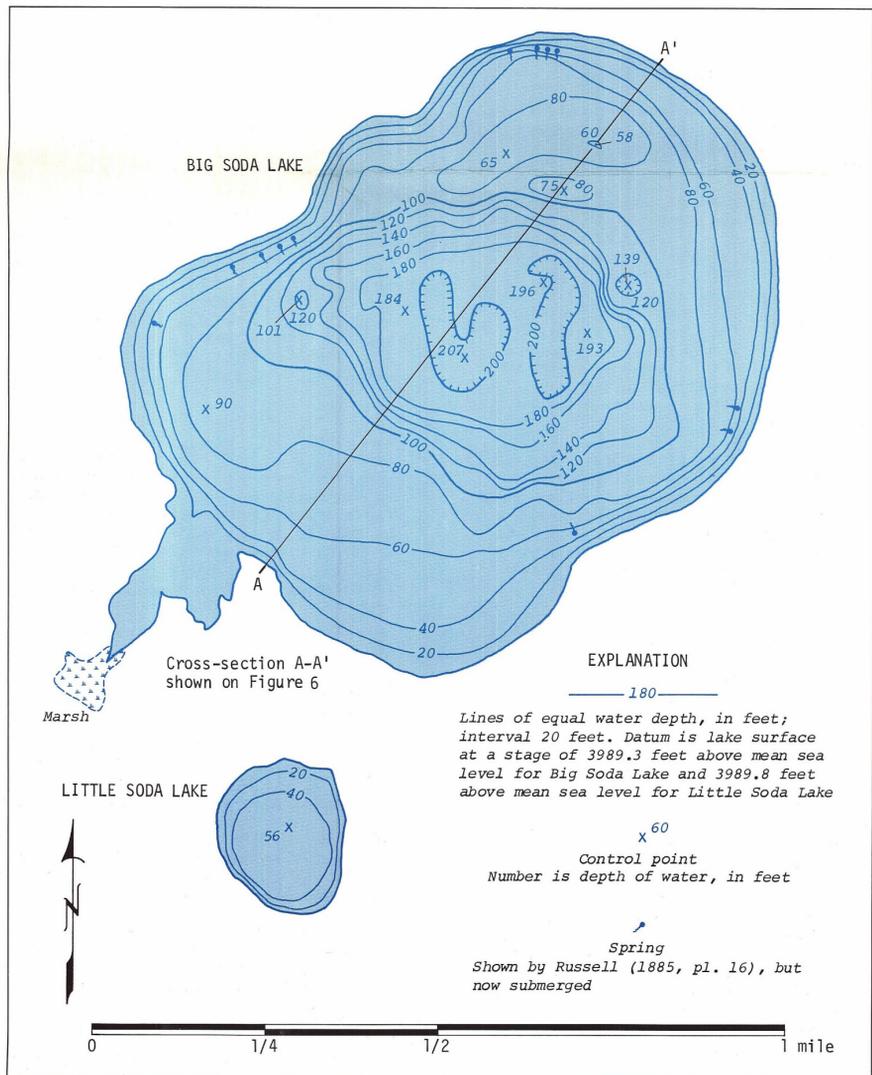


Figure 4.- Reconnaissance bathymetry of the Soda Lakes

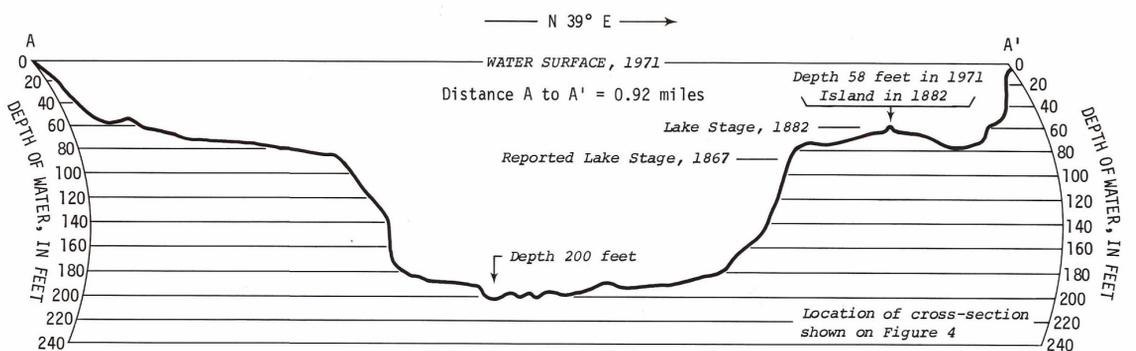


Figure 6.- Cross-section of Big Soda Lake as recorded by the fathometer

Base from aerial photographs, 1:19,000, 1971
Cartography by C. Bosch

HYDROLOGIC RECONNAISSANCE OF BIG AND LITTLE SODA LAKES, CHURCHILL COUNTY, NEVADA

By
F. Eugene Rush
1972

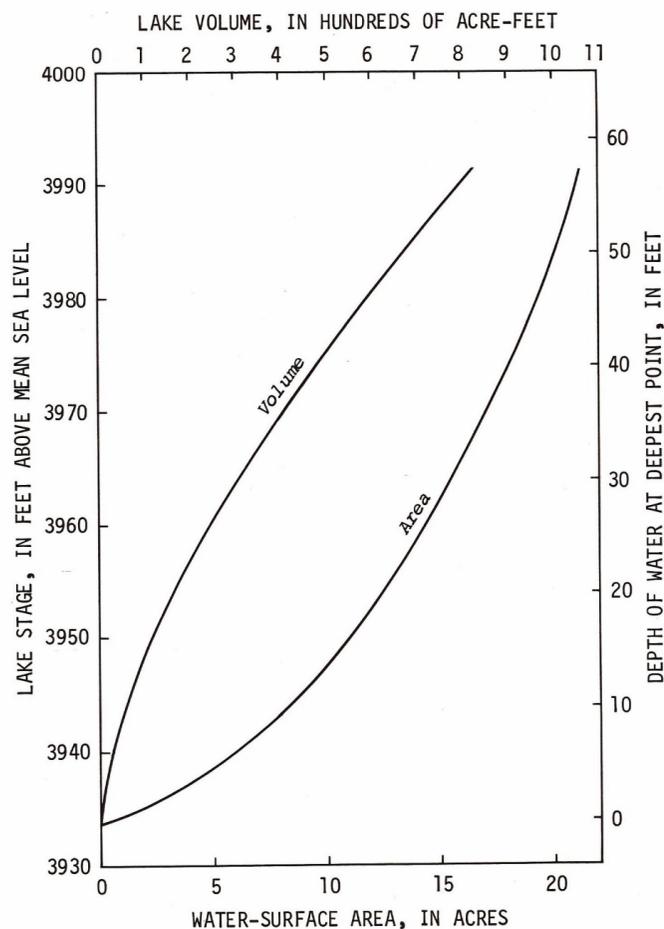


Figure 7.— Stage-area-volume relations for Little Soda Lake

SALT CONTENT

The principal dissolved constituents of Big Soda Lake, in descending order of abundance, are sodium, chloride, and sulfate. Little Soda Lake is different in that of these three dominant ions, sulfate has the highest concentration, followed by sodium. These ions make up approximately 90 percent of the total dissolved solids in both lakes. The naturally occurring ground water in the nearby alluvium, the principal source of water for the Soda Lakes prior to the Newlands Project, was described by Stabler (1904). He reports the following conditions near the Soda Lakes: (1) chloride and sulfate concentrations less than 100 mg/l, (2) springs of good quality in the lake craters, (3) dug well in the crater of Little Soda Lake had dissolved-solids content of 320 mg/l, (4) driven well in crater of Big Soda Lake had dissolved-solids content of 400 mg/l, and (5) depth to water in nearby alluvium generally less than 25 feet.

Changes in the salt content of the Soda Lakes are summarized in table 1. Both the concentrations of salt and the total tonnages of salt are shown to be decreasing with time. The dilution is due principally to increased volumes of water in the lakes. The decrease in total salt content probably is due principally to flushing of salt from the lakes to adjoining aquifers, as ground water moves through the lake.

The removal of soda from Big Soda Lake could not account for the large decrease in total salt tonnages shown in table 1. On the basis of data by Breese (1968, p. 26), it is calculated that a maximum of about 20,000 tons of soda was removed from the lake during 1868-93. Soda extraction continued until the lake began to rise in 1906. It is unlikely that the total tonnage of soda removed would account for any more than a very small part of the 900,000-ton loss of salt during 1869-1967, as determined from table 1.

Data reported by Chatard (1890, p. 48) indicate that Big Soda Lake probably was homogenous; that is, unstratified, under native conditions. More recent data collected by Hutchinson (1937, p. 75) and Breese (1968, table 7) indicate stratification of chemical constituents and water temperature into upper and lower zones, with an intervening transition zone. The upper zone of less salty water (table 1) extends to a depth of about 30 feet below the lake surface. The lower zone of more salty water is below a depth of about 120 feet. The transition zone between 30 to 120 feet is intermediate in composition. The temperature of the

water below a depth of 120 feet was essentially constant at 62°F (17°C) in August 1967 (fig. 8). The chemical stratification probably is caused by a greater dilution of lake water near the surface than at depth as more ground water flows through the upper zone of the lake.

The rate of freshening has decreased in recent years. If we assume that the tonnage of chloride was nearly constant from 1867 to 1906 at about 970,000 tons (table 1) and decreased to 670,000 tons in 1933 and to 630,000 tons in 1967, then from 1906 to 1933 the average annual loss in chloride was about 11,000 tons. For 1933-67, the rate of loss was only about 1,200 tons per year. One implication is that if the present hydrologic conditions of Big Soda Lake persist, the chloride content, and possibly other chemical characteristics as well, may nearly stabilize in a few years.

EVAPORATION AND WATER BALANCE

The net evaporation rate (evaporation of 4.3 feet for fresh water, reduced to about 90 percent by the effects of salinity (Harbeck, 1955, fig. 5), minus precipitation of 0.3 foot) for Big Soda Lake under native conditions may have averaged about 1,000 acre-feet per year (that is, 3.6 feet x 270 acres). At the present stage, the evaporation may average about 1,500 acre-feet (that is, 4 feet x 385 acres). For native conditions, the net evaporation was offset by an equal amount of subsurface inflow; however, under recent equilibrium conditions and with a continuous loss of salt from the lake with high stage, a larger volume of ground water is moving into the lake; part of it leaving as subsurface outflow, carrying dissolved chemical constituents with it. The amount of thermal water entering the lake (Breese, 1968) is not known.

During the period of stage adjustment (1906-30), the volume of Big Soda Lake increased about 20,000 acre-feet, or at an average annual rate of about 800 acre-feet. During the adjustment, the average annual net inflow would have been on the order of 2,000 acre-feet, plus the additional ground-water outflow from the lake.

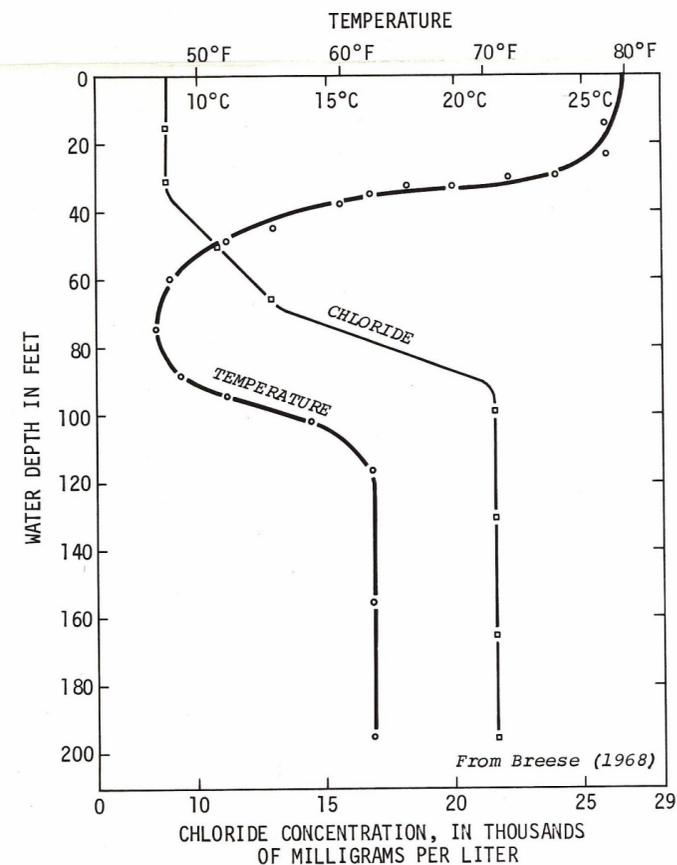


Figure 8.— Variations in water temperature and chloride concentration with depth of Big Soda Lake, August 1967

Table 1.— Summary of water-quality changes for the Soda Lakes

Date	Surface sample			Deep sample			Total lake content (thousands of tons)		
	Density ^{1/} (g/ml)	Chloride (g/l)	Dissolved-solids content (g/l)	Depth (feet)	Density (g/ml)	Chloride (g/l)	Dissolved-solids content (g/l)	Chloride content	
BIG SODA LAKE									
Aug. 1867 ^{2/}	1.0975	39.4	125	--	--	--	--	960	2,800
Sept. 1882 ^{3/}	1.0995	45.7	129	100	1.098	44.3	126	970	2,800
-- 1927 ^{4/}	a1.02	8.77	27.4	--	--	--	--	--	--
July 1933 ^{5/}	1.022	8.20	b26	192	1.066	27.3	b87	670	--
Aug. 1958 ^{6/}	1.017	7.70	24.7	--	--	--	--	--	--
Oct. 1963 ^{7/}	a1.02	7.39	25.7	--	--	--	--	--	--
-- 1967 ^{7/}	1.021	7.60	25.3	197	1.050	21.5	61.8	630	1,900
LITTLE SODA LAKE									
Oct. 1963 ^{7/}	--	.880	6.33	--	--	--	--	.93	6.7
Jan. 1967 ^{7/}	--	.699	5.51	--	--	--	--	.78	6.2

- a. Estimated from dissolved-solids and chloride data.
 b. Estimated from density and chloride data.
 1. Density in July 1915: Big Soda Lake, 1.0563 at 27°C; Little Soda Lake, 1.0051 at 27°C (Lee and Clark, 1916, p. 666).
 2. Based on data from King (1877, v. 2, p. 747).

3. Based on data from Chatard (1890, p. 48).
 4. Based on data from Miller, Hardman, and Mason (1953, p. 34).
 5. Based on data from Hutchinson (1937, p. 75).
 6. Based on data from Whitehead and Feth (1961, table 1).
 7. Based on data from U.S. Bureau of Reclamation.

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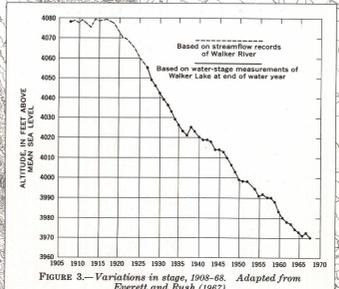
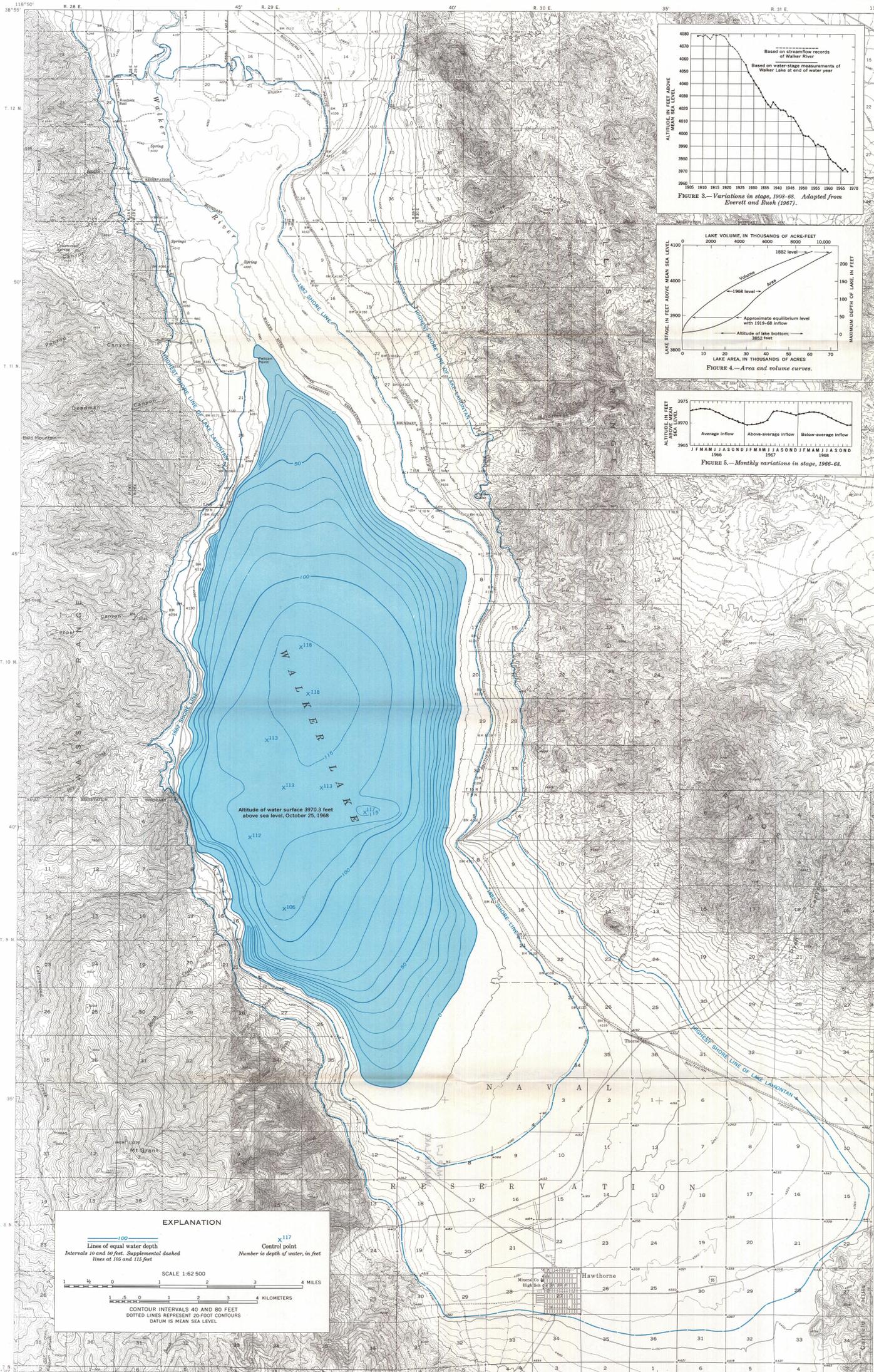


FIGURE 3.—Variations in stage, 1882-68. Adapted from Everett and Rush (1967).

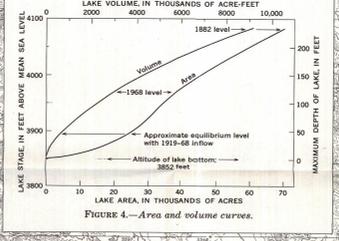


FIGURE 4.—Area and volume curves.

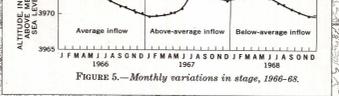


FIGURE 5.—Monthly variations in stage, 1966-68.

INTRODUCTION

Walker Lake is about 60 miles southeast of Carson City, Nev., as shown in figure 1. The town of Hawthorne, Nev., is near the south end of the lake. The lake is used principally for recreation—boating, water skiing, and fishing. However, there is considerable concern for the future of the lake because of its rapid stage recession. The altitude of the lake's surface has receded from about 4,083 feet in 1882 to 3,970 feet in 1968 and the surface area has decreased from nearly 69,000 acres to 38,000 acres. The maximum depth in 1882 was about 118 feet. At the present rate of recession, the lake may not reach new equilibrium until the lake level recedes another 70 feet, at which time the area might be approximately 25,000 acres and the depth 40 feet.

The purpose of this report is (1) to define the area and volume configurations of the lake and (2) to show the effects of recession processes on lake stage, area, volume, and dissolved-solids content.

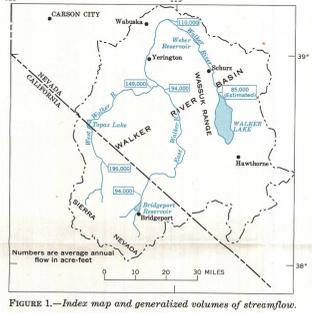


FIGURE 1.—Index map and generalized volumes of streamflow.

SETTING

Walker Lake is the lowest permanent body of water in the Walker River Basin (fig. 1), one of the 14 hydrographic regions of Nevada (Rush, 1968) and part of the Great Basin. The Walker River Basin, which includes about 4,000 square miles of Nevada and California, is drained mostly by Walker River and its two principal tributaries, East and West Walker Rivers. Most of the streamflow originates in the Sierra Nevada of California. Between the water-source areas and Walker Lake the flow is greatly reduced by diversions for irrigation and seepage losses to the ground-water system.

Within the basin, the Sierra Nevada rises to a maximum altitude of 12,440 feet, with crests commonly above 10,000 feet. Topaz Lake (fig. 1), on the West Walker River, has a surface altitude of 4,978 feet; Bridgeport Reservoir on the East Walker River, 6,427 feet. The confluence of East and West Walker Rivers is at an altitude of 4,490 feet; Weber Reservoir spillway, on the Walker River, is at 4,204 feet. The surface of Walker Lake on October 25, 1968, was at an altitude of 3,970.3 feet, the reference datum and stage for this report.

Almost all the streamflow in the basin is derived from precipitation. Precipitation in the basin is greatest in the Sierra Nevada, where the average annual quantity is as much as 50 inches. At Bridgeport, Calif., the average annual precipitation is about 13 inches. Yerington, Nev., and Hawthorne receive an average of about 5 inches per year. The estimated average annual precipitation on the lake is 4 inches (Everett and Rush, 1967, p. 7). Figure 1 shows the generalized distribution of streamflow at selected points.

BASIC DATA

Bathymetric soundings at 877 points were made in July 1957 by the Department of the Navy, Twelfth Naval District, San Bruno, Calif. The lake surface in July 1957 was 3,971.2 feet above sea level. To adjust the data to the 1968 level, 20.9 feet was subtracted from the 1957 data, and the water depths were rounded to the nearest foot.

Lake-stage data have been collected by the U.S. Geological Survey in cooperation with the Nevada Department of Conservation and Natural Resources, beginning in 1908 (no record for period 1909-26), and are published in the following Geological Survey publications.

Period	Publication
1908, 1927-50	Water-Supply Paper 1314, Part 10
1950-60	Water-Supply Paper 1734, Part 10
1960-64	Surface Water Records of Nevada, 1961, 1962, 1963, and 1964
1964-68	Water Resources Data for Nevada, 1965, 1966, 1967, and 1968

The summary of streamflow for the Walker River Basin is based on data also published in the above publications, and on interpretations by Everett and Rush (1967, p. 26), and Lamke and Moore (1965, table 1). Lake-stage data prior to 1908 are from Harding (1965, p. 143-150) and Russell (1885, p. 70).

LAKE AREA AND STAGE

Walker Lake, together with Pyramid and Honey Lakes, are remnants of Lake Lahontan, a larger Quaternary age lake. According to Snyder and others (1964), Lake Lahontan had a maximum area of 8,665 square miles, mostly in northwestern Nevada, and was sustained by a 42,322-square-mile drainage basin. The lake's maximum surface altitude was about 4,380 feet. Walker Lake now occupies part of what was the southernmost arm of Lake Lahontan.

The deepest part of Lake Lahontan was at the present site of Pyramid Lake (100 miles northwest of Walker Lake), where the maximum depth was about 920 feet (Harris, 1970). At the present location of Walker Lake, as well as on the Carson, Smoke Creek, and Black Rock Deserts (broad alluvial areas north of Walker Lake), the maximum depth of water was about 530 feet. With increased dryness, water levels receded, and Walker River Basin became hydrologically isolated from Lake Lahontan.

Figure 2 shows a generalized interpretation of long-term fluctuations in precipitation, starting with the year 800 A.D. Until 1810, the wet periods generally were longer and had greater departures from average conditions than dry periods. This suggests that Walker Lake may have remained at a rather high level prior to 1810. A drought period, 1810 to 1860, was followed by a wet period from about 1860 to 1920.

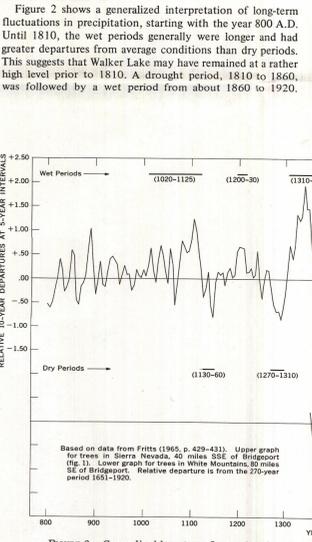


FIGURE 2.—Generalized long-term fluctuations in precipitation as interpreted from data on dated tree-ring widths.

Figure 3 indicates the wet period probably ended about 1918. Morrison (1964, p. 106) described the climate of the area as being very dry in the 1840's, and abnormally wet from 1850 to 1923. The author, therefore, concludes that during the first part of the 19th century, the lake stage probably receded, and thereafter the stage probably rose to a maximum late in the century. Harding (1965, p. 145) indicated that the lake stage may have been as low as 4,035 feet above level in 1845, and that if then rose to about 4,082 feet in 1862 and 4,089 in 1868. The levels in 1862 and 1868 were based by Harding on information that the Indian Agent's home and garden at Schurz, Nev., (fig. 1) were flooded by a rising Walker Lake. James Long, Bureau of Indian Affairs, Stewart, Nev., (oral commun., 1969), after reviewing Bureau documents for that period concluded that the Indian Agent's home was situated at an altitude of about 4,200 feet in the area now occupied by Weber Reservoir (fig. 1), which was the principal area of agricultural development prior to construction of the Weber Reservoir dam, and therefore at least 5 miles from Walker Lake and (2) the Indian Agent's home and garden probably were flooded by Walker River due to rapid spring and summer snowmelt in the Sierra Nevada, rather than by Walker Lake.

By 1882 the water level of Walker Lake was at an altitude of about 4,083 feet, as computed from data presented by Russell (1885, p. 15) and using the Navy's bathymetric data as altitude control. Harding (1965, p. 143) computed a lake-surface altitude of 4,086 feet on the basis of Russell's data. Russell reported a maximum sounded depth of 225 feet, but unknown to him and Harding, this was not at the deepest part of the lake, as determined from the Navy's bathymetric data. In 1882, the lake probably had a maximum depth of 231 feet, an area of about 68,600 acres, and a volume of about 9,000,000 acre-feet (fig. 4).

In 1908, the U.S. Geological Survey made the first of a series of lake-stage observations, at which time the altitude of the lake surface was 4,078 feet, or 5 feet lower than the 1882 stage. The next known observation was not made until 1927, but by that time the lake level had receded 23 feet to an altitude of 4,055 feet. Because a wet period in the Sierra Nevada continued through most of the intervening 20 years (fig. 2), much of the recession probably was caused by large-scale river diversions for irrigation. From about 1918 to 1968, the trend in stage has been a marked recession, averaging about 2 feet per year. Figure 3 shows lake stage for the period 1908-68.

The general stage recession of Walker Lake during most years has been interrupted by small, short-term rises in stage. For example, figure 5 shows monthly stage data for 1966-68, a 3-year period that included above-average, near-average, and below-average inflow.

WATER BUDGET

A water budget for a lake relates the various components of water in flow, outflow, and change in storage:

$$I_{R} + I_{L} + I_{GW} + P = E - \Delta S \quad (1)$$

For Walker Lake, where the change in storage has been an overall marked depletion, equation (1) is modified as follows:

$$I_{R} + I_{L} + I_{GW} + P = E - \Delta S \quad (2)$$

where the elements of inflow are: I_{R} , inflow from Walker River; I_{L} , local surface-water inflow; I_{GW} , ground-water inflow; and P , precipitation directly on the lake surface. Because no known subsurface flow occurs to adjacent valleys the only element of outflow is evaporation from the lake surface, E , and ΔS is the decrease in storage of the lake associated with lake-stage recession.

An approximate average annual water budget for the lake was compiled for the period 1919-68, the 50-year period during which most of the historical recession of stage had occurred.

I_{R} was estimated to average about 85,000 acre-feet per year, on the basis of streamflow records for stations at Wabuska, Nev., and Schurz (fig. 1), on streamflow seepage losses and diversions, and on ground-water recharge and underflow estimated by Everett and Rush (1967).

I_{L} was estimated by Everett and Rush (1967, p. 9) to be 3,000 acre-feet per year. I_{GW} was computed from published information (Everett and Rush, 1967) to be about 3,000 acre-feet per year. P was computed to be about 17,000 acre-feet per year. E was computed by using figures 3 and 4, data from Harding (1965, p. 147), and data from Köhler and others (1954, pl. 1) to be about 200,000 acre-feet per year. ΔS , as computed from figures 3 and 4, was about 110,000 acre-feet per year.

Equation (2) produces the following result:

$$I_{R} + I_{L} + I_{GW} + P = 85,000 + 3,000 + 3,000 + 17,000 = 108,000 \text{ acre-feet per year}$$

$$E - \Delta S = 200,000 - 110,000 = 90,000 \text{ acre-feet per year}$$

Imbalance = 18,000 acre-feet per year

Because of errors in estimates or unresolved hydrologic problems, the equation does not balance. Neither side of the equation was considered more accurate than the other; so the average of the two values, or about 100,000 acre-feet, was selected to approximate both inflow and outflow plus decrease in storage.

During the period 1909-18, the lake stage was nearly constant (fig. 3) and in order to maintain this high stage, average annual inflow from all sources probably was more than 250,000 acre-feet. That part entering the lake from Walker River probably was at least 80 percent of the total inflow, or at least 200,000 acre-feet per year. Inflow from Walker River during the period 1860-1908 probably was even greater, because of the previously discussed rise in lake stage.

DISSOLVED-SOLIDS CONTENT OF THE LAKE

In 1882, Russell (1885, p. 69) collected and analyzed what probably was the first sample of Walker Lake water (table 1). Because of the high stage of the lake at that time, the lake water probably had about its lowest solute concentration since the dry period of early 19th century. Not until 1930 were additional samples collected (Miller and others, 1953, p. 38).

The main factor controlling the variation in concentration of dissolved solids in the lake is the volume of lake water that dilutes the dissolved solids. The concentration of dissolved solids from 1882 to 1918 was probably similar to that sampled in 1882, because of the almost constant lake stage and volume during this period. By 1966 the lake volume had diminished by a factor of 2.9 and the dissolved-solids concentration increased by a factor of 3.4.

The total dissolved-solids content of the lake, in tons, had changed as dramatically as the concentration, as illustrated by figure 6.

A reconnaissance three-dimensional sampling survey of the lake was made by the Geological Survey on October 6, 1965, during which 28 samples were collected. On the basis of these data, the principal conclusions regarding the lake and its inflow at that time are as follows: (1) The lake was almost homogeneous chemically; for example, the chloride concentrations ranged from 2,020 mg/l (milligrams per liter), near the mouth of Walker River, to 2,240 mg/l. Only calcium varied appreciably; it ranged from about 2 to about 9 mg/l. (2) Water temperature decreased with water depth, as illustrated in figure 7.

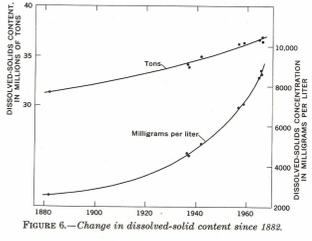


FIGURE 6.—Change in dissolved-solids content since 1882.

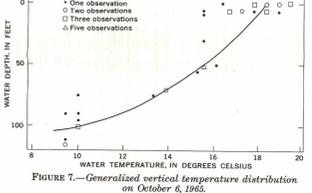


FIGURE 7.—Generalized vertical temperature distribution on October 6, 1965.

OUTLOOK FOR THE FUTURE

To maintain Walker Lake at its 1968 stage of 3,970 feet above sea level, an average annual inflow from Walker River of nearly 140,000 acre-feet would have to be sustained. This would provide a lake having a volume of about 3 million acre-feet and an area of about 38,000 acres, as shown in figure 4. Since 1918 the average annual inflow from the river has been about 85,000 acre-feet; obviously, present-day inflow will not be enough to maintain the present lake stage, area, and volume. The lake stage, therefore, will continue to recede until average annual evaporation is decreased enough to balance, rather than exceed, the average annual inflow from all sources. For example, if the inflow of Walker Lake to the lake is sustained at its present long-term average rate (about 85,000 acre-feet per year), equilibrium would be reached when the lake had an area of roughly 25,000 acres, a stage of about 3,896 feet above sea level, and a volume of about 600,000 acre-feet. This equilibrium stage would be some 70 feet below the 1968 stage and the maximum depth of the lake would be only about 40 feet, compared to a 1968 depth of 118 feet. When general equilibrium is reached, lake stage would continue to fluctuate in a narrow range related to annual variations in inflow.

Projecting the slope of the stage curve in figure 3, the estimated equilibrium stage (3,896 feet) would be reached between the years 2010 and 2060. Because upstream irrigation demands may increase, the average annual inflow from Walker River may decrease, thereby hastening the recession of the lake and consequently lowering its equilibrium level.

What will be the future dissolved-solids content of Walker Lake? The obvious answer is larger tonnages and high concentrations. Projecting trends shown in figure 6, the anticipated amount of dissolved solids at the estimated equilibrium stage (3,896 feet) may be about 40 million tons, and the dissolved-solids concentration would be nearly 50,000 mg/l.

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TABLE 1.—Chemical analyses of Walker Lake water.
[Constituents in milligrams per liter, except as indicated; analyses by U.S. Geological Survey.]

Date of collection	September 1882	February 18, 1966
Water temperature (°C)	15.2	7.5
Silica (SiO ₂)	5
Calcium (Ca)	4.2
Magnesium (Mg)	124
Sodium (Na)	858
Potassium (K)	160
Bicarbonate (HCO ₃)	1,640
Carbonate (CO ₃)	486
Sulfate (SO ₄)	1,530
Chloride (Cl)	588
Fluoride (F)	20
Boron (B)	21
Dissolved solids (calculated)	92,560
Hardness as CaCO ₃	8,610
Specific conductance (micro-mhos per cm at 25°C)	12,000
pH	9.3

^aEstimated.

HYDROLOGIC REGIMEN OF WALKER LAKE, MINERAL COUNTY, NEVADA

By
F. Eugene Rush
1970

Base from U.S. Geological Survey Hawthorne, 1955; Mt. Grant, 1956; and Gilles Canyon and Schurz, 1964.

INTERIOR—GEOLOGICAL SURVEY, WASHINGTON, D.C.—1970-70304

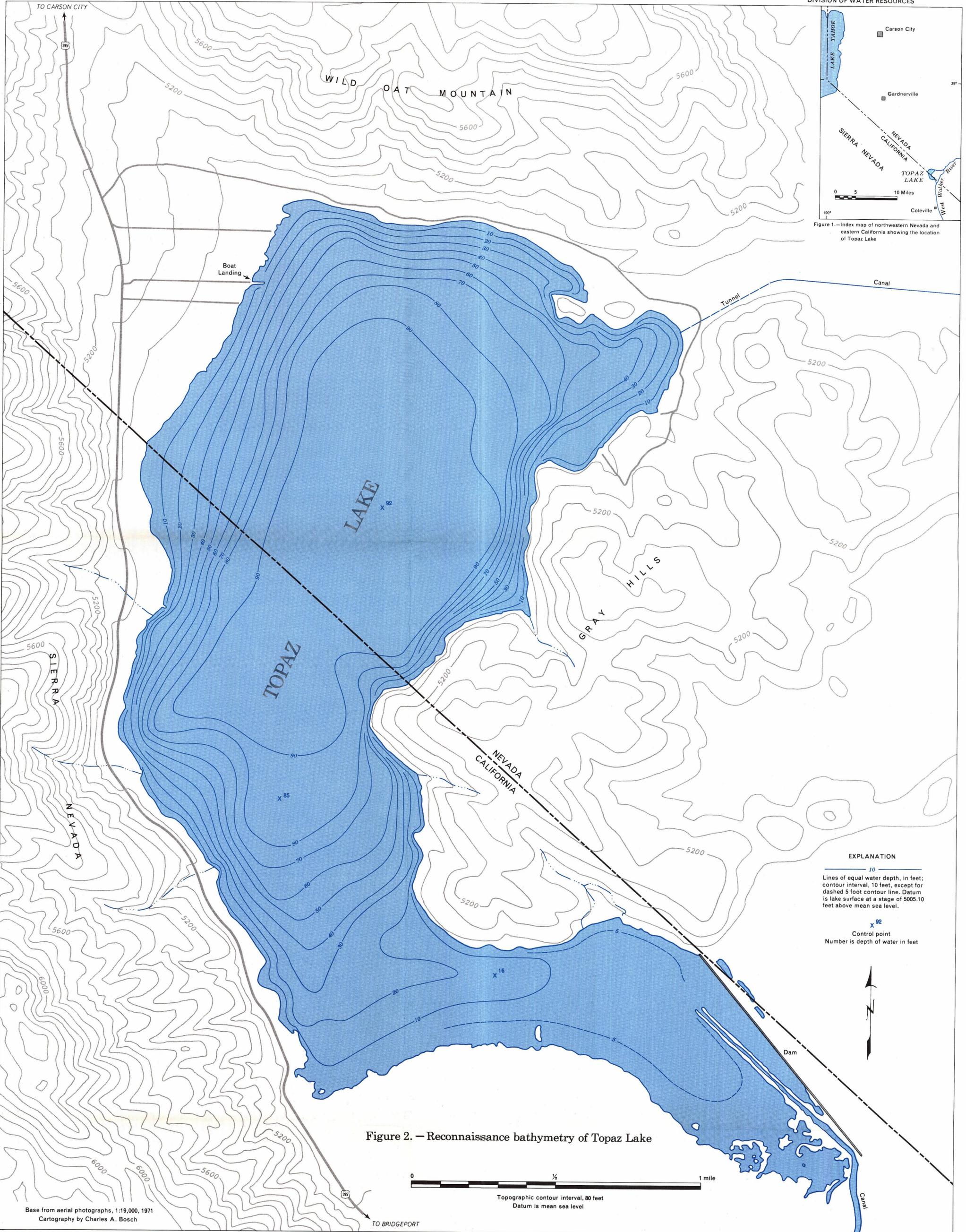


Figure 2. — Reconnaissance bathymetry of Topaz Lake

Figure 1. — Index map of northwestern Nevada and eastern California showing the location of Topaz Lake

BATHYMETRIC RECONNAISSANCE OF TOPAZ LAKE, NEVADA AND CALIFORNIA

By
F. Eugene Rush
and
Victor R. Hill
1972

INTRODUCTION

Topaz Lake is 36 miles south-southwest of Carson City, Nevada, on the California-Nevada State line, as shown in figure 1. The valley in which the lake lies is a topographically closed basin, having an area of about 14 square miles. Mountains surround the valley floor on all but the southeast side, where the topographic divide is lowest. Here the divide is a narrow, but flat alluvial surface at an altitude of about 5,000 feet. The lowest point on the valley floor, the lake bottom, is at an altitude of 4,913 feet.

The mountains on the west are the timber-covered Sierra Nevada that locally reach an altitude of nearly 9,000 feet. The sagebrush- and piñon pine-covered mountain on the north and the sagebrush-covered mountains on the east have altitudes generally less than 6,000 feet.

Under native conditions, a small lake named Alkali Lake occupied the central part of the valley floor. This lake was sustained at a generally low stage by runoff from within the basin. It is physically possible that the West Walker River, south of the basin, could have changed course from time to time and flowed into Topaz Lake basin, across the alluvial part of the divide, described above. Inflow would have stopped when the lake began to spill southeastward out of the basin across the area now occupied by the low dam (fig. 2). Whether this actually happened is not known.

Beginning in 1921, water from the West Walker River was conveyed in a feeder canal across the low alluvial part of the divide to the basin, forming a storage reservoir. The enlarged body of water, which covered most of the valley floor, was renamed Topaz Lake. Stored water is released at the northeast end of the lake through a 1,200-foot long tunnel and a canal which conveys water back to the West Walker River. All water released from the lake passes through the tunnel, because there is no spillway. Downstream, the released water is used principally for irrigation by the Walker River Irrigation District, the developer of Topaz Lake.

The lowest practical lake stage at which water can be released through the tunnel is 4,972.3 feet, which leaves 59 feet of lake depth as dead storage. Releasable storage prior to 1937 was about 45,000 acre-feet. In that year, construction of a dam at the southeast end of the lake (fig. 2) increased this volume to a reported 59,440 acre-feet at a stage of 5,005.0 feet, the maximum operating stage of the lake.

In addition to being used for storage of irrigation water, Topaz Lake is used to an increasing extent for boating and fishing. However, the general absence of sandy beaches probably is discouraging to many potential recreational users.

Variations in lake stage are summarized in figure 3. At the time of the bathymetric survey, the stage of the lake was 5,005.10 feet, or about at its maximum operating stage.

BATHYMETRY

A continuously recording, electronic fathometer was used to measure the depth of Topaz Lake, during 19 traverses. The results of the survey are summarized in figures 2 and 4. The dimensions of the lake at the time of the survey, for the extremes and average in stage for the period 1922-71, and of the unreleasable pool, are summarized in table 1.

In 1940, the U.S. Geological Survey determined the stage-volume relation. The 1971 bathymetric survey produced virtually identical results.

Table 1.--Dimensions of Topaz Lake at various stages

Time or condition	Dimension				
	Depth (feet)	Stage (feet)	Area (acres)	Volume (acre-feet)	Releasable storage (acre-feet)
Time of survey in June 1971	92	5,005.10	2,400	125,000	60,000
Maximum stage	92	5,005.35	2,410	126,000	61,000
Minimum stage	59	4,972.25	1,500	65,000	0
Average stage	76	4,988.90	1,800	92,000	27,000
Maximum stage of unreleasable pool	59	4,972.30	1,500	65,000	0

Other bathymetric surveys have been completed or are planned, as follows:

Lake or Reservoir	Publication
Pyramid Lake	USGS HA-379
Walker Lake	USGS HA-415
Lahontan Reservoir	Nev. DWR Info. Ser. 9
Big and Little Washoe Lakes	Nev. DWR Info. Ser. 10
Big and Little Soda Lakes	Nev. DWR Info. Ser. 11
Topaz Lake	Nev. DWR Info. Ser. 12
Rye Patch Reservoir and Upper and Lower Pitt-Taylor Reservoirs	Nev. DWR Info. Ser. 13
Marlette and Spooner Lakes	Nev. DWR Info. Ser. 14
Weber Reservoir	Nev. DWR Info. Ser. 15
Wild Horse Reservoir	Nev. DWR Info. Ser. 16

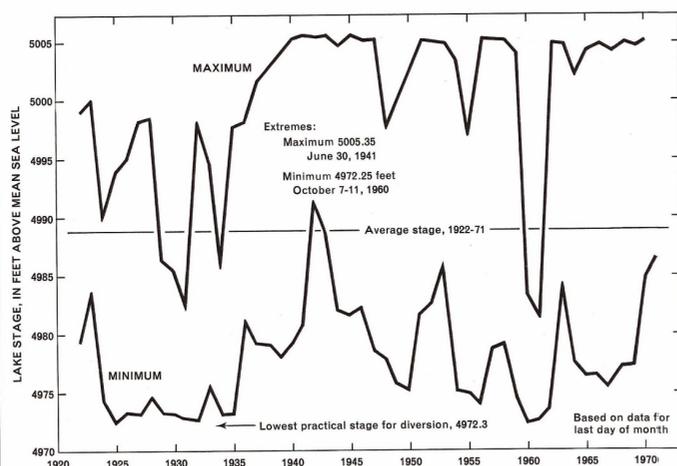


Figure 3.--Variations in stage of Topaz Lake

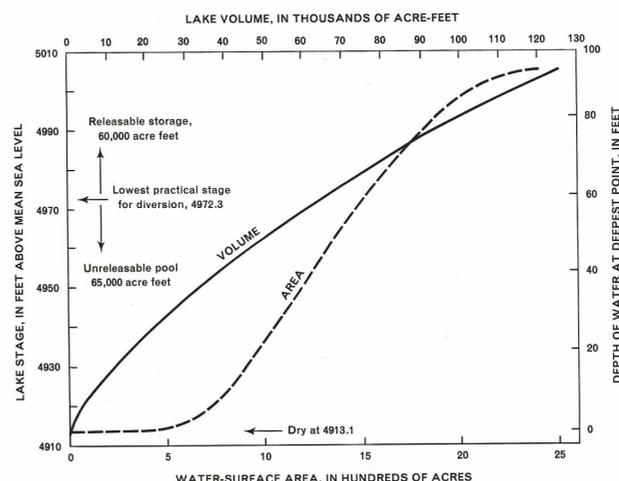


Figure 4.--Stage-area-volume relations for Topaz Lake

118°17'30"
40°35'00"

MATCH TO NORTHERN PART OF RESERVOIR (See front)

INTRODUCTION

Rye Patch, Upper Pitt-Taylor, and Lower Pitt-Taylor Reservoirs are 110 miles northeast of Reno, Nev., and about 35 miles north of Lovelock, Nev. (fig. 1). The reservoirs are on the narrow alluvial floor of a north-trending segment of the Humboldt River Basin at an altitude of about 4,100 feet.

The principal source of water for the reservoirs is snowmelt in the mountains of north-central and northeastern Nevada. The Humboldt River Basin has an area of about 17,000 square miles, of which 95 percent is tributary to the reservoirs. The Humboldt River flows westward and southwestward across northern Nevada to the reservoir sites. Rye Patch Reservoir is an on-channel reservoir, but Upper Pitt-Taylor Reservoir receives water through a canal extending from the Humboldt River; diversion is about 12 miles upstream. Lower Pitt-Taylor Reservoir is filled from Upper Pitt-Taylor Reservoir. To drain the Pitt-Taylor Reservoirs, water is released from the upper reservoir to the lower reservoir, and then to Rye Patch Reservoir. Releases of water from Rye Patch Reservoir are used for irrigation of land near Lovelock (fig. 1), as described by Everett and Rush (1965).

Flow of the Humboldt River a few miles above Rye Patch Reservoir is measured by the Geological Survey at the station, Humboldt River near Inlay, and is summarized in figure 2. The relatively small amount of water diverted to the Pitt-Taylor Reservoirs (fig. 3) is not included in these amounts. Outflow from the reservoirs also is measured by the Geological Survey just below the dam at the station, Humboldt River near Rye Patch.

Irrigation near Lovelock began in 1862; however, full-season water supply was not available because of the erratic natural flow of the Humboldt River. In 1909, work was begun by the Humboldt-Lovelock Irrigation Light & Power Co. on construction of the Pitt-Taylor Reservoirs, named for two of the principals in the company. The design capacity of the reservoir system, to be used for power, irrigation, and domestic use, was 57,000 acre-feet. Water was first turned into the reservoirs in November 1912 when the completed capacity was 30,000 acre-feet. In 1915, levees and dams were raised to 19 feet for the upper reservoir and to 38 feet for the lower reservoir to increase storage capacity to 50,000 acre-feet. In 1935 the upper reservoir has a rated capacity of 29,000 acre-feet and the lower reservoir, 21,000 acre-feet. The reservoir system was sold to the Pershing County Water Conservation District in 1947. Figure 3 shows the diversions to the reservoirs.

Rye Patch Dam construction was begun by the U.S. Bureau of Reclamation in 1935 and was completed and water impoundment began in the following year. The dam is an earth-fill, rock-faced structure with a structural height of 75 feet, a hydraulic height of 63 feet, a crest length of 914 feet, and a volume of 356,000 cubic yards (U.S. Bur. Reclamation, 1961, p. 261-265). The design capacity is 190,000 acre-feet when 12-inch flashboards are in place on top of the spillway gates and the reservoir is full at a stage of 4,134.0 feet above mean sea level. The reservoir is used to some extent for fishing and boating, but its principal use is for storage of irrigation water.

Figure 4 shows the variation in the stage of Rye Patch Reservoir, since its construction. By comparing this graph to water-diversion data for the Pitt-Taylor Reservoirs (fig. 3), it can be seen that since about 1940, water is put into the Pitt-Taylor Reservoirs only during periods when supplemental storage for Rye Patch Reservoir is needed. There are three reasons for this: (1) Pitt-Taylor Reservoirs are less efficient than Rye Patch Reservoir for storage because of their generally lower ratio of stored-water volume to evaporation, as shown in figure 5; (2) water stored in the Pitt-Taylor Reservoirs leaches salts from the reservoir beds, reducing the quality of the stored water, whereas the quality of water stored in Rye Patch Reservoir remains better; and (3) in case of earthquake, Rye Patch Reservoir is considered by the Pershing County Water Conservation District to be a safer reservoir because of higher quality of its dam's design.

To convert staff gage readings to reservoir stage, in feet above mean sea level, the following factors should be added to the staff gage readings: Upper Pitt-Taylor Reservoir, 4,098.0 feet and Lower Pitt-Taylor Reservoir, 4,099.6 feet. These factors are based on the assumption that the stage values, in feet above mean sea level, for Rye Patch Reservoir are essentially correct. These relations were determined during this study by running levels between reservoir water surfaces. Prior to this survey, slightly different factors for the Pitt-Taylor Reservoirs were used by the reservoir operators.

The quality of water released from Rye Patch Reservoir is summarized in table 1.

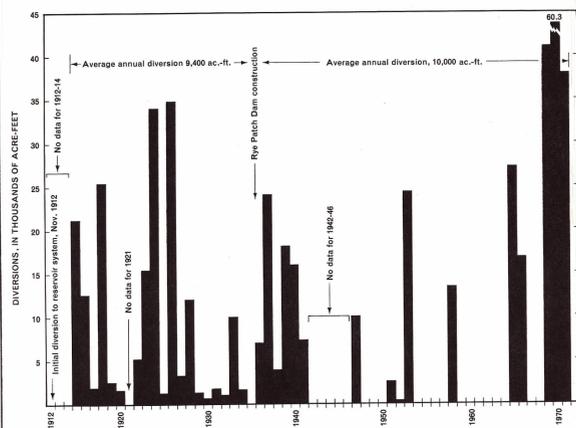
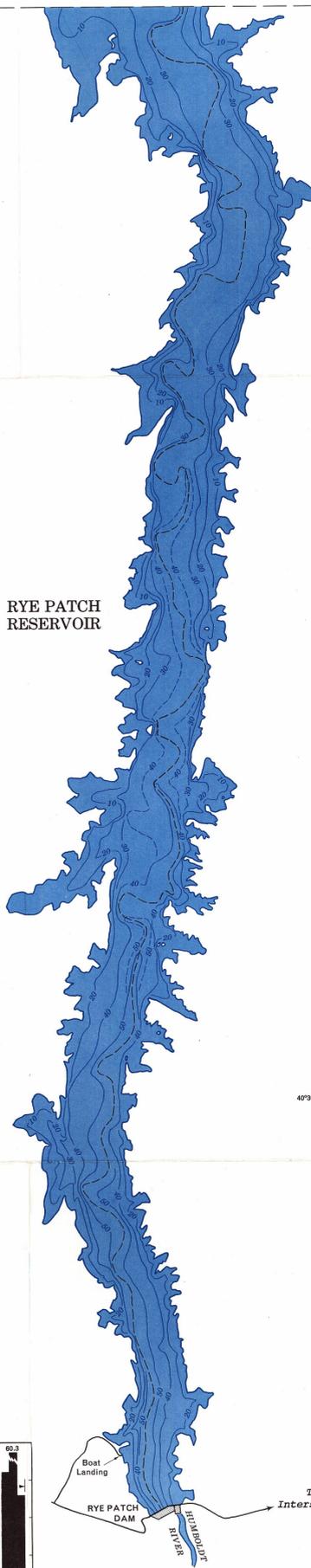


Figure 3.—Annual diversions of water from the Humboldt River to the Pitt-Taylor Reservoirs



RYE PATCH RESERVOIR

Other bathymetric surveys have been completed or are planned, as follows:

Lake or Reservoir	Publication
Pyramid Lake	USGS HA-379
Walker Lake	USGS HA-415
	Nev. DWR
Lahontan Reservoir	Info. Ser. 9
Big and Little Washoe Lakes	Info. Ser. 10
Big and Little Soda Lakes	Info. Ser. 11
Topaz Lake	Info. Ser. 12
Rye Patch Reservoir and Upper and Lower Pitt-Taylor Reservoirs	Info. Ser. 13
Marlette and Spooner Lakes	Info. Ser. 14
Weber Reservoir	Info. Ser. 15
Wildhorse Reservoir	Info. Ser. 16

Table 1.—Summary of water-quality data for Rye Patch Reservoir [At gaging station 1,000 feet below Rye Patch Dam]

Constituent	Concentration range (in milligrams per liter unless otherwise specified)
Water year 1970 (9 samples)	
Discharge when sample collected	13-694 cfs
Temperature	36-71°F
Silica	2.2-21.5°C
Calcium	33-36
Magnesium	38-45
Sodium	15-16
Potassium	125-135
Bicarbonate	16-18
Carbonate	296-312
Sulfate	0-8
Chloride	72-87
Fluoride	86-106
Nitrate	0.5-1.1
Boron	0.0-0.2
Dissolved solids	260-580 ug/l
Hardness	562-604
Specific conductance	161-174
pH	893-950 micromhos
Total phosphorus (as PO ₄)	7.7-8.5
Aldrin, DDD, DDT, dieldrin, endrin, heptachlor, heptachlor epoxide, lindane, 2,4-D, 2,4,5-T, chlordane	0.00 ug/l
Silvex	0.01-0.06 ug/l
Period of record	
Specific conductance, (1951-58, 1959-61, 1962-69):	
Maximum daily	4,010 micromhos on Sept. 2, 1954
Minimum daily	384 micromhos on June 24, 1956
Water temperatures:	
Maximum (1951-54, 1956-58, 1959-61, 1962-68)	78°F (25.5°C) on Sept. 21, 1958
Minimum (1951-54, 1956-58, 1959-61, 1962-67, 1968-69)	33°F (0.5°C) on many winter days

BATHYMETRY

A continuously recording, electronic fathometer was used to measure water depth on the 43 traverses of Rye Patch Reservoir, 21 traverses of Upper Pitt-Taylor Reservoir, and 16 traverses of Lower Pitt-Taylor Reservoir, in June 1971. The results of the surveys are summarized in figures 5, 6, 7, 8, and 9. Dimensions of the reservoirs at reference stages used in figures 5, 6, 7, 8, and 9 are summarized in table 2.

Table 2.—Dimensions of the reservoirs when full^{1/2}

Dimension	Pitt-Taylor Reservoirs	
	Rye Patch Reservoir	Upper Lower
Stage, in feet above mean sea level	4,134	4,158 4,147
Staff gage, in feet	--	60.0 47.4
Maximum depth, in feet	61.5	17.5 26
Area, in acres	11,400	2,070 2,570
Volume, in acre-feet	171,000	24,200 22,200
Limit of drawdown when down-system reservoir is at maximum stage:		
Stage, in feet above mean sea level	4,134	4,134
Maximum depth, in feet	6.5	13
Area, in acres	1,360	360
Volume, in acre-feet	4,800	1,600

1. Pershing County Water Conservation District and the State Engineer may consider the Pitt-Taylor Reservoirs to be full at stages other than those listed in this table.

SEDIMENTATION

The extent of sedimentation within Rye Patch Reservoir has been estimated by two methods: (1) the computed reduction in reservoir storage since the reservoir was formed, and (2) by estimating the amount of sediment fill on the reservoir bottom.

In 1936, the Bureau of Reclamation (1961, p. 261) computed that the reservoir, at a stage of 4,134 feet, had a storage capacity of 190,000 acre-feet. The results of the bathymetric survey of 1971 indicated a capacity of 171,000 acre-feet, or 19,000 acre-feet less. The difference in estimates may be the result of sedimentation, small errors in storage estimates, or both.

Comparing the bathymetric data of 1971 to an unpublished Bureau of Reclamation topographic map of the reservoir site compiled in 1934, sedimentation of from 1 to 5 feet along the axis of the reservoir is indicated. Figure 6 shows supplemental (dashed) contours that represent the configuration of the reservoir bottom in 1934 as compared to 1971 findings (the solid-line contours). The sedimentation, based on this comparison, is estimated to be on the order of 16,000 acre-feet. Sources of sediment would be the inflow of the Humboldt River and the collapse of vertical banks of the reservoir at high stage, as observed during the bathymetric study. This collapse has apparently enlarged the water-surface area of the reservoir from about 11,100 acres, when the reservoir was formed, to its 1971 area of 11,400 acres at a 4,134-foot stage.

From the two methods of estimating sedimentation within Rye Patch Reservoir, it is concluded that between 15,000 and 20,000 acre-feet of sediment has been deposited during the period 1936-71, or between about 30 and 40 million tons, assuming the unit weight of the deposits to be about 90 pounds per cubic foot. According to P. A. Glancy, hydrologist, U.S. Geological Survey (oral commun.), the 36-year sediment yield of the upstream 16,000-square-mile basin, in order to supply this amount to the reservoir from the Humboldt River, assuming a trap efficiency of nearly 100 percent, would be only about 50 to 70 tons per square mile per year. No estimates of sedimentation in the Pitt-Taylor Reservoirs have been made.

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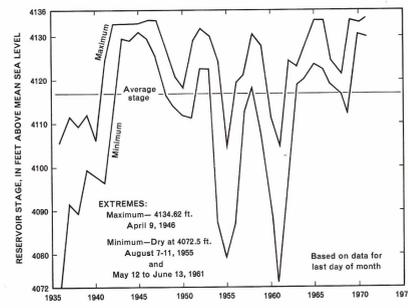


Figure 4.—Annual variation in stage of Rye Patch Reservoir

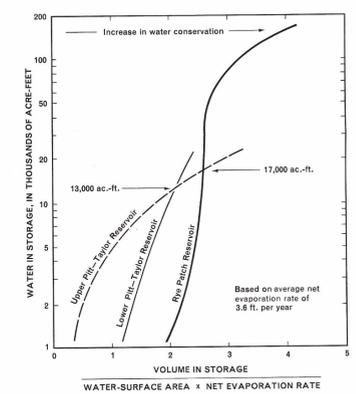


Figure 5.—Storage efficiency of the reservoirs

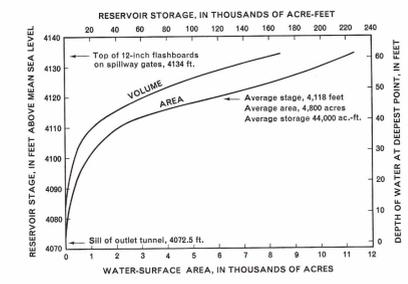


Figure 7.—Stage-area-volume relations for Rye Patch Reservoir, June 1971

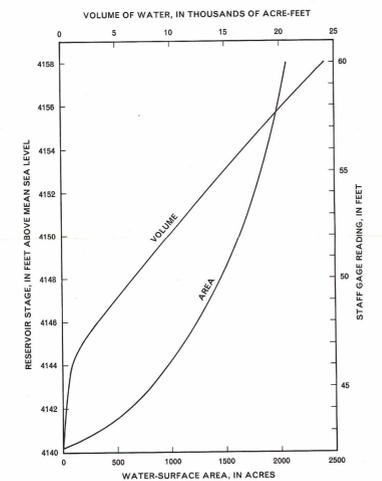


Figure 8.—Stage-area-volume relation for Upper Pitt-Taylor Reservoir

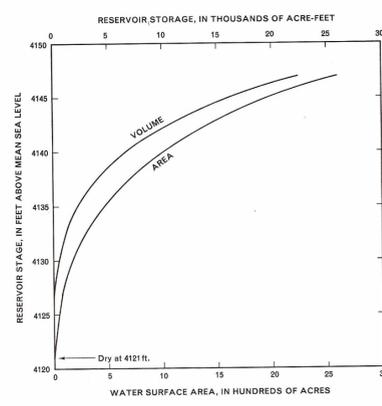
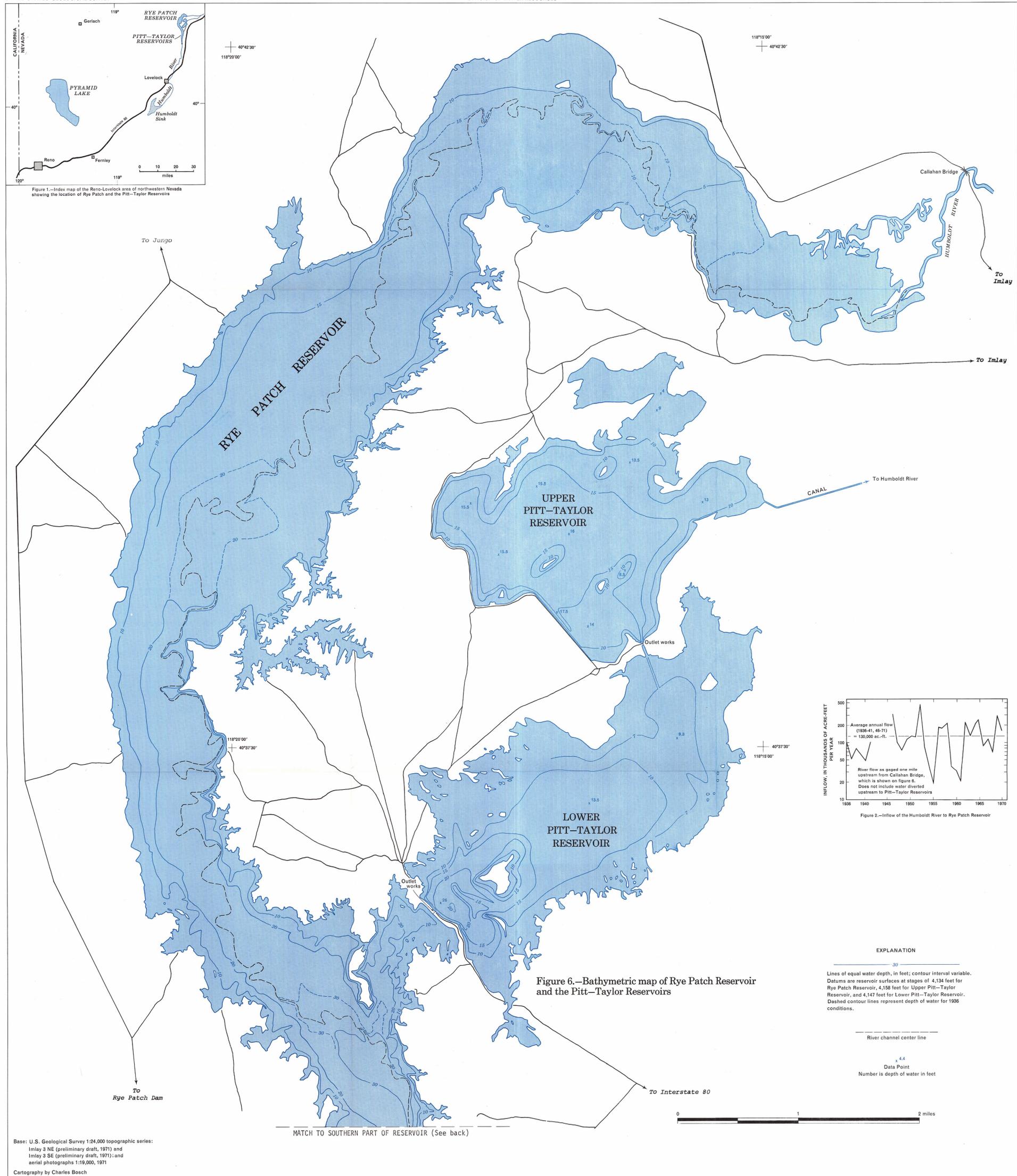


Figure 9.—Stage-area-volume relation for Lower Pitt-Taylor Reservoir



BATHYMETRIC RECONNAISSANCE OF RYE PATCH RESERVOIR AND THE PITT-TAYLOR RESERVOIRS, PERSHING COUNTY, NEVADA

By
F. Eugene Rush and Bruce L. Rice
1972

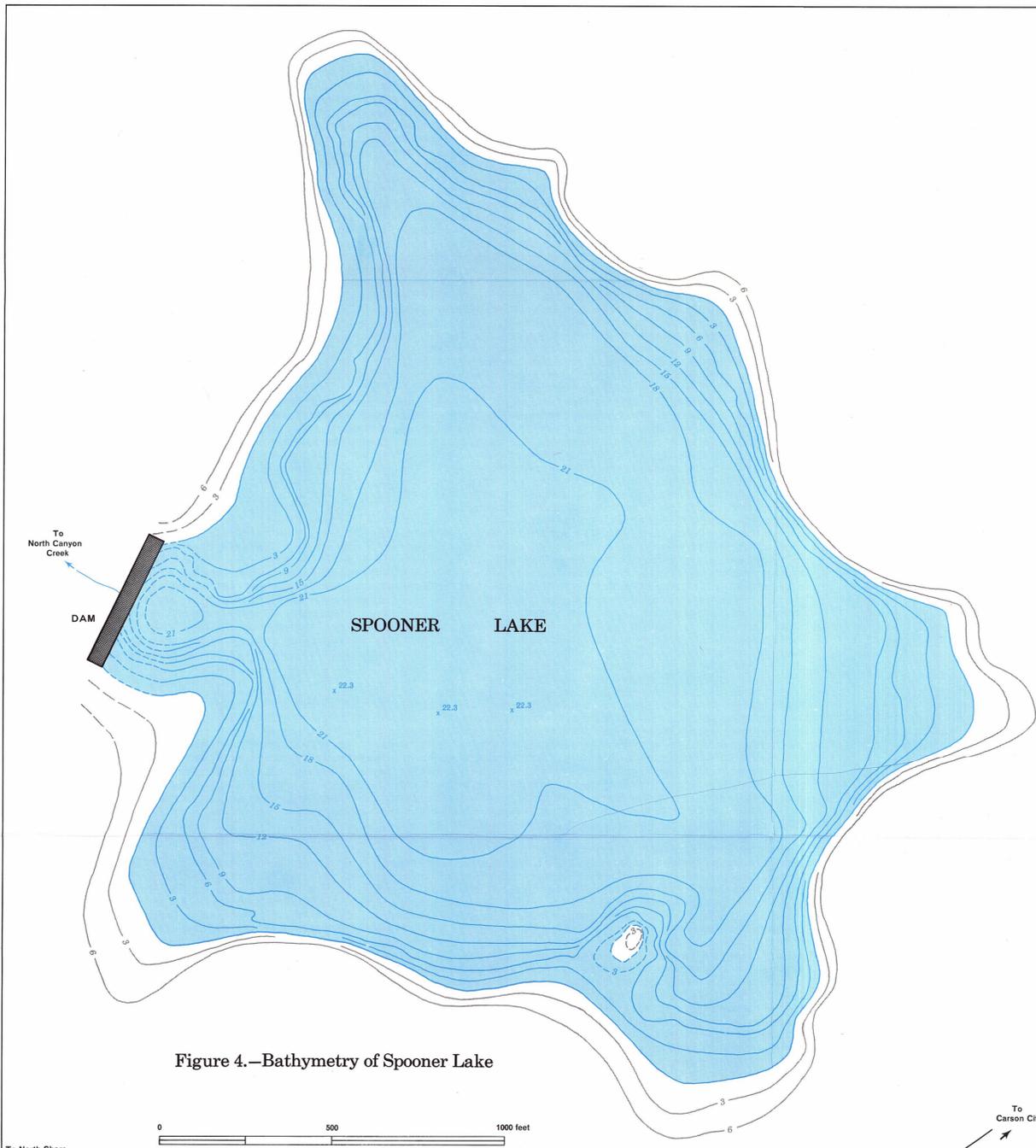
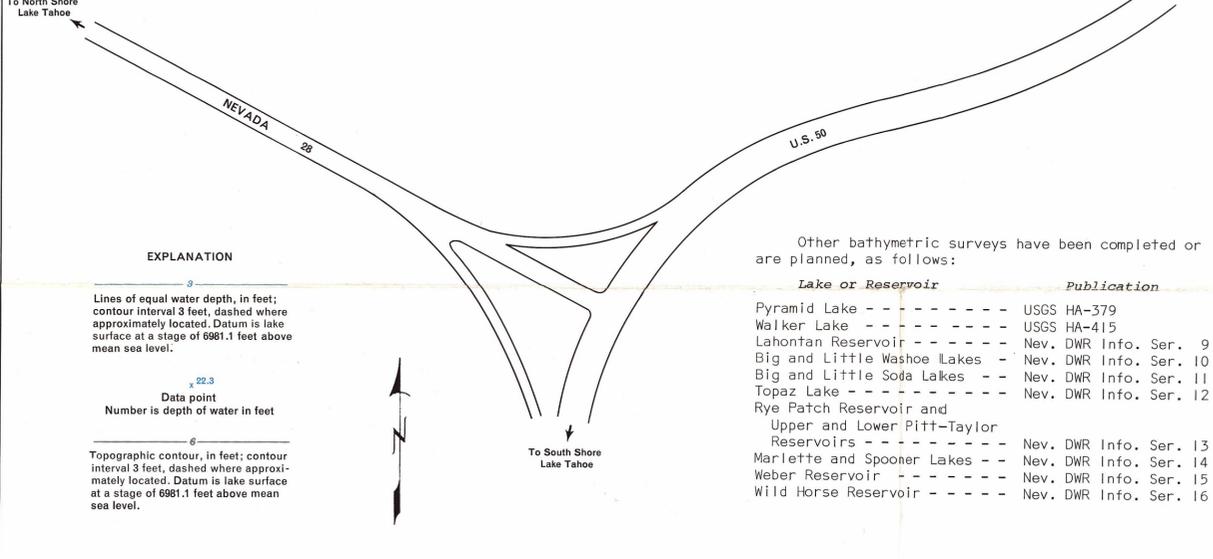


Figure 4.—Bathymetry of Spooner Lake



Dimensions of Spooner Lake at various stages are given in the following tabulation:

Time or condition	Stage (feet above mean sea level)	Area (acres)	Volume (acre-feet)
During bathymetric survey, July 1, 1971	6,976.8	84	1,030
December 2, 1971	6,975.4	80	910
At stage shown in figure 4	6,981.1	95	1,410
Overflow stage of dam	6,982.9	100	1,580

With reference to the above tabulation, it can be seen that the bathymetric map of Spooner Lake (fig. 4) was drawn at a stage about 4 feet higher than actual conditions during the time of the survey. This was done for two reasons: (1) to make the bathymetric map more meaningful by "filling" the lake, and (2) to provide the Nevada Department of Fish and Game with information on the character of the lake if more water were to be stored or if the dam were to be raised or replaced with a higher structure. The data needed to determine this increase in stage above existing field conditions in the summer of 1971 were obtained by instrument surveying peripheral to the lake.

At the time of the survey, water was leaking through or beneath the dam of Spooner Lake at an estimated rate of between 100 and 200 gallons per minute. This leakage is probably about of the same general magnitude as the evaporation from the lake, if both are evaluated on the basis of average annual quantities. Therefore, a somewhat higher stage probably could be maintained, if this leak were eliminated, as the leakage and evaporation probably are the two principal means of discharge from the lake.

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Galloway, J. D., 1947, *Early engineering works contributory to the Comstock*: Nevada Univ. Bull., v. 41, no. 5, 102 p.

Nevada Legislative Commission, 1969, *The Marlette Lake water system; a report on the feasibility and desirability of its retention*: Nevada Legislative Council Bureau Bull. no. 79, 103 p.

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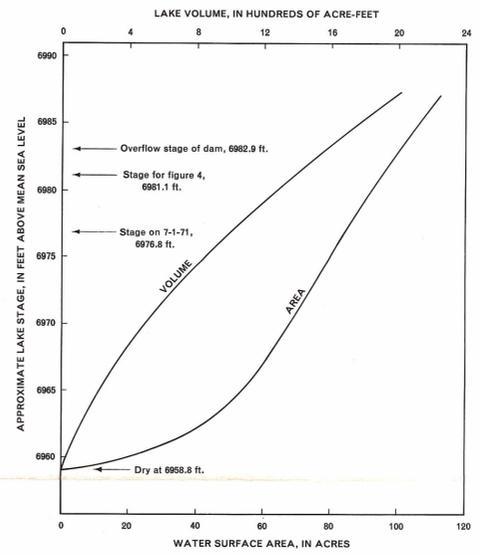


Figure 5.—Stage-area-volume relation for Spooner Lake

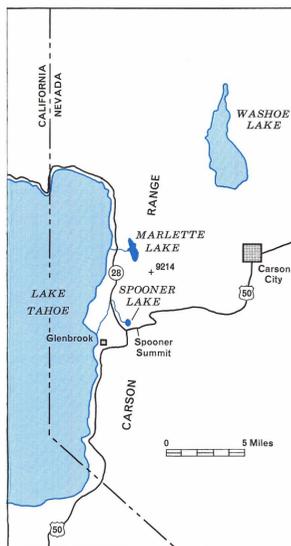


Figure 1.—Index map of the Carson City area showing the locations of Marlette and Spooner Lakes

INTRODUCTION

Marlette and Spooner Lakes are between Carson City and Lake Tahoe (fig. 1) in the Carson Range of the Sierra Nevada. Marlette Lake is at an altitude of 7,838 feet and Spooner Lake, at about 6,980 feet. Marlette Lake is about 1,600 feet higher than Lake Tahoe and about 3,200 feet higher than Carson City. Both lakes drain to Lake Tahoe; Marlette Lake by way of Marlette Creek and Spooner Lake by way of North Canyon and Slaughterhouse Creeks. Marlette Lake is in a consolidated-rock depression (graben), probably containing little sediment. Spooner Lake is in an alpine meadow area, probably underlain by a veneer of alluvium. Both lakes are surrounded by pine-covered mountain ridges and peaks extending as high as 9,000 feet at Marlette Lake and 7,800 feet at Spooner Lake.

The principal source of water for the lakes is snowmelt in their small basins; the Marlette Lake basin has an area of 3.0 square miles and Spooner Lake basin, about 1.2 square miles. Precipitation at Marlette and Spooner Lakes averages about 27 inches per year, or an average annual total of each basin of about 4,500 acre-feet and 1,700 acre-feet, respectively. Of these amounts, an estimated 60 percent runs off to the lakes. The principal use of Marlette Lake in 1971 was as a source of cutthroat trout spawn for the Nevada Department of Fish and Game and is administered by that department. Most of the land adjoining Marlette and Spooner Lakes is within the Lake Tahoe Nevada State Park.

In 1873, according to Galloway (1947) and the Nevada Legislative Commission (1969), Duane L. Bliss and H. M. Yerington (Carson and Tahoe Lumber and Fluming Company) placed a 26-foot high dirt-fill and rock dam (Scott, 1957, p. 301) across the natural outlet from a small lake, then called Goodwin Lake but later named Marlette Lake. Water from the reservoir was conveyed southward to Spooner Summit (fig. 1), by a 6-mile-long V-flume, where it was used to flume lumber 12 miles to a lumber yard a mile south of Carson City. Most of the lumber was transported to the summit by a short railroad (Lake Tahoe Railroad) extending past what is now Spooner Lake to sawmills at Glenbrook on the nearby shore of Lake Tahoe.

In 1876, the Virginia and Gold Hill Water Co. received consent from Bliss and Yerington to draw water from Marlette Lake to be conveyed to the Virginia City area, 17 miles to the northeast. The dam was raised to a reported height of 37 feet, with a length of 213 feet. The top of this dam was at an altitude of 7,828 feet, according to Walter G. Reid, Virginia City (written commun., 1971). With a maximum lake stage of 7,823 feet above sea level, the estimated storage volume was 6,100 acre-feet, according to Galloway (1947) and 5,000 acre-feet according to Reid. The Virginia and Gold Hill Water Co. built a 14-inch by 30-inch flume northward along the west side of the crest of the Carson Range, a distance of 4.38 miles, to the west portal of a tunnel extending eastward through the crest of the Carson Range. The tunnel, 3,994 feet long, connected to another flume which conveyed the water eastward to an inverted siphon and again by flume to Virginia City.

In 1959, the dam at Marlette Lake was raised 15 feet to an altitude of 7,843 feet, bringing the high-water line to 7,838 feet and the storage volume to an estimated 10,400 acre-feet (Reid, written commun., 1971). The tunnel that conveyed water for Virginia City collapsed prior to 1963, preventing the usual movement of water from Marlette Lake to Virginia City and Carson City. In 1966, a diesel-operated pump was installed to pump water from Marlette Lake through a pipeline over the crest of the Carson Range to Red House diversion dam, near Hobart Reservoir, 2 miles northeast of Marlette Lake. Pumping has been infrequent.

Spooner Lake was created prior to 1927 by the construction of a small earthen dam across a narrow part of Spooner Meadow, situated near and due north of the intersection of U.S. Highway 50 and State Route 28, as shown in figure 4. The lake was built by Charles L. Fulstone to store irrigation water. In the past, the lake was called Fulstone Reservoir and Spooner Meadow Reservoir.

Altitudes of the dam and lake are based on the approximate altitude of a nearby feature, a large rock at the north end of the dam. The top of this rock has an altitude of about 6,990.7 ± 1.0 feet.

BATHYMETRY

A continuously recording, electronic fathometer was used to measure the depths of Marlette and Spooner Lakes, making a total of 19 traverses on Marlette Lake and 17 on Spooner Lake. The results of the survey are summarized in figures 2, 3, 4, and 5. The dimensions of Marlette Lake at high-water lines with 1873, 1876, and the 1959 dams are summarized in the following tabulation:

Spillway level, 1873 dam	
Stage, approximate (feet)	7,813 ± 1
Area (acres)	270 ± 13
Volume (acre-feet)	3,400 ± 300
Maximum depth (feet)	19.4 ± 1
Spillway level, 1876 dam	
Stage (feet)	7,823
Area (acres)	525
Volume (acre-feet)	6,400
Maximum depth (feet)	29.4
Spillway level, 1959 dam	
Stage (feet)	7,838
Area (acres)	381
Volume (acre-feet)	11,800
Maximum depth (feet)	44.4

It can be seen from the data presented in this report that previous estimates of lake volume for the old and present dams were low with reference to the data resulting from this bathymetric survey.

The bottom of Marlette Lake, or that area below a depth of 35 feet, generally can be characterized as a very gently sloping plane. The plane, dipping to the northwest, may have been formed as a horizontal playalike surface that has been tilted by faulting along the western margin of the lake. The short time since faulting has limited the modification of the surface by sedimentation or erosion.

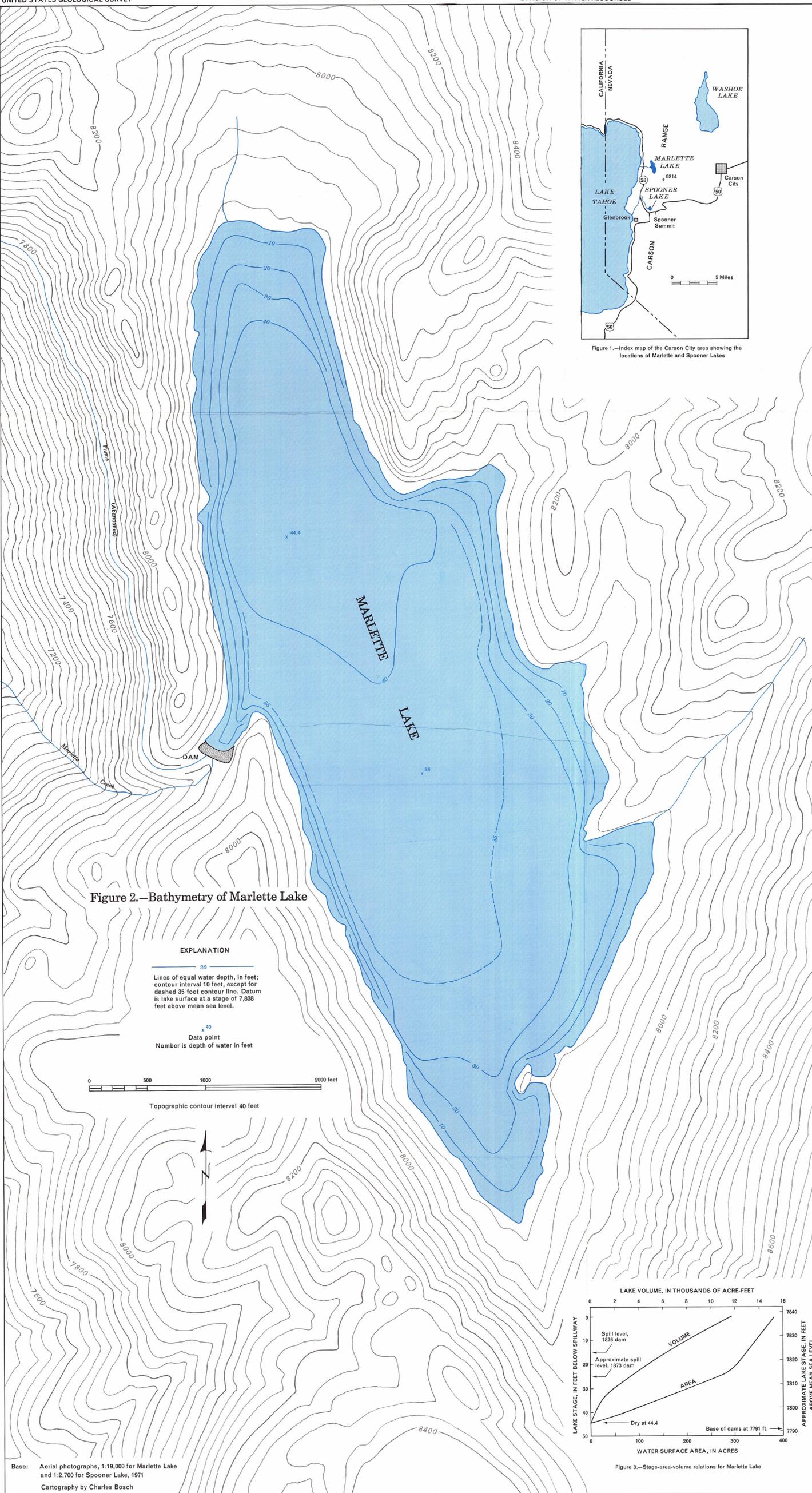


Figure 2.—Bathymetry of Marlette Lake

EXPLANATION

— 20 —
Lines of equal water depth, in feet; contour interval 10 feet, except for dashed 35-foot contour line. Datum is lake surface at a stage of 7,838 feet above mean sea level.

x 40
Data point
Number is depth of water in feet

0 500 1000 2000 feet
Topographic contour interval 40 feet

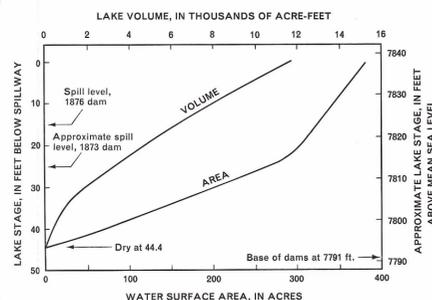


Figure 3.—Stage-area-volume relations for Marlette Lake

Base: Aerial photographs, 1:19,000 for Marlette Lake and 1:2,700 for Spooner Lake, 1971
Cartography by Charles Bosch

BATHYMETRIC RECONNAISSANCE OF MARLETTE AND SPOONER LAKES, WASHOE COUNTY AND CARSON CITY, NEVADA

By
F. Eugene Rush, Bruce R. Scott,
Patrick A. Glancy, Terrance L. Katzer
and Edwin P. Ament
1972

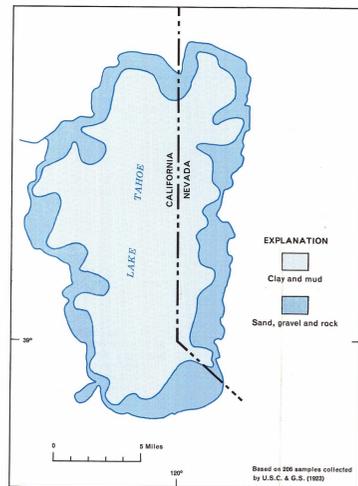


Figure 2.—General distribution of lake-bottom sediments

Table 2.—Summary for Lake Tahoe

Feature	Description
Reference stage (U.S. Bureau of Reclamation datum)	6,229 feet above mean sea level
Water-surface area	194 square miles (124,000 acres)
Nevada part	61 square miles
California part	133 square miles
Maximum depth	1,646 feet (Hyne and others, 1972, p. 1435, report 1,627 feet as deepest encountered)
Rank:	
In North America	3d deepest
In world	10th deepest
Volume	125 million acre-feet
Stage:	
Natural outlet	6,223 feet
Recorded variations (1901-71):	
Maximum	6,231.26 feet in July 1907
Minimum	6,211.74 feet in December 1934
Mean	6,226.5 feet
Maximum prehistoric (approximate)	7,000 feet (Hyne and others, 1972, p. 1435)
Length	22 miles
Width	12 miles
Shoreline (approximate)	75 miles
Nevada	30 miles
California	45 miles
Shoreline ownership (U.S. Forest Service, 1972):	
Private	55 miles
Federal	12 miles
States	8 miles
Water temperature (mostly from Crippen and Pavelka, 1970; and Lake Tahoe Area Council, 1963):	
At surface:	
Minimum (winter)	40-45°F
Maximum (summer)	65-75°F
At depth below 500 feet	Nearly constant 39°F
Freezing	Only to a minor extent in shallow, protected inlets, except Emerald Bay, which occasionally freezes over.
Water quality (Crippen and Pavelka, 1970; and Lake Tahoe Area Council, 1963):	
Dissolved solids	60-70 mg/l
Clarity	Secchi disc ¹ visible to about 120 feet; 90 percent light penetration to 120-130 feet.
Emerald Bay:	
Area	460 acres
Maximum depth	213 feet
Fannette Island:	
Area	2.5 acres
Height (approximate)	100 feet
Age (Hyne and others, 1972)	10,000 years
Origin (Hyne and others, 1972)	Scoured by Tioga glaciation

¹ White, 8-inch diameter disc.

Table 3.—Summary for dam at lake outlet

Feature	Description
Site	At Tahoe City, outlet of lake to Truckee River
Initial dam:	
Date	1870 ¹
Builder	A. W. Von Schmidt ¹
Materials	Wood crib and rock
Present dam (mostly from U.S. Bureau of Reclamation, 1961, p. 534-542):	
Construction period	1909-13 ¹
Probable builders	Floriston Pulp and Paper Company and Truckee River General Electric Company ¹
Structural height	16 feet
Hydraulic height	11 feet
Top width	11 feet
Maximum base width	19 feet
Crest length	109 feet
Outlet works	Seventeen 5-foot by 4-foot gates
Design:	
Outflow capacity	2,630 cfs
Active stage	6,223-6,229.1 feet (legal limits by Federal Court decree)
Active capacity	744,600 acre-feet (USGS)

¹ Published information on dam-construction dates and builders is inconsistent. Information presented hopefully is correct.

The basin was formed about 2 million years ago by faulting and volcanism. The basin is a graben, or down-dropped block. The lake level has been lowered by erosion of a natural dam created by tilting and faulting (Hyne and others, 1972, p. 1435).

Precipitation, the source of water in Lake Tahoe, is summarized in figure 3. The greatest precipitation is on the western and southwestern mountains. Most of the precipitation falls in the winter and spring as snow. Snow-melting and subsequent runoff to the lake occur principally in the period April-July, as shown by a graph of the runoff distribution of the Upper Truckee River (fig. 4), a major tributary to the lake at Tahoe Keys. Outflow to Truckee River occurs each year, as summarized in figure 5. Precipitation, runoff, and evaporation from the lake surface are the principal factors controlling the interrelation between the outflow to the Truckee River and the fluctuations in lake stage (fig. 6).

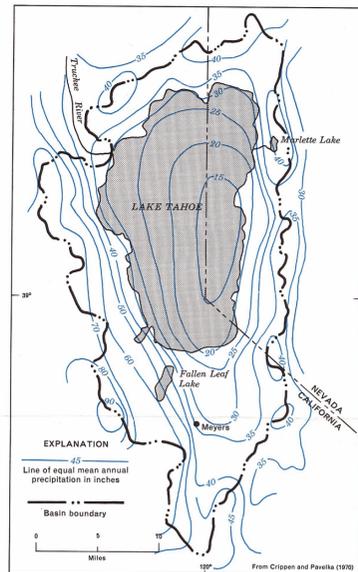


Figure 3.—Approximate mean annual precipitation in the Lake Tahoe Basin

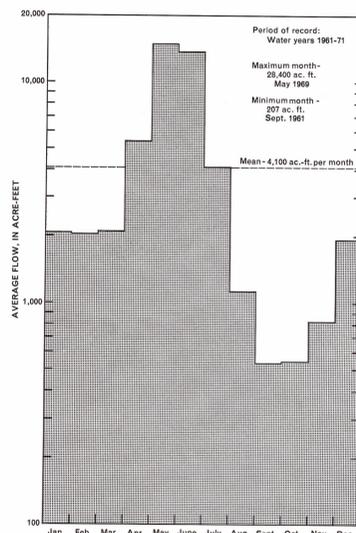


Figure 4.—Mean monthly flow of Upper Truckee River near Meyers, California

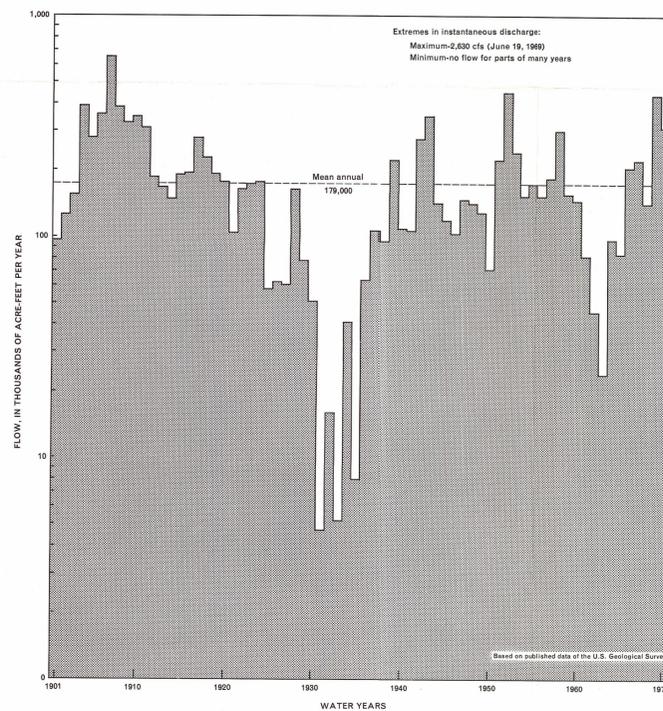


Figure 5.—Flow from Lake Tahoe to the Truckee River

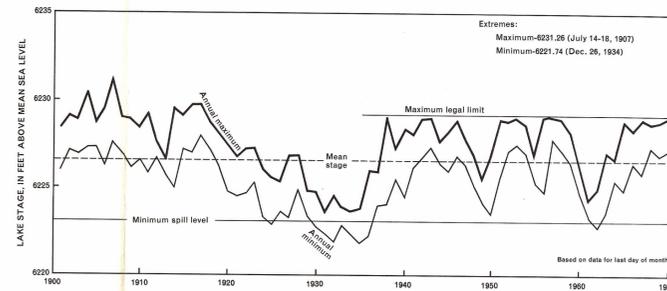


Figure 6.—Annual variations in stage of Lake Tahoe

BATHYMETRY

The bathymetric map of Lake Tahoe is based on a soundings map of the U.S. Coast and Geodetic Survey (1923). The U.S. Coast and Geodetic Survey map shows water depths for about 1,800 points. The deepest sounding in the lake, 1,646 feet and the only sounding greater than 1,600 feet, was about 6 miles due east of the Truckee River outlet and along the axis of the lake.

Several submerged mounds, as shown, range in height up to about 400 feet. Hyne and others (1972, p. 1441) describe the mounds as composed of slumped sediments, whereas Goldman and Court (1968) suggested a possible volcanic origin.

Most of the steep slopes shown by the close spacing of the bathymetric contours on the east and west sides of the lake are of fault origin. In addition, a fault scarp extends southwestward from Stateline Point, at the north end of the lake, toward the deepest point in the lake. About 2 miles north of Emerald Bay at Rubicon Point, the lake depth increases 1,300 feet in about the same horizontal distance, producing a bottom slope of about 45° from horizontal.

Stage-area-volume relations are shown in figure 7.

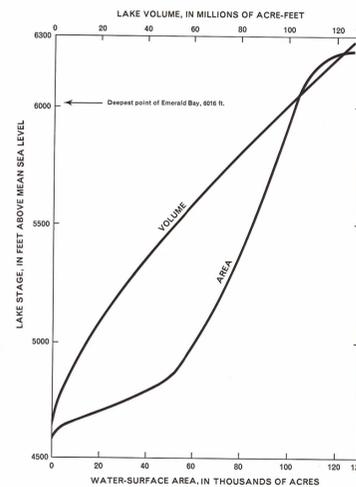


Figure 7.—Stage-area-volume relation for Lake Tahoe

WATER BUDGET

A water budget for a lake relates the various components of water inflow, outflow, and change in storage:

$$I_{sw} = O_{sw} \pm \Delta S \quad (1)$$

For Lake Tahoe, equation (1) is modified as follows:

$$I_{sw} + I_{gw} + P = O_{sw} + E + D \pm \Delta S \quad (2)$$

where the elements of inflow are: I_{sw} , inflow of all surface water; I_{gw} , inflow of ground water; and P , precipitation directly on the lake surface. The elements of outflow are: O_{sw} , surface-water outflow from the lake to the Truckee River; E , evaporation from the lake surface; and D , diversions from the lake. ΔS is change in lake storage associated with net stage change for the budget period. Ground-water seepage from the lake is believed to be negligible to nonexistent.

For the purposes of this reconnaissance, an approximate mean annual water budget for the lake can be computed by omitting ground-water inflow (I_{gw}), diversions from the lake (D), and storage change (ΔS), because these elements, though pertinent, are small in relation to the other budget elements and the total water in storage. Equation (2) therefore is modified as follows:

$$I_{sw} + P = O_{sw} + E \quad (3)$$

A 71-year period, 1901-71 was used as a base for computation. Based largely on estimates of Crippen and Pavelka (1970, p. 36), equation (3) becomes (rounded):

$$310,000 + 220,000 = 180,000 + 350,000,$$

or approximately 530,000 acre-feet of total mean annual inflow and outflow. However, because the 71-year period of record is believed to be somewhat wetter than normal, this total, when adjusted to the base periods 1919-69 and 1931-60 of average wetness, probably is more nearly 500,000 acre-feet per year.

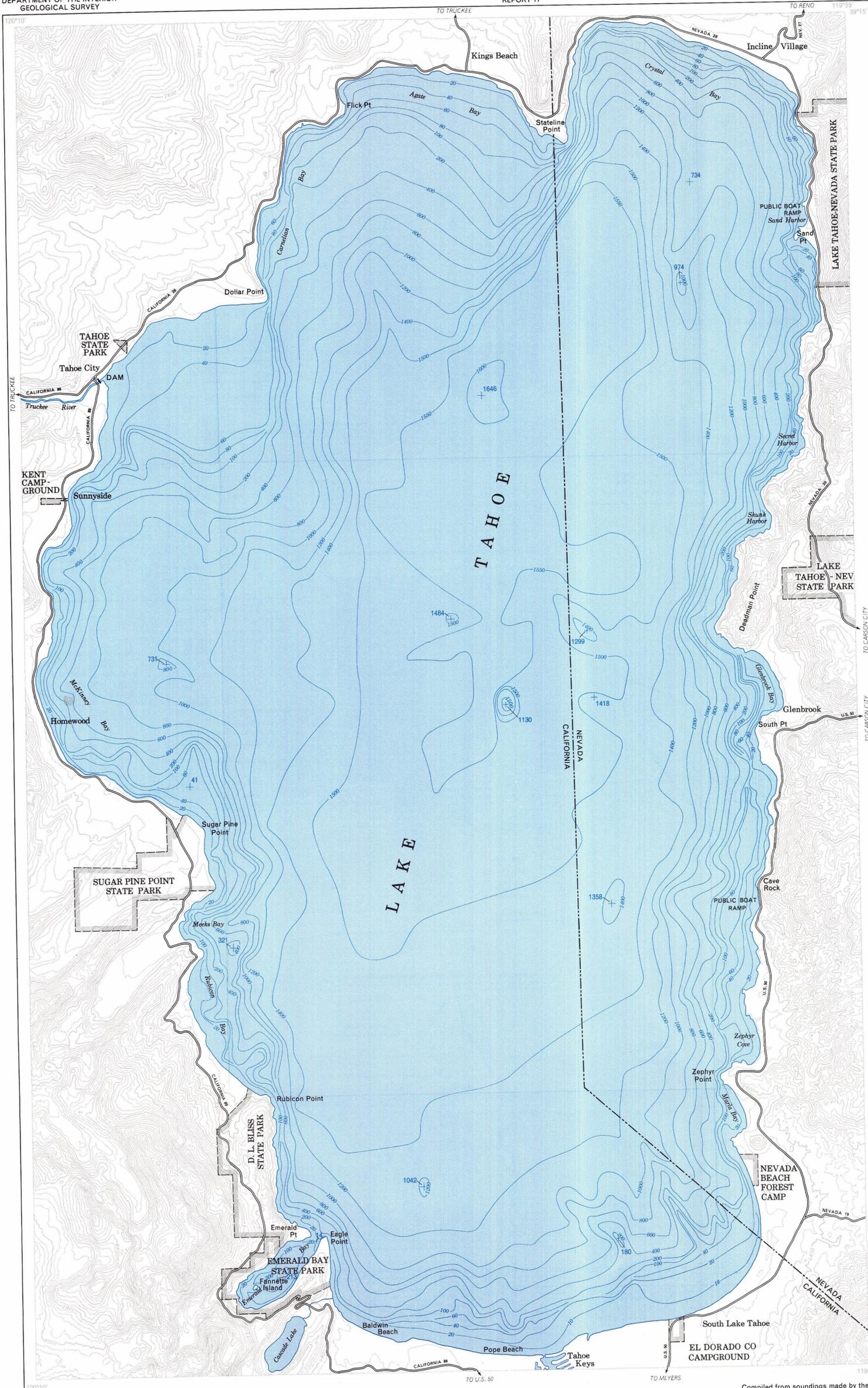
The lakes and reservoirs of this series are listed as follows:

Lake or Reservoir	Publication
Pyramid Lake	USGS HA-3791
Walker Lake	USGS HA-415
Lahontan Reservoir	Nev. DWR Info. Ser. 92
Big and Little Washoe Lakes	Nev. DWR Info. Ser. 10
Big and Little Soda Lakes	Nev. DWR Info. Ser. 11
Topaz Lake	Nev. DWR Info. Ser. 12
Rye Patch Reservoir and Upper and Lower Pitt-Taylor Reservoirs	Nev. DWR Info. Ser. 13
Marlette and Spooner Lakes	Nev. DWR Info. Ser. 14
Weber Reservoir	Nev. DWR Info. Ser. 15
Wild Horse Reservoir	Nev. DWR Info. Ser. 16
Lake Tahoe	Nev. DWR Info. Ser. 17

¹ U.S. Geological Survey Hydrologic Atlas.
² Nevada Division of Water Resources Information Series Report.

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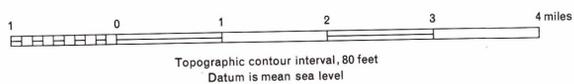


Base: U.S. Geological Survey
1:62,500 topographic series:
Carson City, Nev., 1956
Fallen Leaf Lake, Calif., 1955
Freel Peak, Nev.-Calif., 1956
Tahoe, Calif., 1955

EXPLANATION

Lines of equal water depth, in feet; contour interval variable.
Datum is lake surface at a stage of 6,229 feet above mean sea level (U.S. Bureau of Reclamation datum of 1929, supplementary adjustment of 1959)

321
+
Data point
Number is depth of water in feet



Compiled from soundings made by the
U.S. Coast and Geodetic Survey (1923)
Cartography by Charles A. Bosch

INTRODUCTION

Lake Tahoe is on the Nevada-California State line, 10 miles west of Carson City, Nevada, as shown in figure 1. The natural outflow from Lake Tahoe, high in the Sierra Nevada, is to the Truckee River, which flows through Reno to Pyramid Lake. Since the beginning of the 20th century, a part of the Truckee River flow has been diverted through a canal of the Newlands Project (U.S. Bureau of Reclamation) for use in the Carson River Basin near Fallon, Nevada (60 miles east of Carson City).

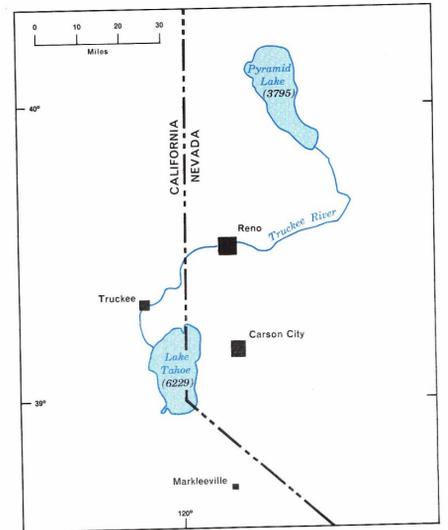


Figure 1.—Index map of northwestern Nevada and eastern California showing the location of Lake Tahoe (Numbers in parenthesis are water-surface altitudes)

Table 1 is an informational summary for the Lake Tahoe Basin. In addition, the U.S. Forest Service (1972) has published data on land-use capability. This report shows that approximately three-fourths of the land has either a high relative erosion potential or disturbance hazard. The 1970 U.S. Census indicates a total population in the basin of about 26,000. Table 2 summarizes facts about Lake Tahoe. The facts presented describe the lake at a stage of 6,229 feet above mean sea level, U.S. Bureau of Reclamation Lake Tahoe datum, which is the maximum stage regulated by use of a small dam (table 3). The Lake Tahoe datum is 1.14 feet higher than the sea level datum of 1929 used elsewhere in the area, as determined by the U.S. Geological Survey in November 1960. Reference stages of 6,223 feet and 6,225 feet have been used in other studies (Matthews and Schwarz, 1969), but for a series of bathymetric reconnaissances, of which this study is a part, the decision was made to evaluate each lake and reservoir when "full".

Table 1.—Summary for the Lake Tahoe Basin

Feature	Description
Basin area (mostly from Crippen and Pavelka, 1970)	506 square miles (324,000 acres)
Land area (approximate)	306 square miles (196,000 acres)
Water area (approximate)	200 square miles (128,000 acres)
Lakes in basin:	
Tahoe (stage 6,229 feet)	194 square miles (124,000 acres)
Fallen Leaf	1,400 acres
Marlette	381 acres
Upper and Lower Echo	330 acres
Cascade	210 acres
Spoooner	97 acres
Numerous small lakes and ponds	600 acres
Length (north-south)	40 miles
Width	18 miles
Highest altitude (Freel Peak)	10,881 feet
Lowest altitude (deepest point of lake)	4,583 feet

Generalized rock distribution in order of outcropping abundance (adapted from Crippen and Pavelka, 1970)

Granitic rocks (Sierra Nevada batholith)	East, south, and southwest ranges.
Glacial deposits	Mostly south and west of Lake Tahoe.
Volcanic rocks	Scattered, but mostly in northwest range.
Metamorphic rocks	Southwest range.
Lake beds	Along west, south, and north shores up to an altitude of about 7,000 feet.

Underlying the floor of Lake Tahoe with a thickness of at least 400 feet (Hyne and others, 1972, p. 1435). Bottom sediments are summarized in figure 2.

Land ownership January 1972 (U.S. Forest Service, 1972)	Square miles	Percent
Federal Government:		
In Nevada	28	9
In California	153	50
State Government:		
Nevada	9.5	3
California	5.6	2
Private	110	36
Total (rounded)	306	100

Weather at Tahoe City (Crippen and Pavelka, 1970):

Temperature (°F):	
Maximum	94
Minimum	-15
Mean annual	42
Average frost-free period	86 days
Mean annual precipitation	31 inches
Mean annual snowfall	18 feet

BATHYMETRIC RECONNAISSANCE OF LAKE TAHOE, NEVADA AND CALIFORNIA

By F. Eugene Rush

1973

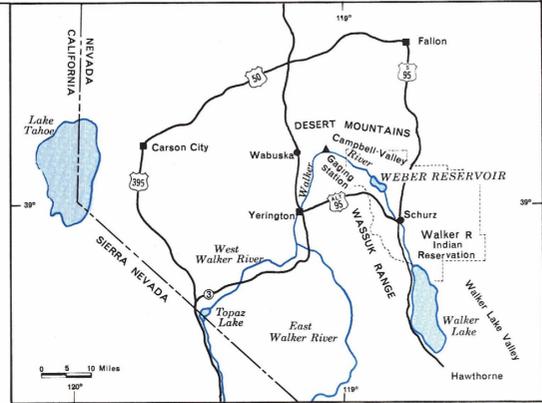
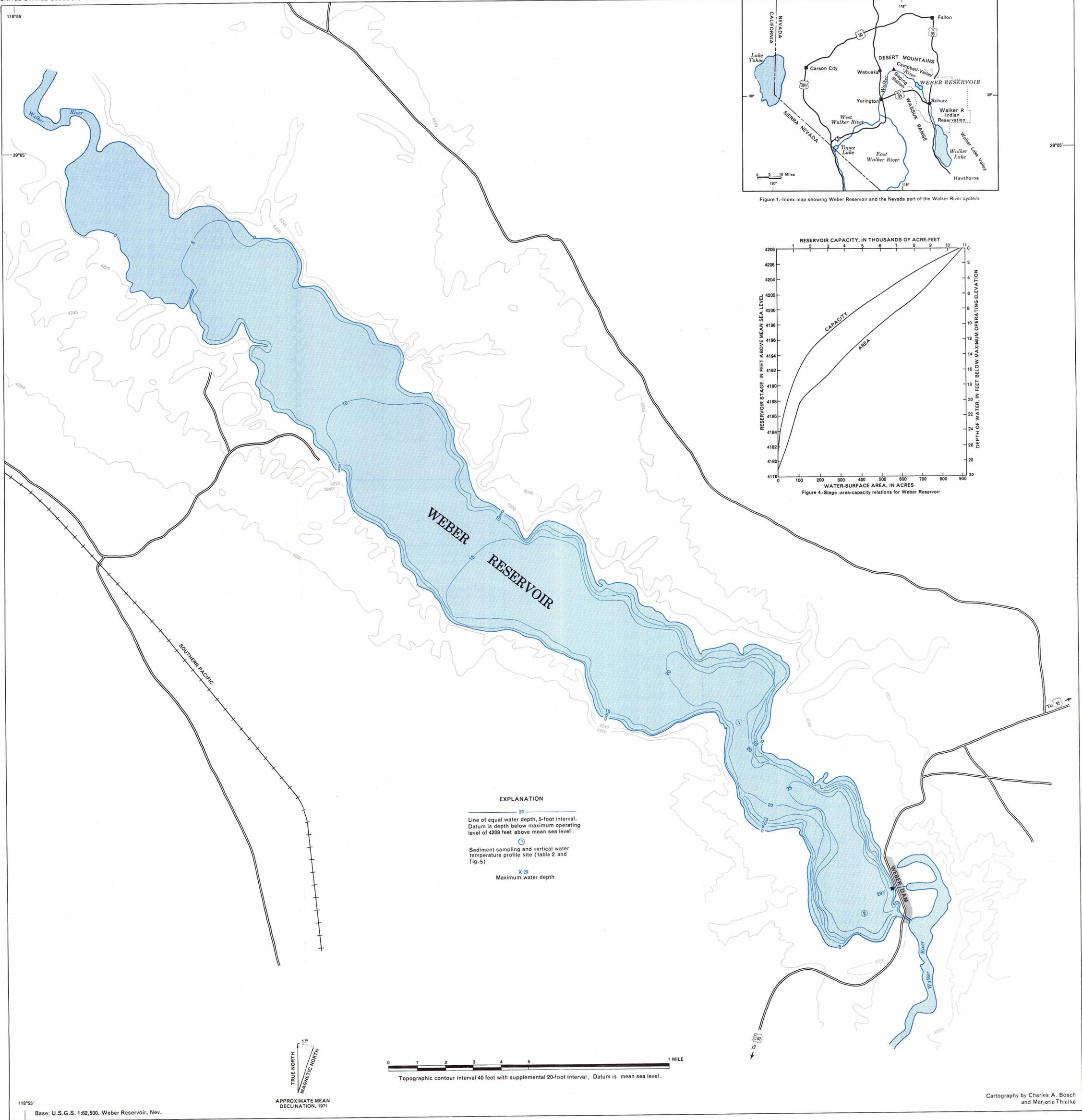


Figure 1.-Index map showing Weber Reservoir and the Nevada part of the Walker River system

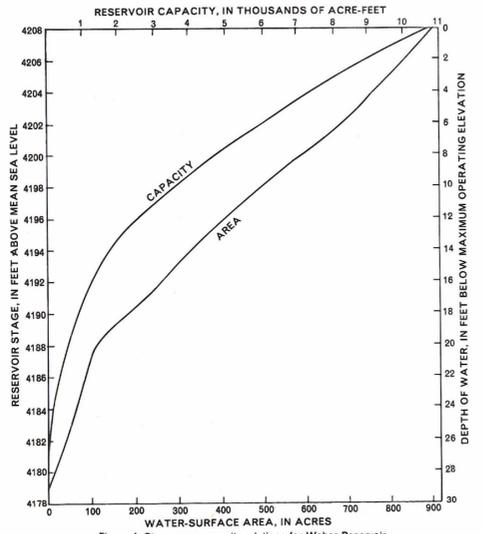


Figure 4.-Stage-area-capacity relations for Weber Reservoir

EXPLANATION

— 20 —
Line of equal water depth, 5-foot interval. Datum is depth below maximum operating level of 4208 feet above mean sea level.

①
Sediment sampling and vertical water temperature profile site (table 2 and fig. 5)

X 29
Maximum water depth

TRUE NORTH
MAGNETIC NORTH
APPROXIMATE MEAN DECLINATION, 1971

0 1 2 3 4 5 1 MILE
Topographic contour interval 40 feet with supplemental 20-foot interval. Datum is mean sea level.

118°55'
Base: U.S.G.S. 1:62,500, Weber Reservoir, Nev.

Cartography by Charles A. Bosch and Marjorie Thielke

BATHYMETRIC RECONNAISSANCE OF WEBER RESERVOIR, MINERAL COUNTY, NEVADA
By
T. L. Katzer and Lynn Harmsen
1973

INTRODUCTION

Weber Reservoir is in Campbell Valley (northern end of Walker Lake Valley) on the Walker River about 7 road miles northwest of Schurz and about 16 miles east of Yerington in Mineral County, Nevada (fig. 1). Weber Reservoir is totally within the Walker River Indian Reservation and is at an elevation of about 4,200 feet. The Desert Mountains that form the northern boundary of Campbell Valley have a maximum elevation of 6,404 feet. The northern end of the Wassuk Range, that forms the southern boundary of Campbell Valley, has a maximum elevation of 8,159 feet (fig. 1).

Shortages of Walker River irrigation water on the Walker River Indian Reservation created an interest in upstream storage, and in 1915, Frank Weber, an engineer, made the first reservoir site investigation. The National Industrial Recovery Act of 1933 allocated funds to the U.S. Indian Service (now the Bureau of Indian Affairs) for the building of the reservoir. On Sept. 21, 1933, construction started, and 10 months later on July 27, 1934, storage began, although it was more than a year later before the dam was completely finished. The reservoir is operated by the Bureau of Indian Affairs to provide summer irrigation water to the Walker River Indian Reservation.

The primary source of water for Weber Reservoir is the Walker River, which heads in the high Sierra Nevada (fig. 1). Flow of the Walker River is somewhat depleted by extensive irrigation upstream from Weber Reservoir. A continuous-record streamflow gaging station is at the head of Campbell Valley on the Walker River near Wabuska (fig. 1), about 16 miles upstream from Weber Dam. This station is the furthest downstream gage on the Walker River. Table 1 lists the annual flows at this site for water years 1924-71 (a water year is the 12-month period, October 1 to September 30). There are streamflow records prior to 1924; however, the years 1924-71 represent the period of manmade modifications of the hydrologic system that exist today. Streamflow at this site for the 48-year period 1924-71 has averaged about 118,000 acre-feet per year, and this value agrees closely with the runoff of 120,000 acre-feet per year of average wetness for the base period 1919-69. Streamflow measurements made about 2 miles upstream from Weber Reservoir suggest that the flows into the reservoir may be on the order of 5 to 15 percent less than the flows measured at the Wabuska gaging station. The loss is due largely to evapotranspiration in excess of local runoff along the river.

Table 1.--Streamflow of the Walker River at the Wabuska gaging station for water years 1924-71

Water year	Streamflow (acre-feet)	Water year	Streamflow (acre-feet)
1924	52,600	1948	31,070
1925	a 20,000	1949	36,520
1926	29,200	1950	30,330
1927	100,000	1951	158,600
1928	46,900	1952	379,000
1929	18,300	1953	121,800
1930	14,500	1954	43,340
1931	9,340	1955	34,620
1932	59,800	1956	277,000
1933	35,900	1957	88,350
1934	21,000	1958	227,300
1935	46,410	1959	70,590
1936	a 110,000	1960	26,260
1937	a 114,000	1961	23,780
1938	a 470,000	1962	37,260
1939	a 80,000	1963	169,200
1940	62,960	1964	51,460
1941	179,900	1965	123,200
1942	a 280,000	1966	107,600
1943	240,100	1967	237,100
1944	a 70,000	1968	90,710
1945	331,900	1969	403,200
1946	170,900	1970	134,900
1947	84,410	1971	93,720

Average annual streamflow (rounded) 118,000 acre-feet

a. Estimated streamflow based on streamflow records on the East Walker River near Mason and the West Walker River near Hudson.

Figure 2 shows the monthly flow distribution at the streamflow gaging station near Wabuska for the water years 1940-71. In general, this flow distribution is also applicable to the monthly inflow to Weber Reservoir.

Weather Bureau records at Schurz and Yerington suggest that the average precipitation on the reservoir is about 5 inches per year. The full reservoir has a surface area of 900 acres (1972 survey); thus, the estimated inflow from precipitation is nearly 380 acre-feet per year (rounded). Using an average evaporation rate of 4 feet per year (Kohler and others, 1959, p. 13), annual evaporation loss is estimated to be about 3,600 acre-feet per year for a full reservoir, which is nine times the average annual precipitation inflow. Everett and Rush (1967) used a precipitation inflow rate of 400 acre-feet per year and an evaporation loss rate of 4,000 acre-feet per year. These figures were based on an average reservoir surface area of 1,000 acres.

Lack of storage prevents the reservoir from being fully effective in the control of downstream flooding, although minor flooding can be reduced. Figure 3 shows the annual variation in contents of Weber Reservoir for the period of continuous record, water years 1956-72.

BATHYMETRY

A continuously recording, electronic fathometer was used to measure the depth of the reservoir on 38 traverses. The reservoir was at maximum operating stage, 4,208.0 feet (considered full with 2 feet of freeboard left on radial spillway gates) during the survey on May 22-23, 1972. Figure 4 shows the stage-area-capacity relations. The new stage-area-capacity figures indicate a water-surface area of 900 acres and a storage capacity of 10,700 acre-feet at stage 4,208 feet, or about 2,400 acre-feet less storage than the original area-capacity figures. This represents an 18 percent loss of computed storage capacity. The maximum reservoir depth found was 29 feet.

SEDIMENTATION

The difference between the two capacities is probably the result of sedimentation and errors in the estimates of storage. The original figures (from the files of the Bureau of Indian Affairs, Stewart, Nev.) show that the minimum elevation on the reservoir floor was at elevation 4,176 feet, which was 3 feet lower than that found by this study. The thickness of fill probably is not uniform but may average somewhere between 2 to 3 feet, or on the order of 2,000 to 2,500 acre-feet. This computes to be an average sedimentation rate in the reservoir for the 38 years of operation of about 60 acre-feet per year. Because of upstream storage reservoirs and natural sediment traps, and because much of the area contributes little sediment, no meaningful erosion rate for the entire watershed could be computed.

Sedimentation data for the reservoir are virtually nonexistent, except for a few samples taken for this study, the results of which are listed in table 2. The data for May 23, 1972 show (1) an increase of sediment concentration with depth of water, and (2) that more sediment was moving into the reservoir (236 mg/l) than was moving out (82 mg/l).

Table 2.--Suspended sediment quantities at selected sites

Location	Date	Concentration (milligrams per liter)	Streamflow at time of sample (cubic feet per second)	Tons of sediment per day during time of sampling
Walker River 2 miles upstream from Weber Reservoir	April 27, 1972	65	80	14
Walker River at outlet of Weber Reservoir	May 23, 1972	80	90	67
Walker Reservoir point sampling:				
Site 1 (bathymetric map)				
12 feet below water surface	do.	47	--	--
20 feet below water surface	do.	57	--	--
Site 2 (bathymetric map)				
3 feet below water surface	do.	29	--	--
13 feet below water surface	do.	54	--	--
26 feet below water surface	do.	164	--	--

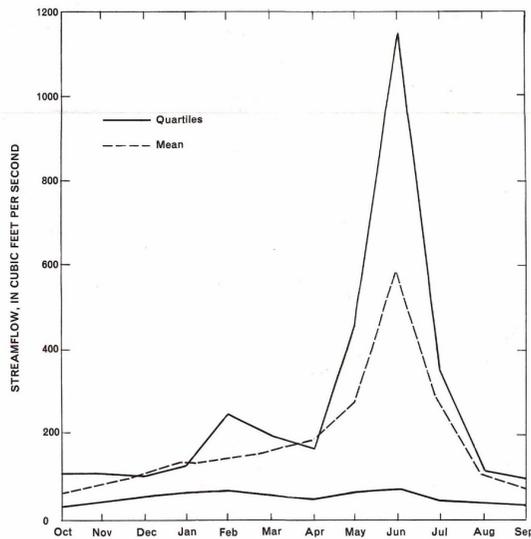


Figure 2.--Mean monthly flow distribution, Walker River near Wabuska, water years 1940-71 (Quartiles show 25 percent of the monthly flows were higher and lower than indicated)

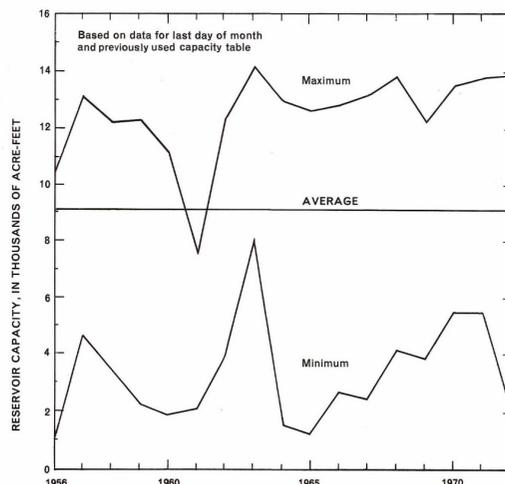


Figure 3.--Annual variation in contents of Weber Reservoir, water years 1956-1972, (furnished by Bureau of Indian Affairs, Stewart, Nev.)

WATER QUALITY

Table 3 lists the results of U.S. Geological Survey chemical analyses of water samples collected during the 1971 water year at the Walker River gaging station near Wabuska. Although this sampling site is several miles above the reservoir, it should be generally representative of the water quality entering the reservoir. Figure 5 shows vertical temperature profiles at two sites in the reservoir taken on May 23, 1972. The water-inflow temperature at the head of the reservoir at that time was 21°C; the outflow temperature, taken as water was being released from the bottom of the reservoir, was 14.5°C.

Table 3.--Summary of water-quality data for river (Data for Walker River gaging station near Wabuska, 16 miles upstream from reservoir, fig. 1)

Constituent	Concentration range ¹ (in milligrams per liter unless otherwise specified)
Water year 1971	
(11 samples)	
Temperature	0.5 - 21.5°C
Calcium	19 - 49
Magnesium	4.6 - 13
Sodium	19 - 80
Bicarbonate	101 - 267
Chloride	6.1 - 32
Nitrate (as N)	0.0 - 1.1
Orthophosphate (as P)	0.09 - 0.40
Dissolved solids	194 - 437
Hardness (as CaCO ₃)	66 - 176
Specific conductance	222 - 677 micromhos

¹ Range in flow at time of sampling, 29 to 1,010 cubic feet per second.

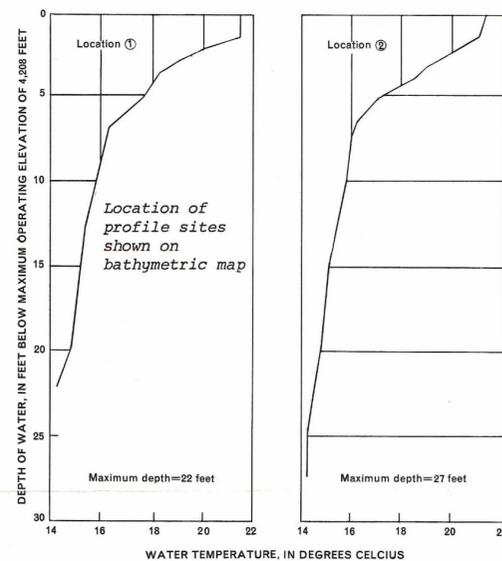


Figure 5.--Vertical water-temperature profiles at two sites in Weber Reservoir, May 23, 1972

OTHER BATHYMETRIC SURVEYS

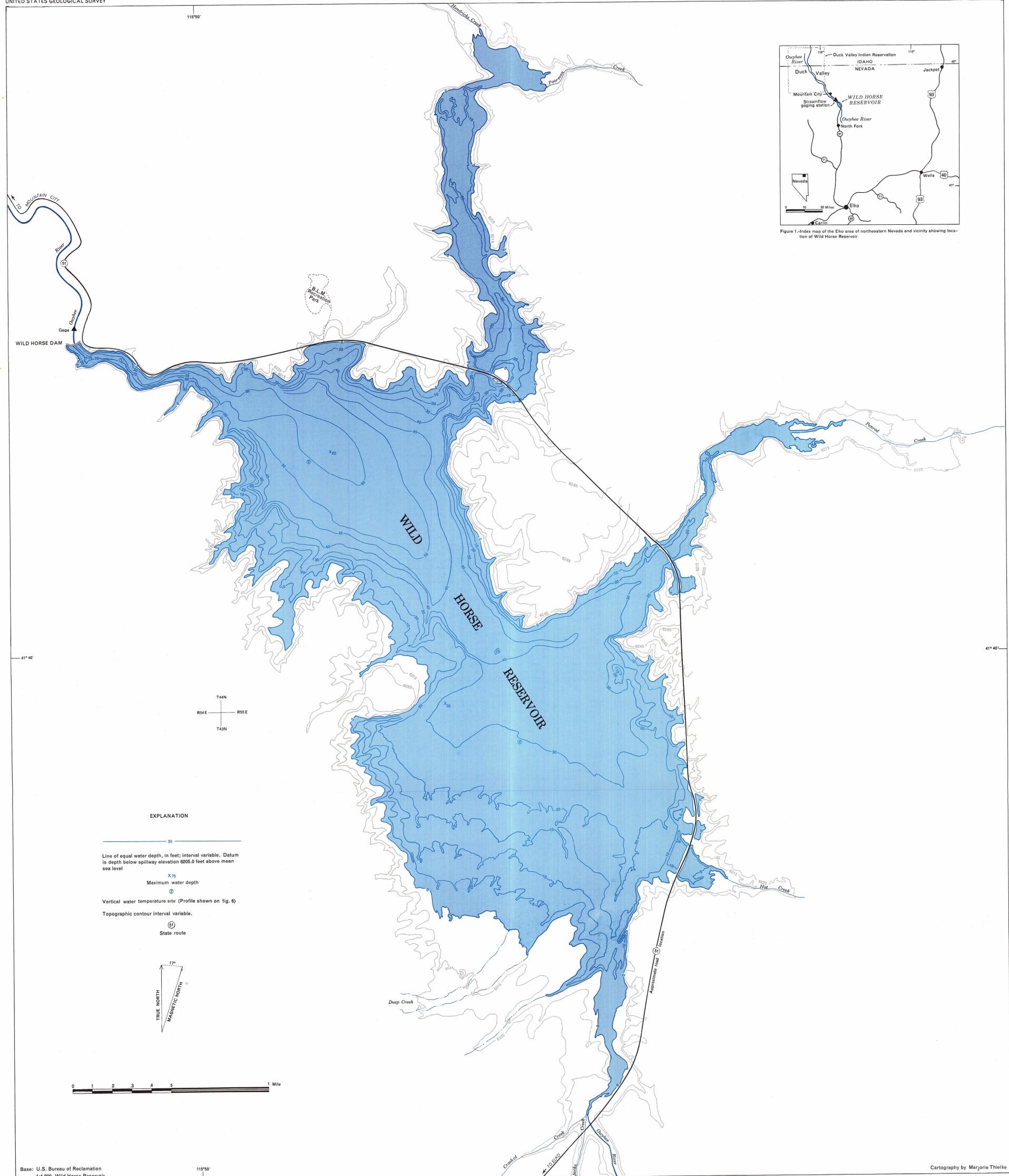
Other bathymetric surveys have been completed or planned, as follows:

Lake or Reservoir	Publication
Pyramid Lake	USGS HA-379 ¹
Walker Lake	USGS HA-415
Lahontan Reservoir	Nev. DWR Info. Ser. 9 ²
Big and Little Washoe Lakes	Nev. DWR Info. Ser. 10
Big and Little Soda Lakes	Nev. DWR Info. Ser. 11
Topaz Lake	Nev. DWR Info. Ser. 12
Rye Patch Reservoir and Upper and Lower Pitt-Taylor Reservoirs	Nev. DWR Info. Ser. 13
Marlette and Spooner Lakes	Nev. DWR Info. Ser. 14
Weber Reservoir	Nev. DWR Info. Ser. 15
Wild Horse Reservoir	Nev. DWR Info. Ser. 16
Lake Tahoe	Nev. DWR Info. Ser. 17

¹ U.S. Geological Survey Hydrologic Atlas.
² Nevada Division of Water Resources Information Series Report.

REFERENCES CITED

Everett, D. E., and Rush, F. E., 1967, A brief appraisal of the water resources of the Walker Lake area, Mineral, Lyon, and Churchill Counties, Nevada: Nevada Dept. Conserv. and Nat. Resources, Water Resources - Recon. Ser. Rept. 40.
Kohler, M. A., and others, 1959, Evaporation maps of the United States: U.S. Weather Bureau Tech. Paper no. 37, p. 13.



BATHYMETRIC RECONNAISSANCE OF WILD HORSE RESERVOIR, ELKO COUNTY, NEVADA

By
T. L. Katzer and Lynn Harmsen
1973

Base: U.S. Bureau of Reclamation
1:4,800, Wild Horse Reservoir

Cartography by Marjorie Thielke

BATHYMETRIC RECONNAISSANCE OF WILD HORSE RESERVOIR,
ELKO COUNTY, NEVADA

By T. L. Katzner and Lynn Harmsen

INTRODUCTION

Wild Horse Reservoir is in Elko County, on the northeastern flank of the Independence Mountains, about 62 miles north of Elko and 12 miles southeast of Mountain City, on State Route 51 (fig. 1). The drainage area of the Owyhee River, the main tributary to Wild Horse Reservoir, at the gaging station immediately below the reservoir is about 209 square miles. The Owyhee River is one of Nevada's few north-flowing streams and is a tributary to the Snake River (not shown in fig. 1). Wild Horse Reservoir, when full, is at an elevation of 6,205 feet above mean sea level and the surrounding mountains commonly have peaks above an elevation of 7,000 feet.

The need for supplemental irrigation water for Duck Valley on the Western Shoshone Indian Reservation led to the first reservoir site investigation by Halbert T. Johnson in 1916. The National Industrial Recovery Act of 1933 allocated funds to the U.S. Indian Service (now the Bureau of Indian Affairs) for the building of a dam. Construction started in June 1936, ended in June 1937, and storage began March 18, 1938. According to the Bureau of Indian Affairs, at spillway elevation 6,189.2 feet, the reservoir had a water-surface area of 1,860 acres and a capacity of 33,500 acre-feet.

According to the U.S. Bureau of Reclamation, the aggregate used in the construction of the first dam proved to be of poor quality. The Bureau of Indian Affairs, therefore, decided to build a new dam to remove a possible safety hazard and allow the reservoir to be enlarged to about double its former capacity. Construction by the U.S. Bureau of Reclamation on the second dam started in September 1967 and was completed in June 1969. This present dam is a double-curvature, thin-arch, concrete structure. The crest is 435 feet long, the spillway is 75 feet wide at elevation 6,205.0 feet. The 1972 survey determined that at spillway elevation 6,205.0 feet, the surface area is 2,830 acres, and the capacity is 73,500 acre-feet.

The reservoir is operated to provide irrigation water for about 21,000 acres of Indian land (W. H. Hoy, Bureau of Indian Affairs, Stewart, Nev., oral commun., 1973). Figure 2 shows the reservoir's annual variation in contents for the water years 1938-71 (a water year is from October 1 to September 30). The original capacity table was used prior to 1969. Table 1 presents the annual releases and spills from the reservoir, as recorded by a downstream gaging station on the Owyhee River (fig. 1). Figure 3 shows the mean monthly flow distribution for the Owyhee River below the reservoir, before and after dam construction.

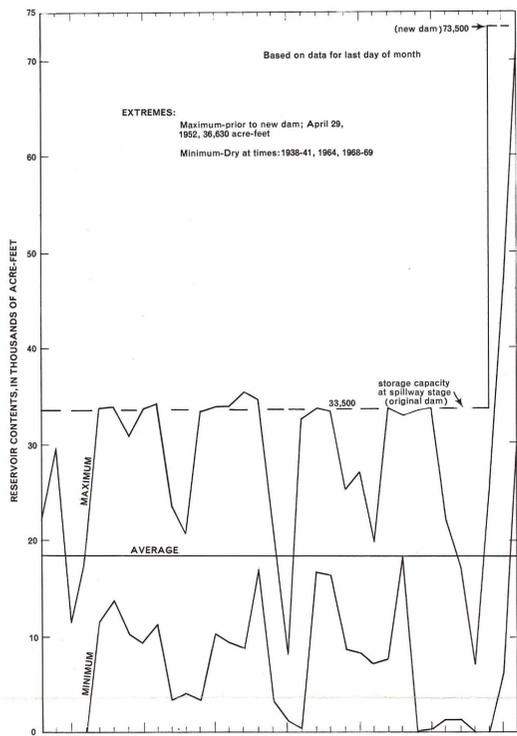


Figure 2.—Annual variation in contents of Wild Horse Reservoir, Water years 1938-71.

Table 1.—Outflow from Wild Horse Reservoir in the Owyhee River, 1939-72 (Measured at the gaging station near Gold Creek)

Water year	Streamflow in acre-feet (rounded)	Water year	Streamflow in acre-feet (rounded)
1939	39,700	1957	39,300
1940	15,100	1958	37,200
1941	a 9,000	1959	19,600
1942	33,200	1960	12,600
1943	b 63,700	1961	38,200
1944	27,200	1962	32,900
1945	37,200	1963	14,500
1946	45,800	1964	52,900
1947	22,900	1965	23,000
1948	17,900	1966	25,900
1949	26,800	1967	13,500
1950	33,800	1968	11,000
1951	46,000	1969	40,000
1952	62,100	1970	13,800
1953	30,300	1971	37,600
1954	21,800	1972	54,600
1955	9,800		
1956	15,300	Average	30,100

a. Minimum.
b. Maximum.

Weather Bureau records at North Fork (fig. 1) suggest that the average annual precipitation on the reservoir is about 10 inches. According to Kohler and others (1959, pl. 2), the average annual evaporation from the reservoir is about 3.5 feet (net evaporation, about 2.7 feet). Based on these figures, the construction of the new dam has increased the average inflow from direct precipitation on the reservoir from about 1,500 acre-feet to about 2,300 acre-feet per year at spillway stage, and the water lost by evaporation has also increased from about 6,500 acre-feet to about 10,000 acre-feet at spillway stage.

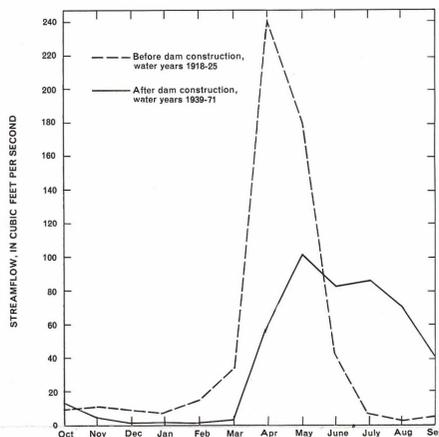


Figure 3.—Mean monthly flow distribution, Owyhee River near Gold Creek, water years 1918-25 and 1939-71.

BATHYMETRY

A continuous recording, electronic fathometer was used to measure the depth of the reservoir on 48 traverses. The Nevada Fish and Game Department provided a boat which was operated by Enforcement Officer Robert Poling. The bathymetric survey was made on May 25-26, 1972, when the reservoir was spilling at stage 6,205.6 feet.

Figure 4 is a graph of the stage-area-volume relations. The results of this survey indicate that at spillway elevation 6,205 feet, the water-surface area is 2,830 acres and the storage capacity is 73,500 acre-feet. A comparison of these results and data furnished by the Bureau of Indian Affairs, Stewart, Nev., shows that at spillway stage the water-surface area computed in this survey to be 7 percent less and the capacity to be 3 percent greater than those previously used. The differences are within the accuracy limits of this reconnaissance. The stage-area-volume tables of the Bureau of Indian Affairs show that the original reservoir, at spillway stage 6,189.2 feet, had a water-surface area of 1,860 acres and a capacity of 33,500 acre-feet. The 1972 survey shows that at the original reservoir spillway stage the surface area was 1,870 acres and the capacity was 36,700 acre-feet. The differences in capacity are again considered to be within the accuracy limits of the two surveys. Bathymetric data were used to develop the stage-area-volume relations below elevation 6,185 feet, and existing topographic maps were used above elevation 6,185 feet.

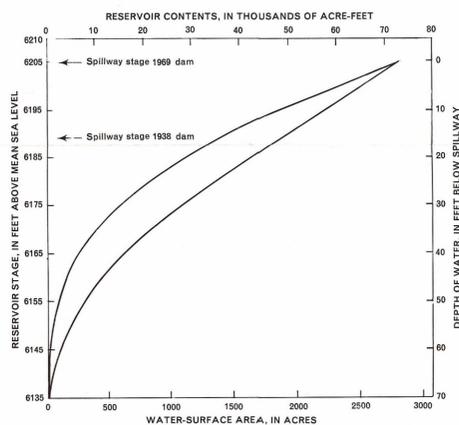


Figure 4.—Stage-area-capacity relations of Wild Horse Reservoir (1972 Survey)

SEDIMENTATION AND WATER QUALITY

Sediment movement and deposition data are lacking for Wild Horse Reservoir. The results of this survey, however, indicate that there has been no identifiable loss of storage due to sedimentation, and therefore sediment transport into Wild Horse Reservoir is assumed to have been minor. The lack of identifiable sedimentation in the reservoir is mainly due to the relatively stabilized soils, small size of the drainage area above the reservoir, and small runoff.

Table 2 presents the results of the Nevada State Health Department water sampling and analyses of the Owyhee River during 1971. The sampling site was about 10 miles downstream from the reservoir.

Table 2.—Water-quality data for sampling site downstream from Wild Horse Reservoir

[Data for Owyhee River immediately upstream from Mountain City (fig. 1), furnished by the Nevada State Health Department]

	Concentration in milligrams per liter unless otherwise specified	
	May 24	August 30
Approximate flow, cubic feet per second	250	80
Temperature	12.5°C (54.5°F)	18.5°C (65.3°F)
pH (units)	7.9	8.3
Dissolved oxygen	8.9	8.3
Chloride (Cl)	1	4
Orthophosphate (PO ₄)	.16	.40
Nitrate (NO ₃)	.2	0
Dissolved solids	121	137
Alkalinity	74	94
Bicarbonate	90	115
Carbonate	0	0

Figure 5 shows vertical temperature profiles at four sites in the reservoir.

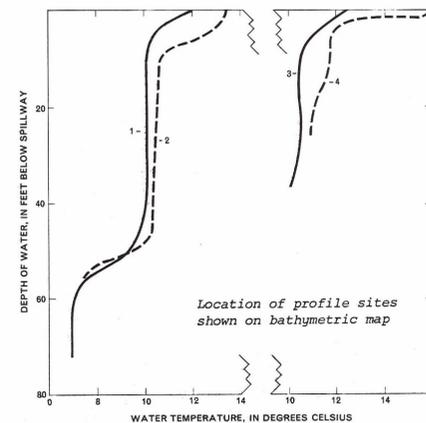


Figure 5.—Vertical water-temperature profiles for Wild Horse Reservoir, May 26, 1972

OTHER BATHYMETRIC SURVEYS

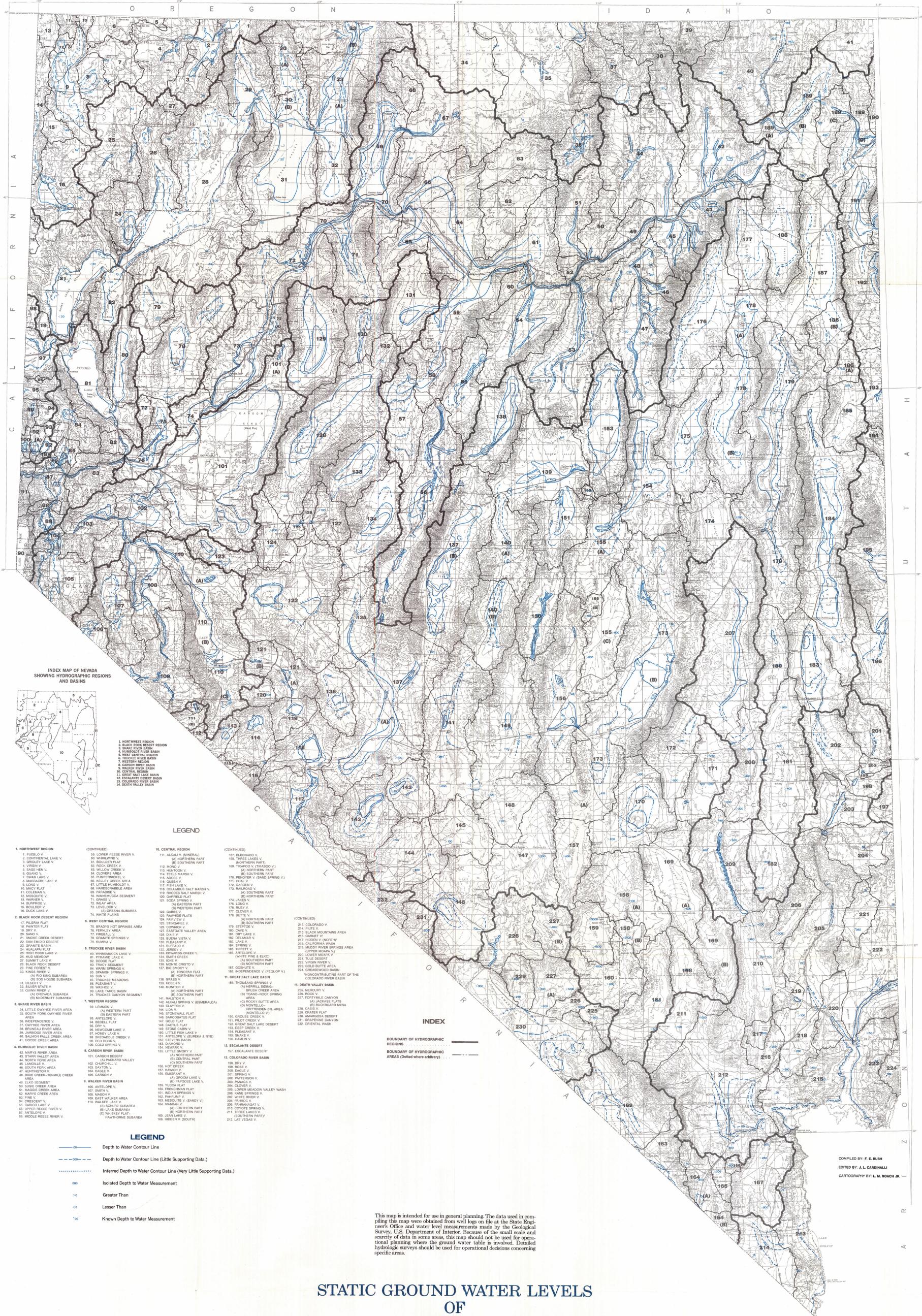
Reconnaissance bathymetric surveys of this series are listed below:

Lake or Reservoir	Publication
Pyramid Lake	USGS HA-379 ¹
Walker Lake	USGS HA-415 ²
Lahontan Reservoir	Nev. DWR Info. Ser. 9 ²
Big and Little Washoe Lakes	Nev. DWR Info. Ser. 10
Big and Little Soda Lakes	Nev. DWR Info. Ser. 11
Topaz Lake	Nev. DWR Info. Ser. 12
Rye Patch Reservoir and Upper and Lower Pitt-Taylor Reservoirs	Nev. DWR Info. Ser. 13
Marlette and Spooner Lakes	Nev. DWR Info. Ser. 14
Weber Reservoir	Nev. DWR Info. Ser. 15
Wild Horse Reservoir	Nev. DWR Info. Ser. 16
Lake Tahoe	Nev. DWR Info. Ser. 17

¹ U.S. Geological Survey, Hydrologic Atlas.
² Nevada Division of Water Resources, Information Series Report.

REFERENCE CITED

Kohler, M. A., and others, 1959, Evaporation maps of the United States: U.S. Weather Bureau Tech. Paper No. 37, p. 13.



INDEX MAP OF NEVADA
SHOWING HYDROGRAPHIC REGIONS
AND BASINS

LEGEND

- 1. NORTHWEST REGION
 - 1. PUEBLO V.
 - 2. CONTINENTAL LAKE V.
 - 3. BRIDLEY LAKE V.
 - 4. WINDY V.
 - 5. SAGE HEN V.
 - 6. QUAKY V.
 - 7. SWAN LAKE V.
 - 8. MASSACRE LAKE V.
 - 9. LONG V.
 - 10. MACK FLAT
 - 11. COLUMBIAN V.
 - 12. MOSQUITO V.
 - 13. HORNED V.
 - 14. BURRIS V.
 - 15. BOULDER V.
 - 16. DUCK LAKE V.
- 2. BLACK ROCK DESERT REGION
 - 17. PILGRIM FLAT
 - 18. DRY V.
 - 19. SAND V.
 - 20. SMOKE CREEK DESERT
 - 21. GRANITE DESERT
 - 22. SAN EMILIO DESERT
 - 23. GRANITE BASIN
 - 24. HUALAPAI FLAT
 - 25. HIGH ROCK LAKE V.
 - 26. MID MEADOW
 - 27. SUMMIT LAKE V.
 - 28. BLACK ROCK DESERT
 - 29. PINE FOREST V.
 - 30. KNIFE RIVER V.
 - 31. DESEY V.
 - 32. SILVER STATE V.
 - 33. JUNIPER RIVER V.
 - 34. OREGONA SUBAREA
 - 35. MOUNTAIN SUBAREA
- 3. SNAKE RIVER BASIN
 - 36. LITTLE OYHIEE RIVER AREA
 - 37. SOUTH FORK OYHIEE RIVER AREA
 - 38. INDEPENDENCE V.
 - 39. OYHIEE RIVER AREA
 - 40. BRUNEAU RIVER AREA
 - 41. CARBIDE RIVER AREA
 - 42. JOCKEY CREEK AREA
- 4. HUMBOLDT RIVER BASIN
 - 43. MARYS RIVER AREA
 - 44. STARR VALLEY AREA
 - 45. NORTH FORK AREA
 - 46. LANDLIDE V.
 - 47. SOUTH FORK AREA
 - 48. HUNTINGTON V.
 - 49. ROSE CREEK-TENNILE CREEK AREA
 - 50. ELKO SEGMENT
 - 51. SAGE CREEK AREA
 - 52. MARYS CREEK AREA
 - 53. PINE V.
 - 54. CARDO LAKE V.
 - 55. UPPER REESE RIVER V.
 - 57. ANTELOPE V.
 - 58. LOWER REESE RIVER V.
- 5. WEST CENTRAL REGION
 - 59. LOWER REESE RIVER V.
 - 60. WINDY RIVER V.
 - 61. BOULDER FLAT
 - 62. ROCK CREEK V.
 - 63. WILLOW CREEK V.
 - 64. CLOVER AREA
 - 65. PUMPERNICKEL V.
 - 66. KELLEY CREEK AREA
 - 67. LITTLE HUMBOLDT V.
 - 68. HINDENBERG AREA
 - 69. WANNAMUCKA SEGMENT
 - 70. HANSHIRE V.
 - 71. MALAY AREA
 - 72. DOWNEY V.
 - 73. WESTERN PLAINS
 - 74. WEST PLAINS
 - 75. BRADY'S HOT SPRINGS AREA
 - 76. PENNEY AREA
 - 77. FRENCH V.
 - 78. GRANITE SPRINGS V.
 - 79. KUMVA V.
- 6. TRUCKEE RIVER BASIN
 - 80. WINDMUCKA LAKE V.
 - 81. HUMBOLDT LAKE V.
 - 82. DOOSE FLAT
 - 83. BRADY'S HOT SPRINGS V.
 - 84. WARM SPRINGS V.
 - 85. TRUCKEE MEADOWS
 - 86. PLEASANT V.
 - 87. TRUCKEE CANYON SEGMENT
 - 88. LAKE TAHOE BASIN
 - 89. WAGNER V.
- 7. WESTERN REGION
 - 90. LEMMON V.
 - 91. WESTERN PART
 - 92. ANTELOPE V.
 - 93. BEBBS FLAT
 - 94. DRY V.
 - 95. MICHIGAN LAKE V.
 - 96. HONEY LAKE V.
 - 97. HONEY LAKE V.
 - 98. ANTELOPE V. (EQUIV. & NYE)
 - 99. RED ROCK V.
 - 100. COLD SPRING V.
 - 101. NEWBY V.
 - 102. LITTLE SMOKY V.
 - 103. SOUTHERN PART
 - 104. HOT CREEK
 - 105. KAMOH V.
 - 106. EMIGRANT V.
 - 107. GROCK LAKE V.
 - 108. YUCCA FLAT
 - 109. FRENCHMAN FLAT
 - 110. INDIAN SPRINGS V.
 - 111. PIRRIUM V.
 - 112. MESQUITE V. (BANKY V.)
 - 113. WARDEN V.
 - 114. SOUTHERN PART
 - 115. JEAN LAKE V.
 - 116. HOBBS V. (SOUTH)
- 8. CENTRAL REGION
 - 117. ALKALI V. (MINERAL)
 - 118. SOUTHERN PART
 - 119. MOBY WASH V.
 - 120. HUNTING V.
 - 121. YELLS WASH V.
 - 122. ADOLF V.
 - 123. QUEEN V.
 - 124. FISH LAKE V.
 - 125. COLUMBIAN SALT MARSH V.
 - 126. RICHMOND FLAT
 - 127. SOUTH SPRING V.
 - 128. GARDNER V.
 - 129. WESTERN PART
 - 130. GARDNER V.
 - 131. SOUTHERN PART
 - 132. SOUTHERN PART
 - 133. LONG V.
 - 134. CLOVER V.
 - 135. BUTTE V.
 - 136. GARDNER V.
 - 137. STEPTOE V.
 - 138. LONG V.
 - 139. DRY LAKE V.
 - 140. GANNETT V.
 - 141. CALIFORNIA WASH
 - 142. CALIFORNIA WASH
 - 143. MADDY RIVER SPRINGS AREA
 - 144. UPPER MADDY V.
 - 145. LOWER MADDY V.
 - 146. TULE DESERT
 - 147. VIRGIN RIVER V.
 - 148. GOLD BUTTE AREA
 - 149. GREENWOOD BASIN
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 - 151. HERRILL BASIN
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 - 154. ROCKY BUTTE AREA
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 - 159. BARCORTAUS FLAT
 - 160. GOLD FLAT
 - 161. STONE CANYON V.
 - 162. LITTLE FISH LAKE V.
 - 163. ANTELOPE V. (EQUIV. & NYE)
 - 164. STEVENS BASIN
 - 165. DANMAY V.
 - 166. NEWBY V.
 - 167. LITTLE SMOKY V.
 - 168. SOUTHERN PART
 - 169. HOT CREEK
 - 170. KAMOH V.
 - 171. EMIGRANT V.
 - 172. GROCK LAKE V.
 - 173. YUCCA FLAT
 - 174. FRENCHMAN FLAT
 - 175. INDIAN SPRINGS V.
 - 176. PIRRIUM V.
 - 177. MESQUITE V. (BANKY V.)
 - 178. WARDEN V.
 - 179. SOUTHERN PART
 - 180. JEAN LAKE V.
 - 181. HOBBS V. (SOUTH)
- 9. GREAT SALT LAKE BASIN
 - 182. THOUSAND SPRINGS V.
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 - 204. GROCK LAKE V.
 - 205. YUCCA FLAT
 - 206. FRENCHMAN FLAT
 - 207. INDIAN SPRINGS V.
 - 208. PIRRIUM V.
 - 209. MESQUITE V. (BANKY V.)
 - 210. WARDEN V.
 - 211. SOUTHERN PART
 - 212. JEAN LAKE V.
 - 213. HOBBS V. (SOUTH)
- 10. COLORADO RIVER BASIN
 - 214. ELDOBRADO V.
 - 215. THREE LAKE V.
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 - 217. SOUTHERN PART
 - 218. SOUTHERN PART
 - 219. PENNEY V. (SAND SPRING V.)
 - 220. COAL V.
 - 221. GARDNER V.
 - 222. WESTERN PART
 - 223. SOUTHERN PART
 - 224. LONG V.
 - 225. CLOVER V.
 - 226. BUTTE V.
 - 227. GARDNER V.
 - 228. STEPTOE V.
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 - 241. HERRILL BASIN
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 - 244. ROCKY BUTTE AREA
 - 245. MONTICELLO V.
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 - 247. LIDA V.
 - 248. STONEMALL FLAT
 - 249. BARCORTAUS FLAT
 - 250. GOLD FLAT
 - 251. STONE CANYON V.
 - 252. LITTLE FISH LAKE V.
 - 253. ANTELOPE V. (EQUIV. & NYE)
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 - 268. WARDEN V.
 - 269. SOUTHERN PART
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 - 272. MCHERRY V.
 - 273. ROCKY MOUNTAIN BASIN
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 - 299. NEWBY V.
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 - 303. KAMOH V.
 - 304. EMIGRANT V.
 - 305. GROCK LAKE V.
 - 306. YUCCA FLAT
 - 307. FRENCHMAN FLAT
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 - 309. PIRRIUM V.
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 - 311. WARDEN V.
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 - 313. JEAN LAKE V.
 - 314. HOBBS V. (SOUTH)

INDEX

- BOUNDARY OF HYDROGRAPHIC REGIONS
- BOUNDARY OF HYDROGRAPHIC AREAS (Dotted where arbitrary)

LEGEND

- Depth to Water Contour Line
- - - - - Depth to Water Contour Line (Little Supporting Data.)
- Inferred Depth to Water Contour Line (Very Little Supporting Data.)
- 000 Isolated Depth to Water Measurement
- > Greater Than
- < Lesser Than
- * Known Depth to Water Measurement

This map is intended for use in general planning. The data used in compiling this map were obtained from well logs on file at the State Engineer's Office and water level measurements made by the Geological Survey, U.S. Department of Interior. Because of the small scale and scarcity of data in some areas, this map should not be used for operational planning where the ground water table is involved. Detailed hydrologic surveys should be used for operational decisions concerning specific areas.

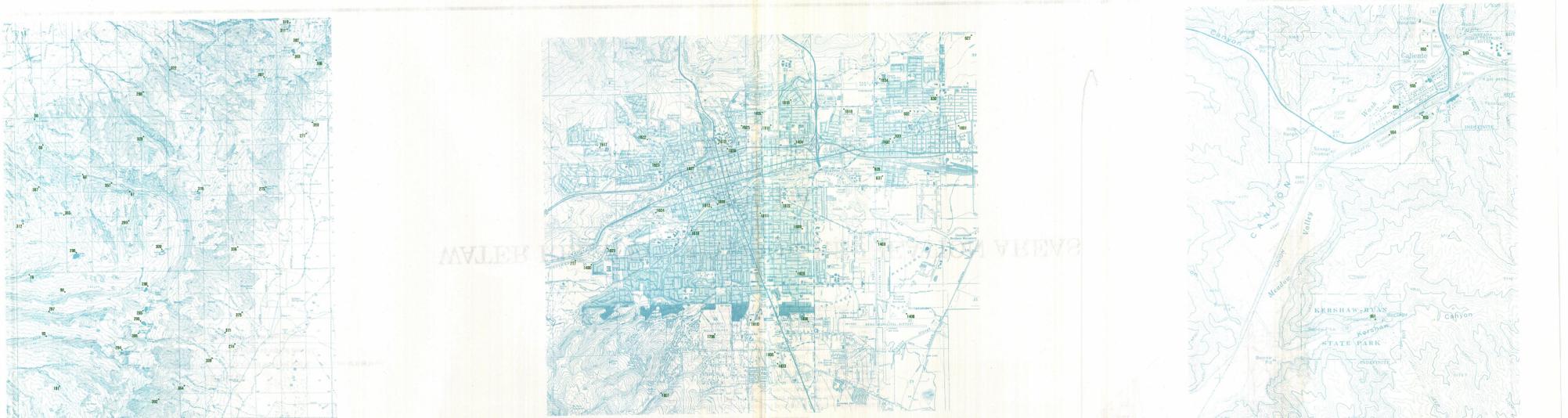
STATIC GROUND WATER LEVELS
OF
NEVADA

PREPARED AS PART OF THE NEVADA STATE WATER PLAN
DIVISION OF WATER RESOURCES
STATE ENGINEERS OFFICE

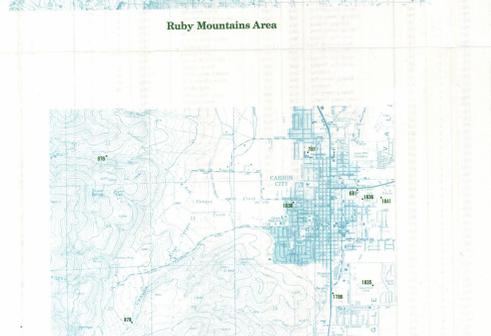
1974

SCALE 1:750,000
1:500,000 (approximate)

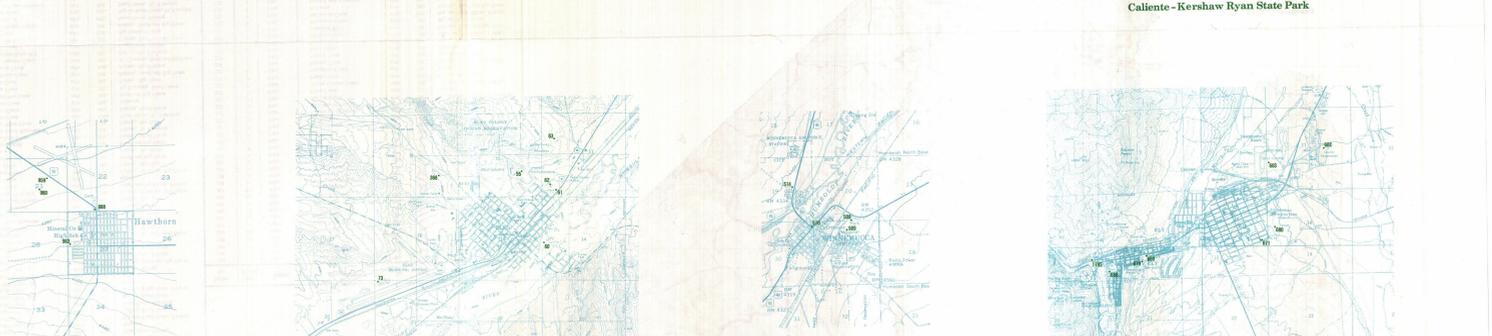
COMPILED BY: F. E. RUSH
EDITED BY: A. L. CARDONALI
CARTOGRAPHY BY: L. M. BOACH, JR.



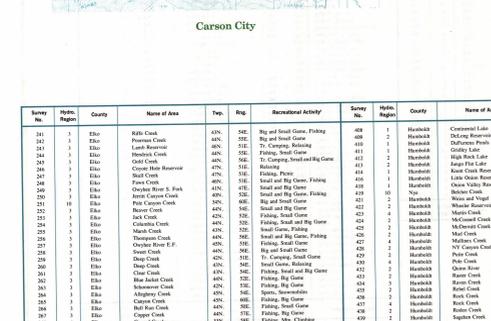
Reno-Sparks



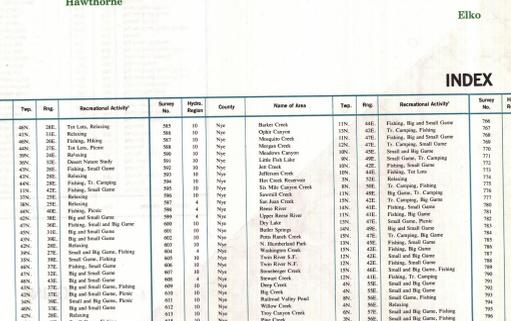
Ruby Mountains Area



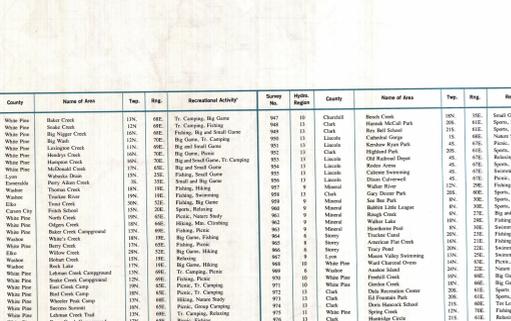
Caliente - Kershaw Ryan State Park



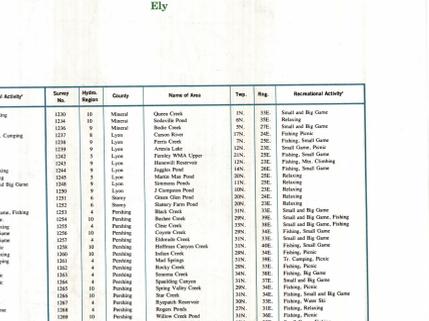
Carson City



Hawthorne



Elko

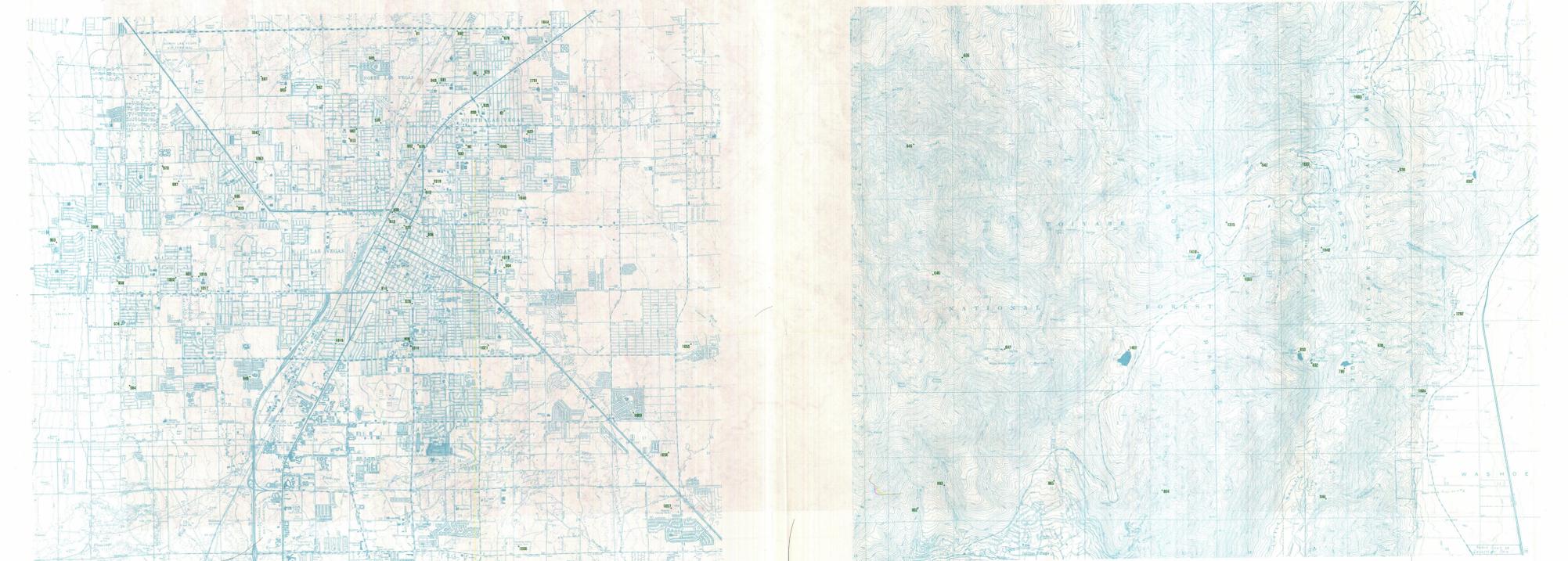


Winnemucca

Ely

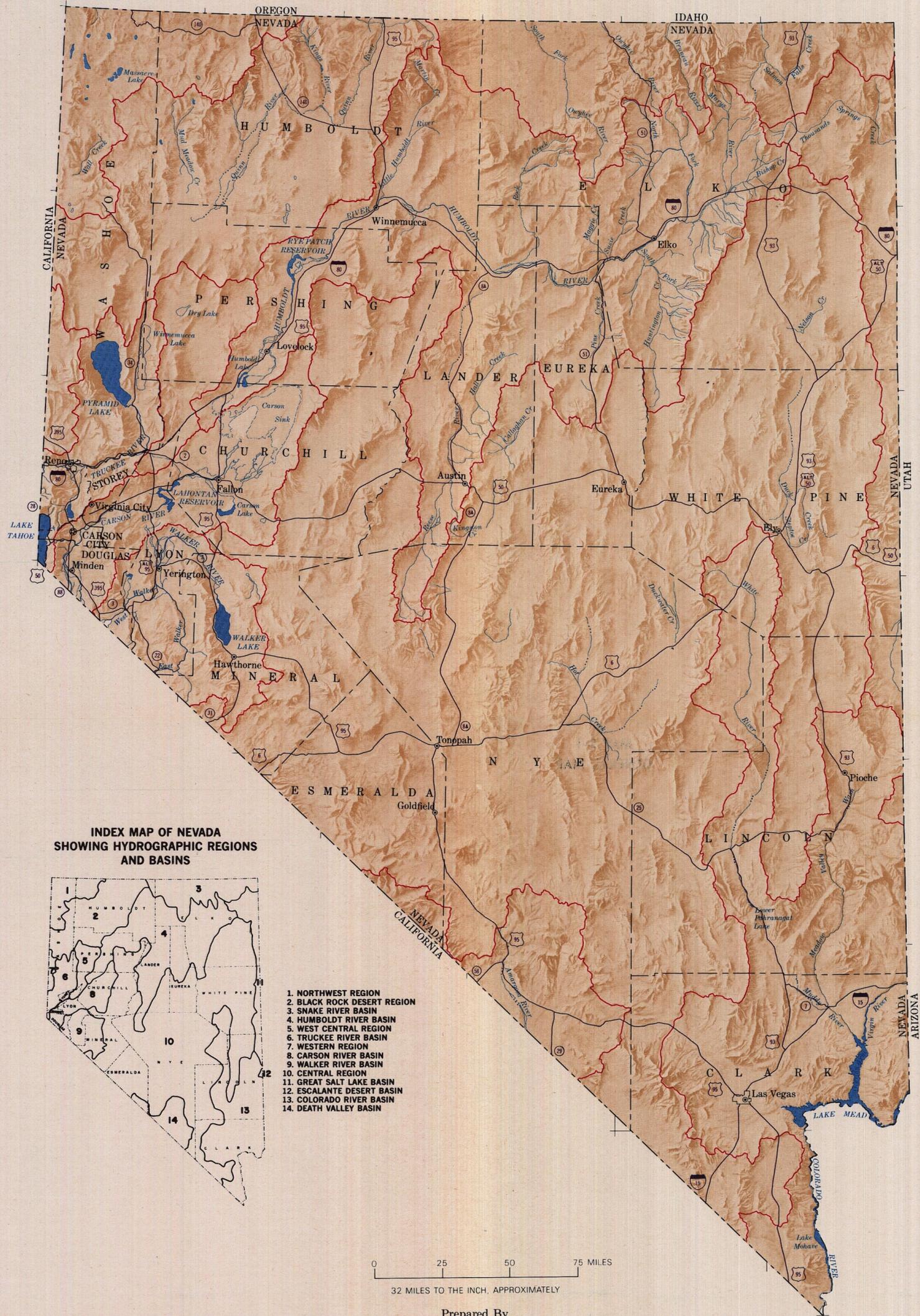
INDEX

Table with 20 columns: Survey No., Hydro Region, County, Name of Area, Type, Recreational Activity, Survey No., Hydro Region, County, Name of Area, Type, Recreational Activity, Survey No., Hydro Region, County, Name of Area, Type, Recreational Activity, Survey No., Hydro Region, County, Name of Area, Type, Recreational Activity. The table lists numerous recreation areas across Nevada, including names like 'Bainville Park', 'Cottonwood Canyon', 'Hawthorne', 'Elko', 'Winnemucca', 'Carson City', 'Reno-Sparks', 'Caliente', and 'Ely'.



Las Vegas-North Las Vegas

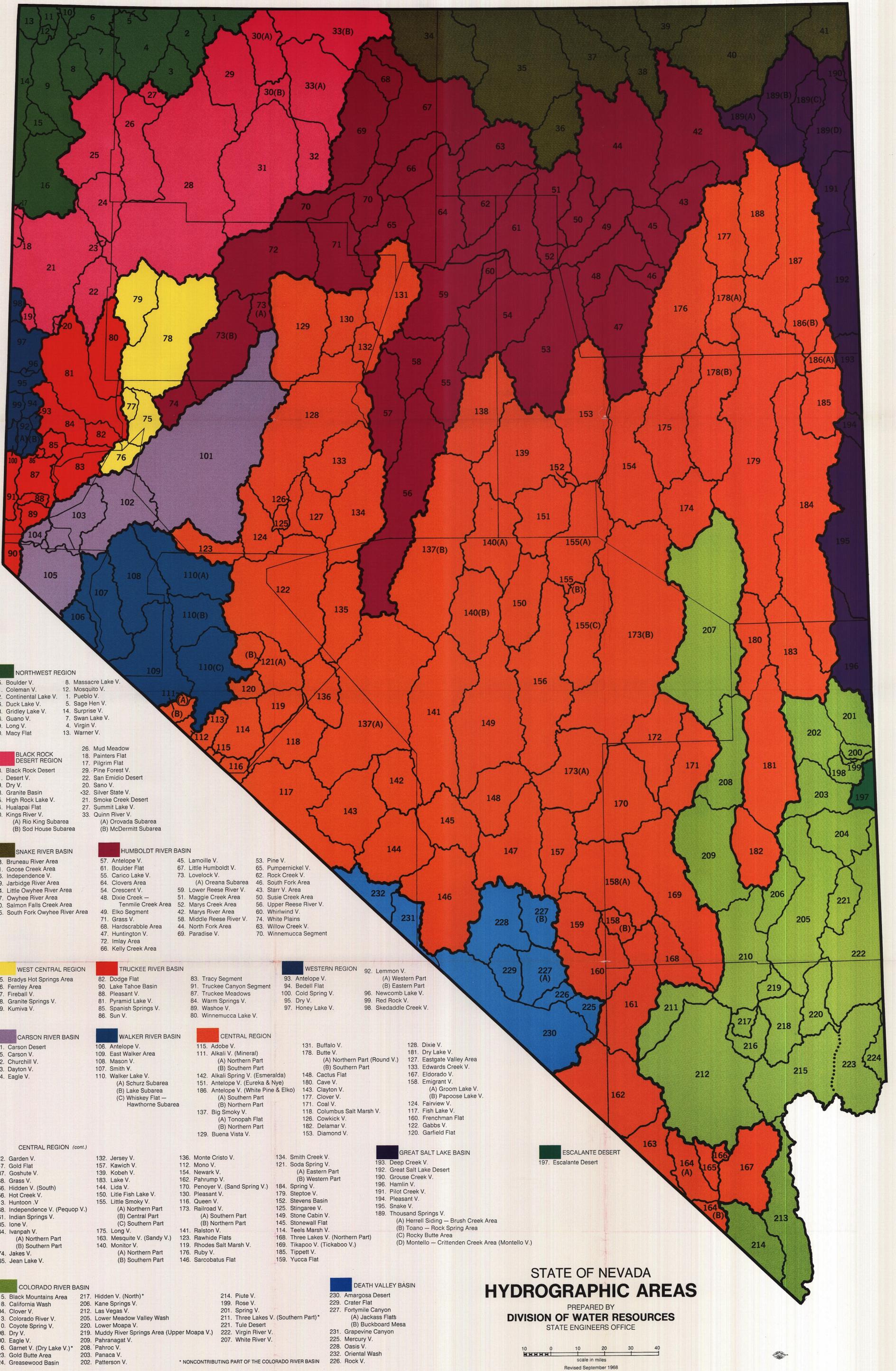
Mt. Rose-Slide Mtn. Areas

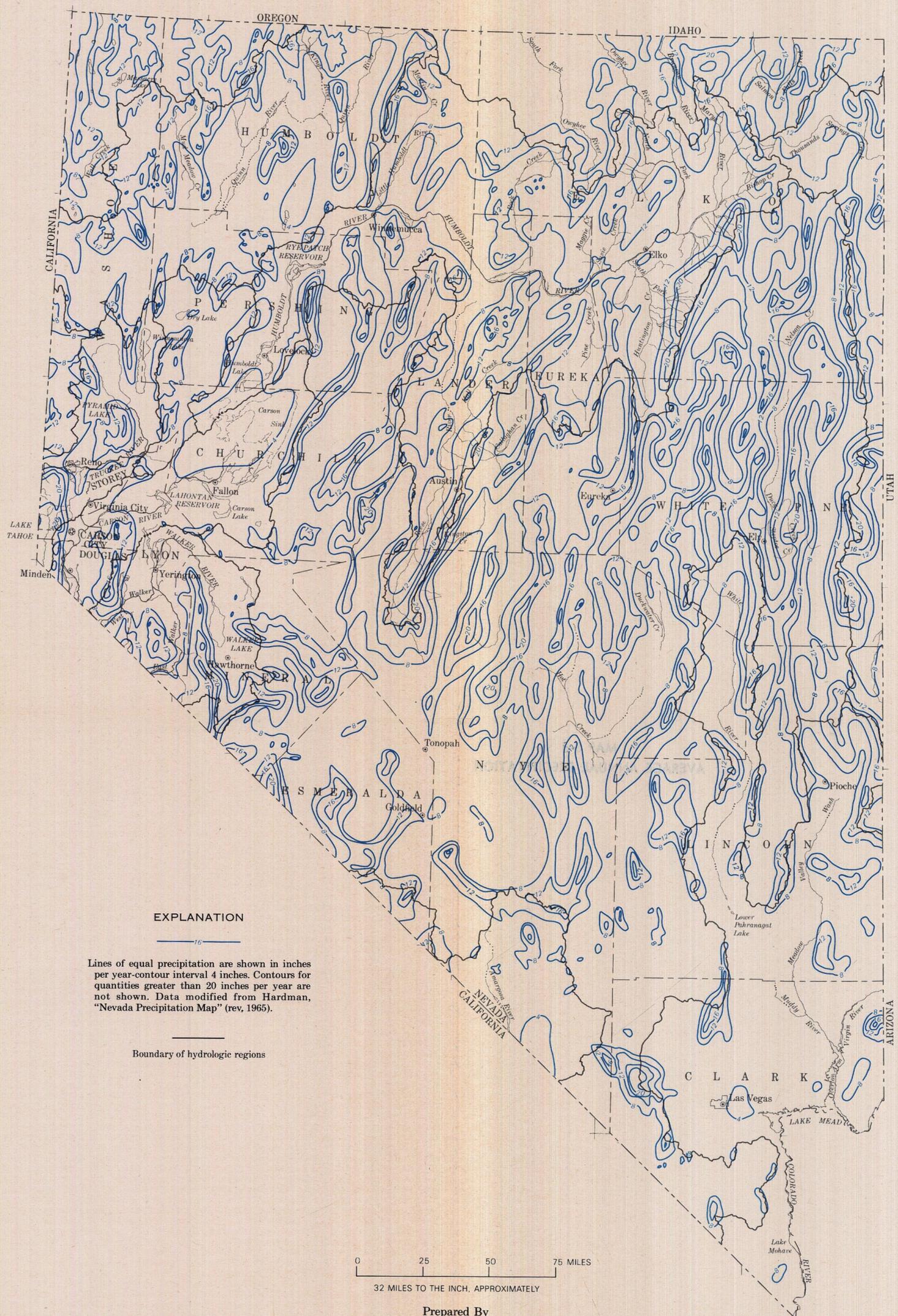


Prepared By
STATE OF NEVADA
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1971
GENERAL MAP OF NEVADA

Edited by B. R. Scott

Cartography by L. M. Roach, Jr.
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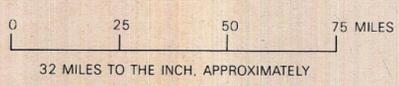




EXPLANATION

— 16 —
 Lines of equal precipitation are shown in inches per year—contour interval 4 inches. Contours for quantities greater than 20 inches per year are not shown. Data modified from Hardman, "Nevada Precipitation Map" (rev. 1965).

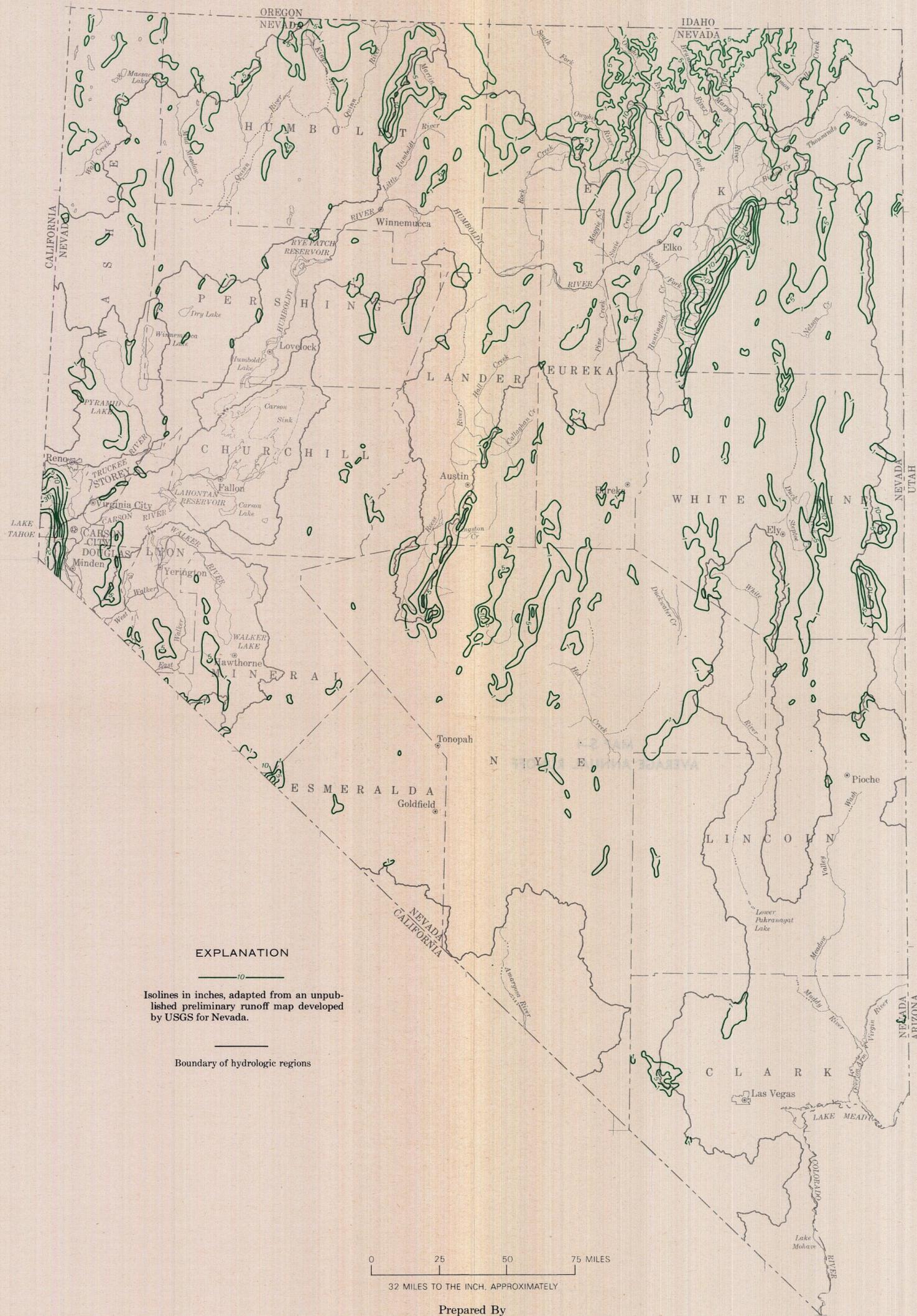
— — — — —
 Boundary of hydrologic regions



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AVERAGE ANNUAL PRECIPITATION

Edited by B. R. Scott

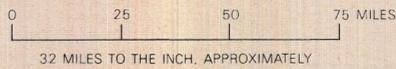
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EXPLANATION

— Isohyes in inches, adapted from an unpublished preliminary runoff map developed by USGS for Nevada.

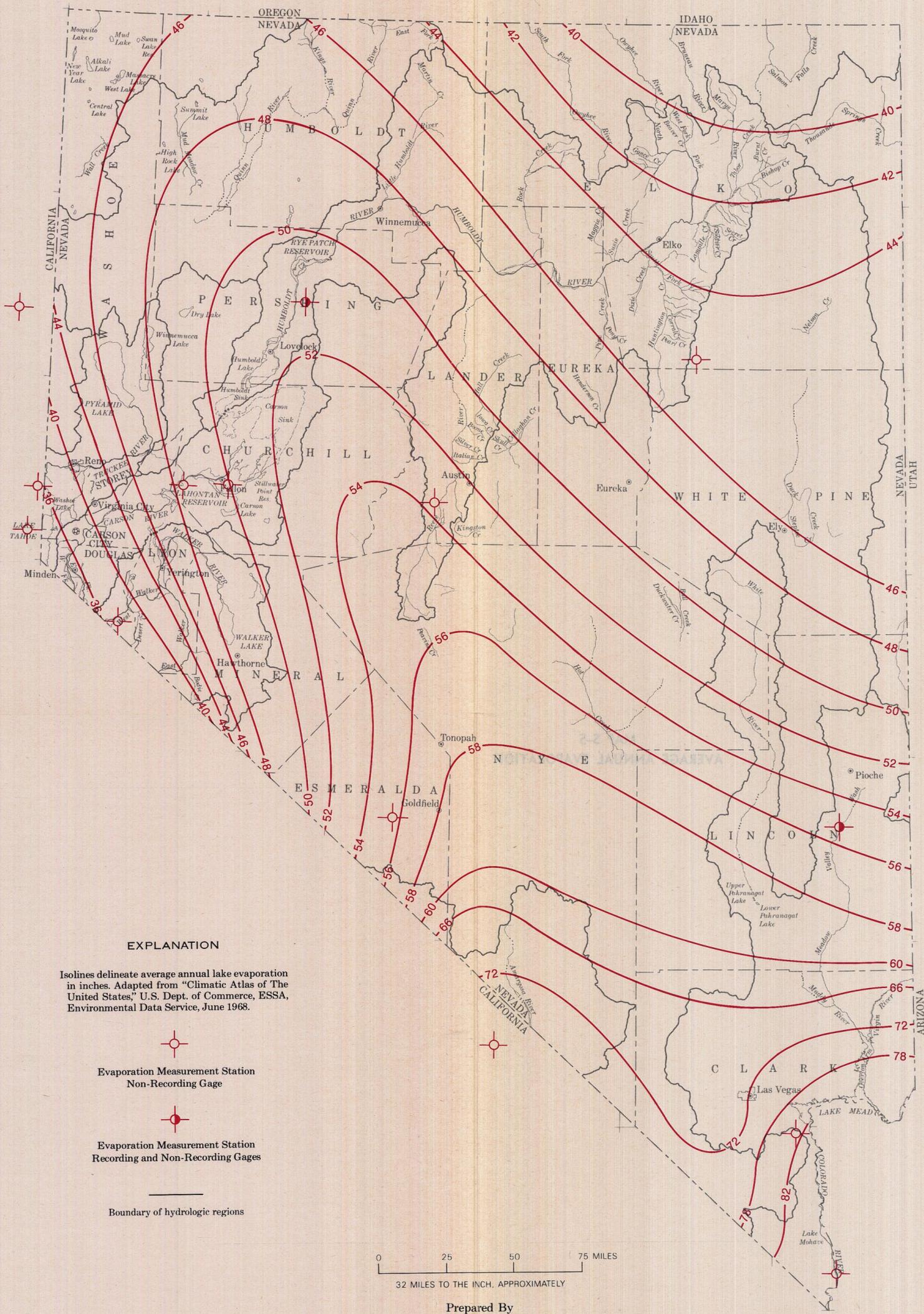
--- Boundary of hydrologic regions



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AVERAGE ANNUAL RUNOFF

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EXPLANATION

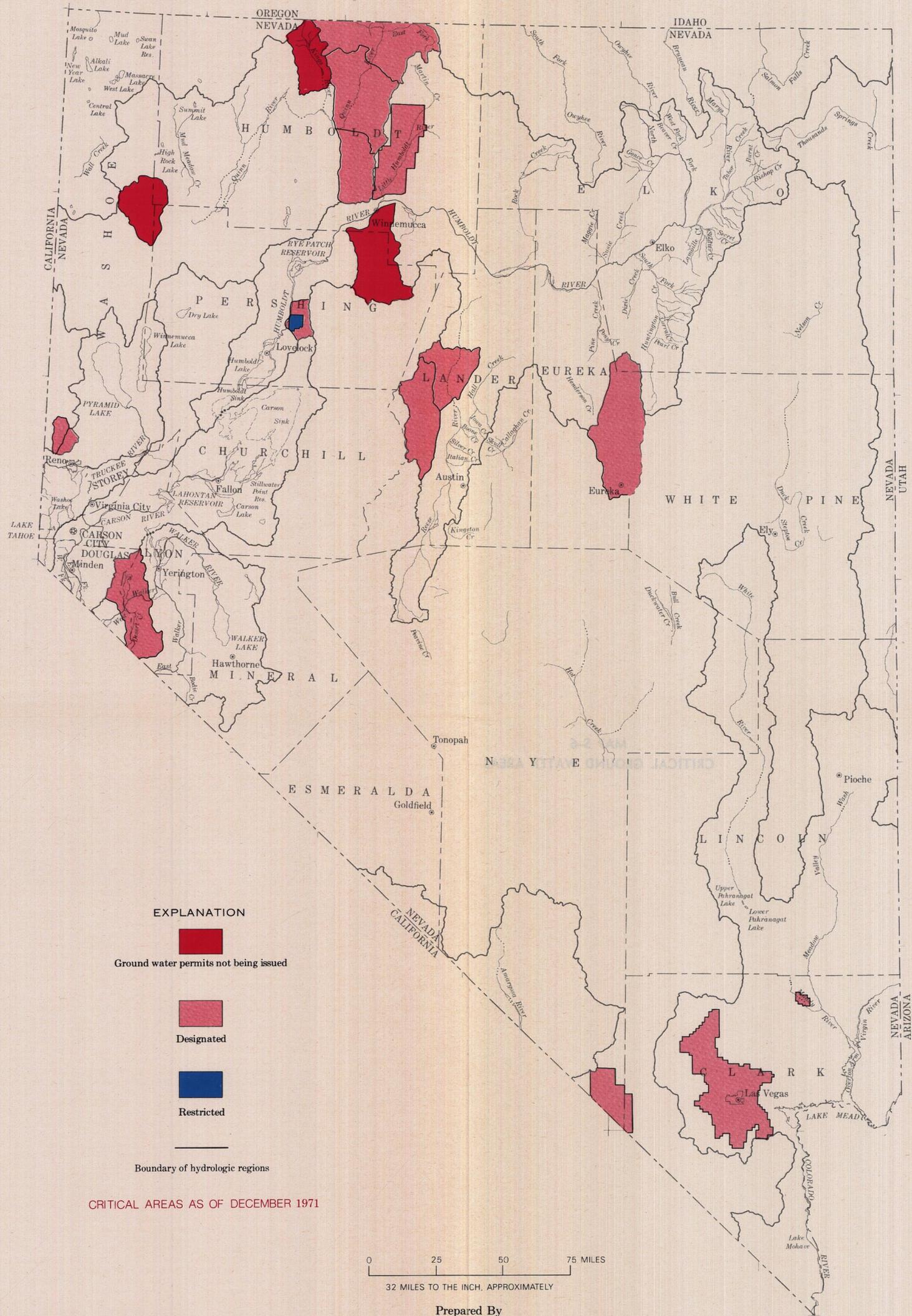
Isolines delineate average annual lake evaporation in inches. Adapted from "Climatic Atlas of The United States," U.S. Dept. of Commerce, ESSA, Environmental Data Service, June 1968.

-  Evaporation Measurement Station
Non-Recording Gage
-  Evaporation Measurement Station
Recording and Non-Recording Gages
-  Boundary of hydrologic regions

0 25 50 75 MILES
32 MILES TO THE INCH, APPROXIMATELY

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AVERAGE ANNUAL EVAPORATION

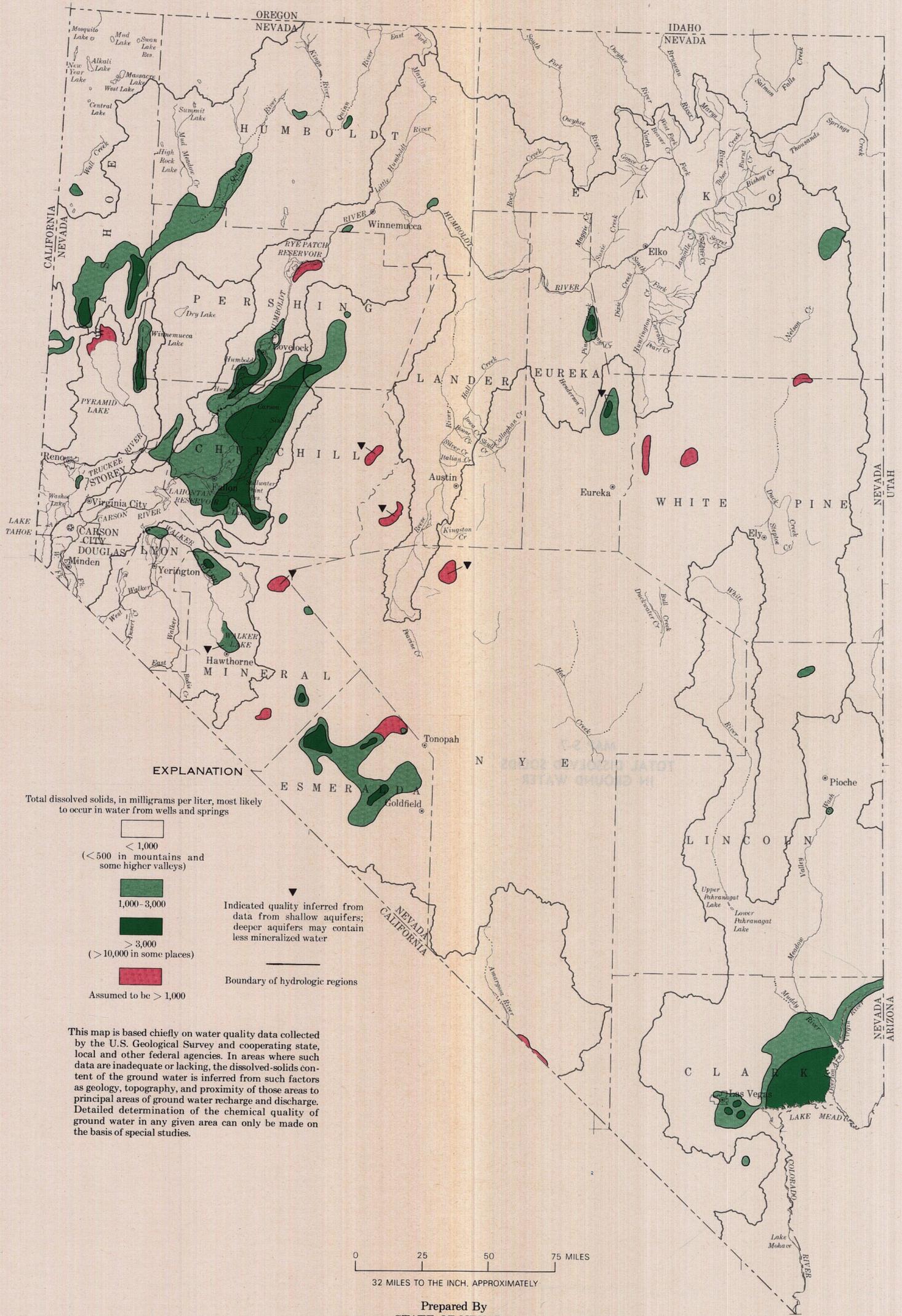
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CRITICAL GROUND WATER AREAS

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BASED ON STATE-FEDERAL COMPREHENSIVE FRAMEWORK STUDIES



EXPLANATION

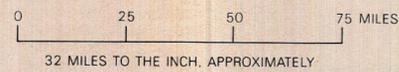
Total dissolved solids, in milligrams per liter, most likely to occur in water from wells and springs

- < 1,000
(< 500 in mountains and some higher valleys)
- 1,000 - 3,000
- > 3,000
(> 10,000 in some places)
- Assumed to be > 1,000

Indicated quality inferred from data from shallow aquifers; deeper aquifers may contain less mineralized water

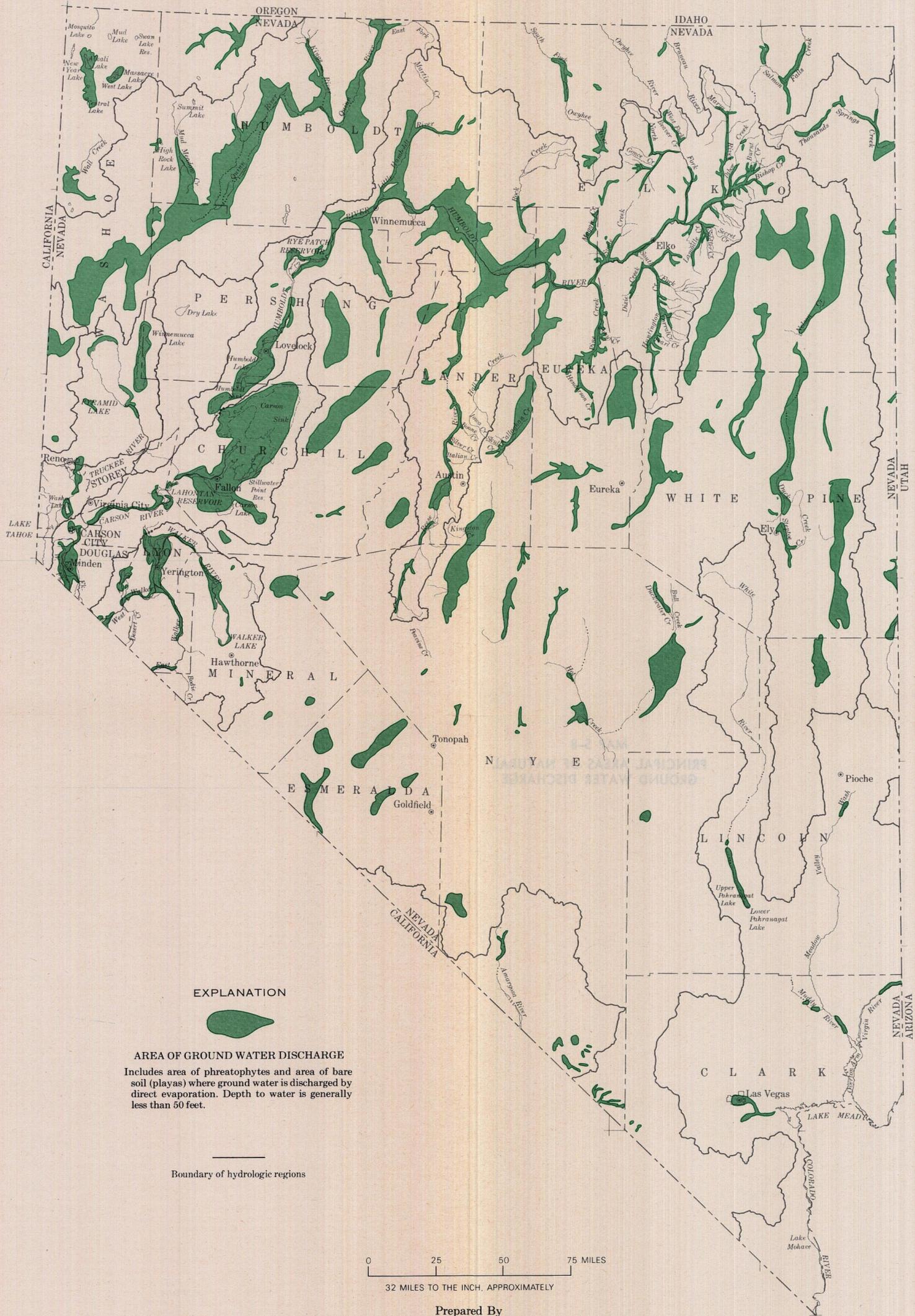
Boundary of hydrologic regions

This map is based chiefly on water quality data collected by the U.S. Geological Survey and cooperating state, local and other federal agencies. In areas where such data are inadequate or lacking, the dissolved-solids content of the ground water is inferred from such factors as geology, topography, and proximity of those areas to principal areas of ground water recharge and discharge. Detailed determination of the chemical quality of ground water in any given area can only be made on the basis of special studies.



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TOTAL DISSOLVED SOLIDS IN GROUND WATER

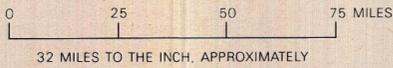
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EXPLANATION

AREA OF GROUND WATER DISCHARGE
 Includes area of phreatophytes and area of bare soil (playas) where ground water is discharged by direct evaporation. Depth to water is generally less than 50 feet.

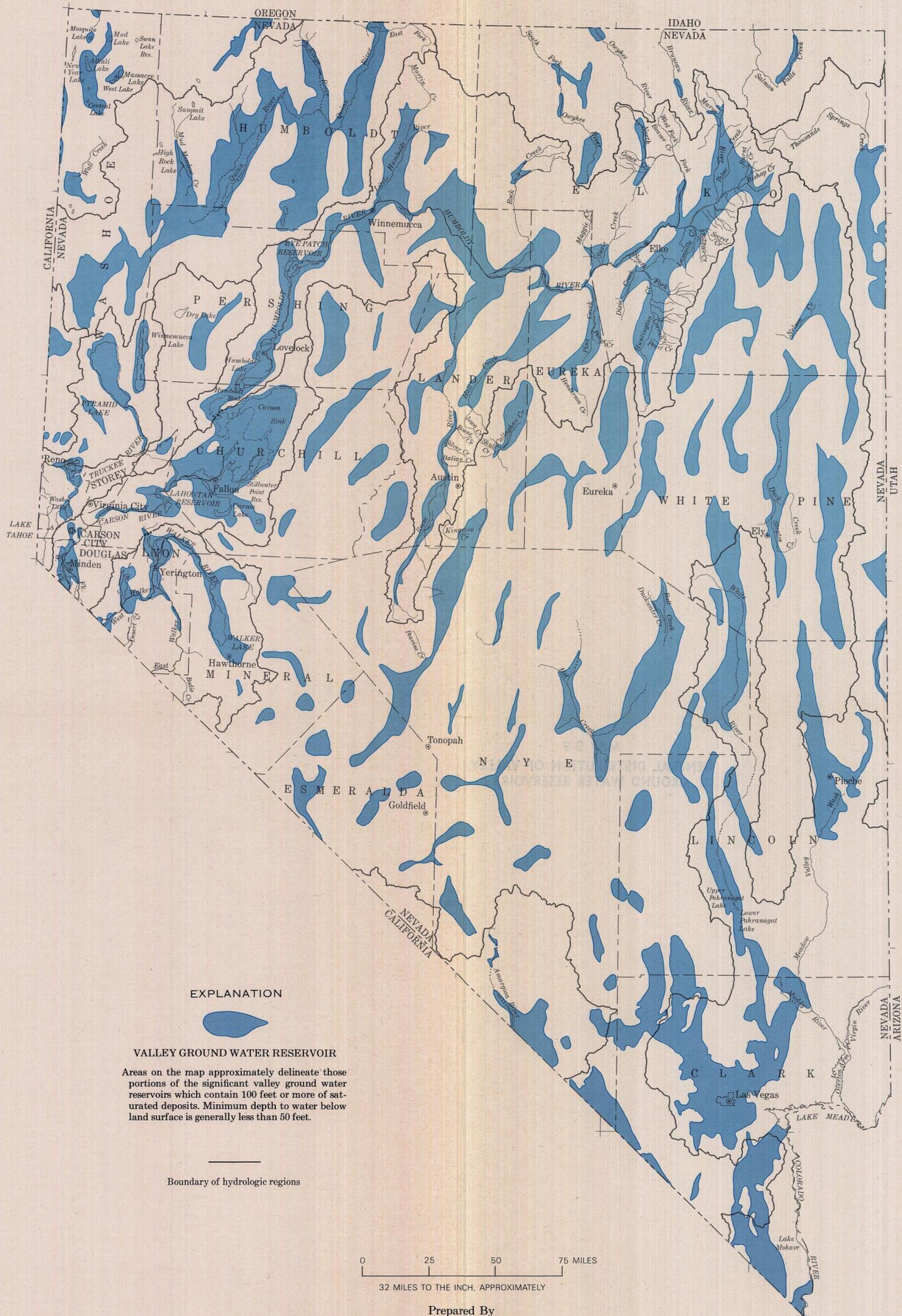
Boundary of hydrologic regions



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**PRINCIPAL AREAS OF NATURAL
 GROUND WATER DISCHARGE**

Edited by B. R. Scott

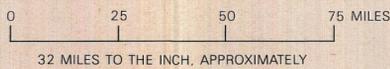
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EXPLANATION

VALLEY GROUND WATER RESERVOIR
 Areas on the map approximately delineate those portions of the significant valley ground water reservoirs which contain 100 feet or more of saturated deposits. Minimum depth to water below land surface is generally less than 50 feet.

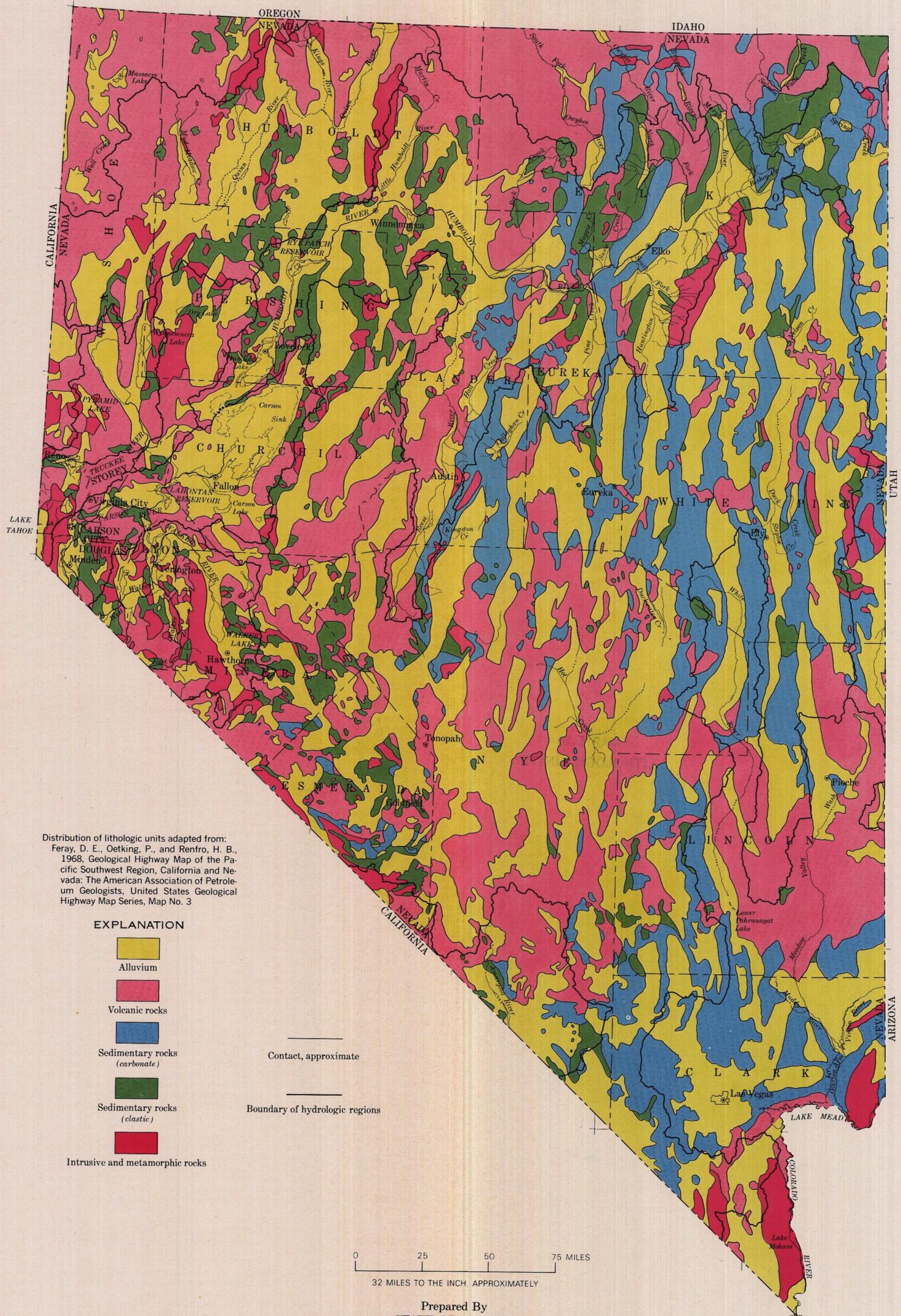
— Boundary of hydrologic regions



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**GENERAL DISTRIBUTION OF VALLEY
 GROUND WATER RESERVOIRS**

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Distribution of lithologic units adapted from:
 Feray, D. E., Oetking, P., and Renfro, H. B.,
 1968. Geological Highway Map of the Pa-
 cific Southwest Region, California and Ne-
 vada: The American Association of Petrole-
 um Geologists, United States Geological
 Highway Map Series, Map No. 3

EXPLANATION

- Alluvium
 - Volcanic rocks
 - Sedimentary rocks
(carbonate)
 - Sedimentary rocks
(clastic)
 - Intrusive and metamorphic rocks
- Contact, approximate
 Boundary of hydrologic regions

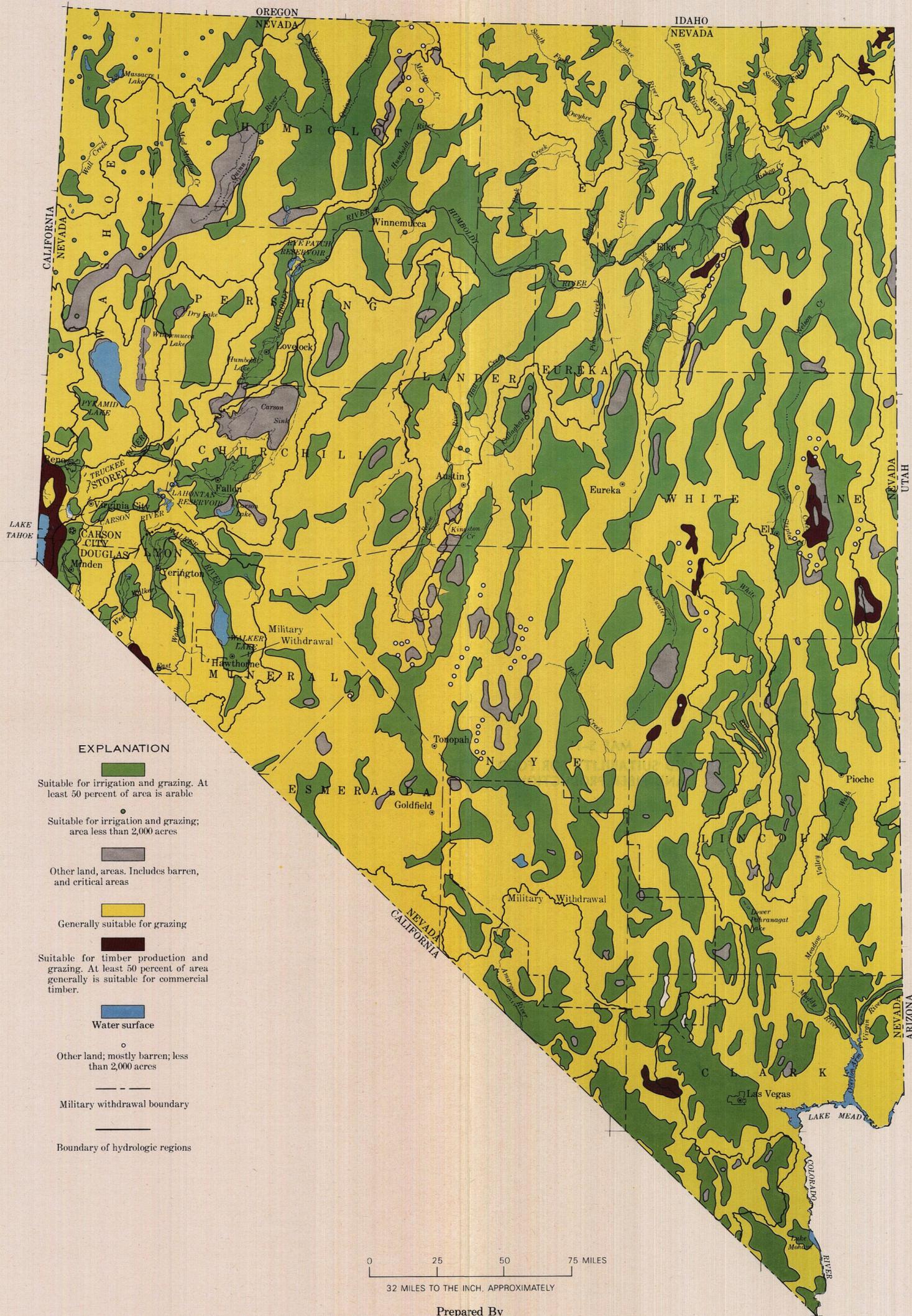
0 25 50 75 MILES
 32 MILES TO THE INCH, APPROXIMATELY

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**GENERAL DISTRIBUTION OF THE PRINCIPAL
 LITHOLOGIC UNITS IN NEVADA**

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EXPLANATION

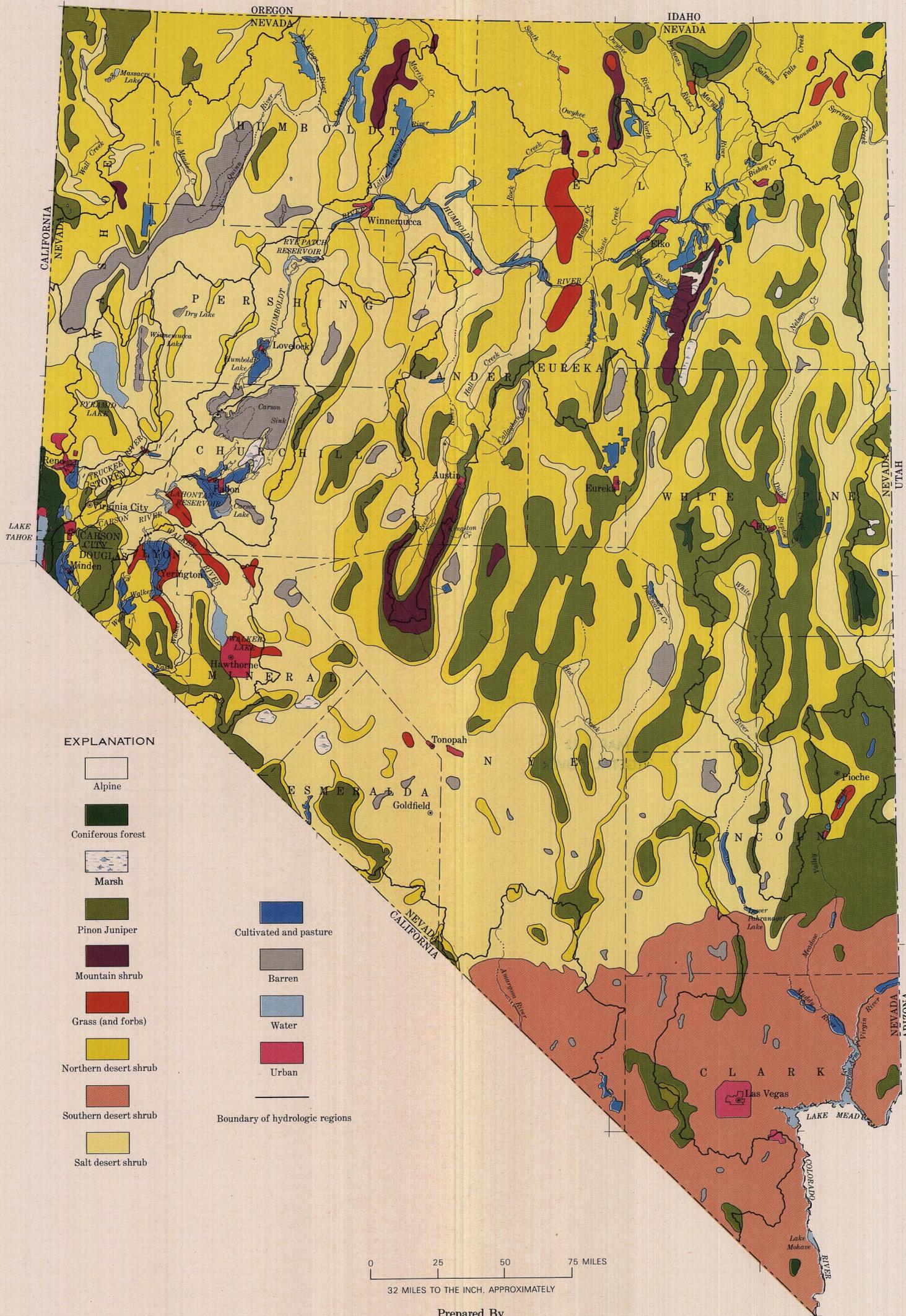
- Suitable for irrigation and grazing. At least 50 percent of area is arable
- Suitable for irrigation and grazing; area less than 2,000 acres
- Other land areas. Includes barren, and critical areas
- Generally suitable for grazing
- Suitable for timber production and grazing. At least 50 percent of area generally is suitable for commercial timber.
- Water surface
- Other land; mostly barren; less than 2,000 acres
- Military withdrawal boundary
- Boundary of hydrologic regions

0 25 50 75 MILES
32 MILES TO THE INCH, APPROXIMATELY

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**LAND SUITABILITY FOR
FOOD AND FIBER PRODUCTION**

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EXPLANATION

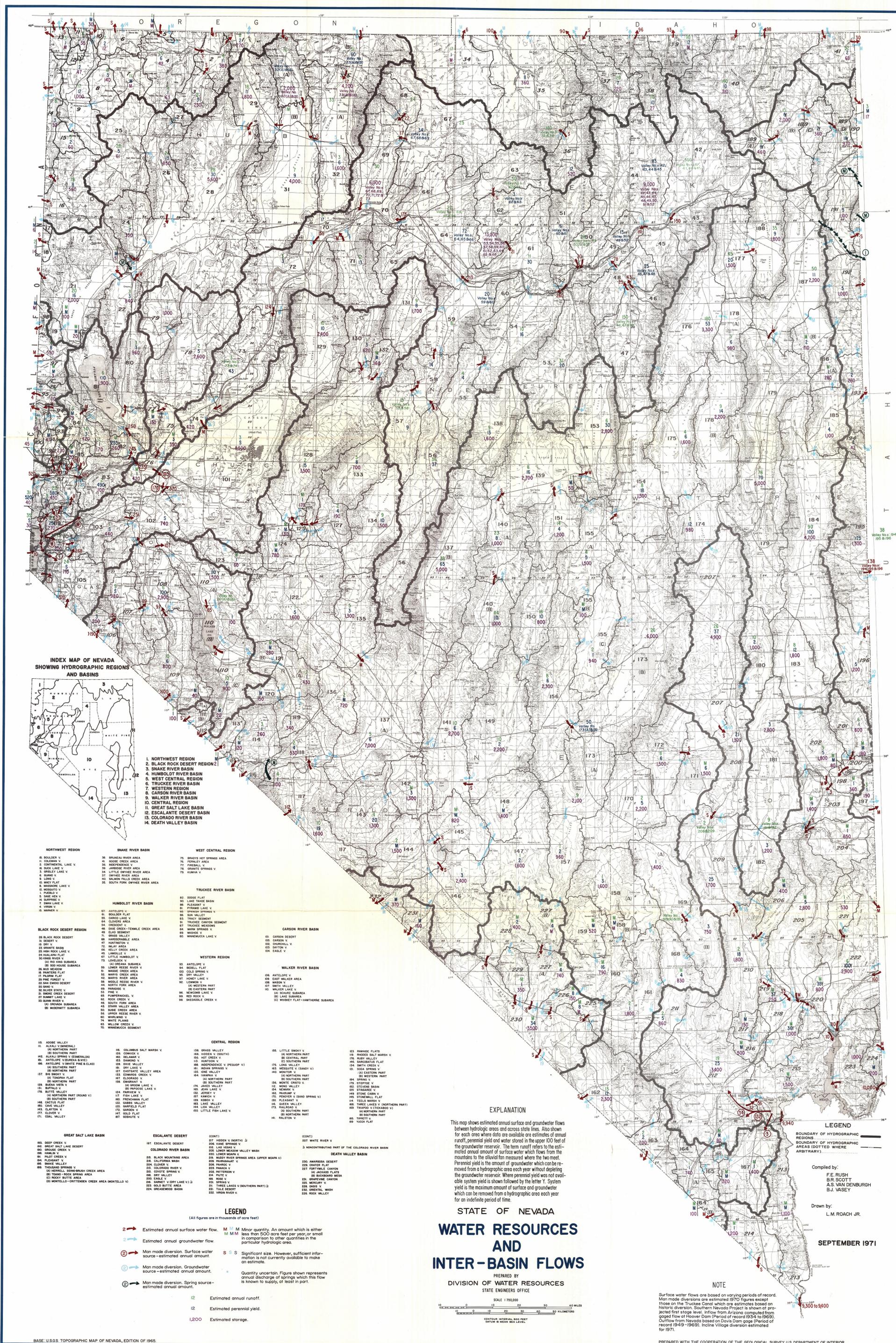
- | | | | |
|--|-----------------------|--|--------------------------------|
| | Alpine | | Cultivated and pasture |
| | Coniferous forest | | Barren |
| | Marsh | | Water |
| | Pinon Juniper | | Urban |
| | Mountain shrub | | Boundary of hydrologic regions |
| | Grass (and forbs) | | |
| | Northern desert shrub | | |
| | Southern desert shrub | | |
| | Salt desert shrub | | |

0 25 50 75 MILES
 32 MILES TO THE INCH, APPROXIMATELY

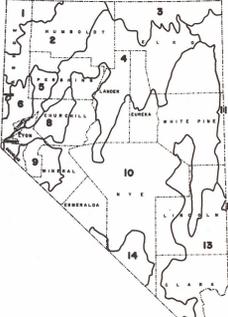
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 VEGETAL COVER MAP

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INDEX MAP OF NEVADA SHOWING HYDROGRAPHIC REGIONS AND BASINS



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 - 3. SNAKE RIVER BASIN
 - 4. HUMBOLDT RIVER BASIN
 - 5. WEST CENTRAL REGION
 - 6. TRUCKEE RIVER BASIN
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 - 11. COLORADO RIVER BASIN
 - 12. DEATH VALLEY BASIN
- | | | | | | | | |
|---|--|--|---|--|---|--|--|
| <p>NORTHWEST REGION</p> <ul style="list-style-type: none"> 1. BOULDER V. 2. COLLEMAN V. 3. CONTINENTAL LAKE V. 4. COOK LAKE V. 5. SPRAY LAKE V. 6. COOK LAKE V. 7. LONG V. 8. MARY FLAT 9. MASSACHUSETT LAKE V. 10. MONTICLO V. 11. SUPPER V. 12. SWAN LAKE V. 13. VIRBON V. 14. WARD V. | <p>SNAKE RIVER BASIN</p> <ul style="list-style-type: none"> 15. BRUNEAU RIVER AREA 16. HOOVER CREEK AREA 17. HOOVER CREEK AREA 18. JORDON CREEK AREA 19. LITTLE OWENS RIVER AREA 20. LITTLE OWENS RIVER AREA 21. LITTLE OWENS RIVER AREA 22. SALMON FALLS CREEK AREA 23. SOUTH FORK OWENS RIVER AREA | <p>WEST CENTRAL REGION</p> <ul style="list-style-type: none"> 24. BRADYS HOT SPRINGS AREA 25. FENLEY AREA 26. GRANITE SPRINGS V. 27. KUMON V. | <p>TRUCKEE RIVER BASIN</p> <ul style="list-style-type: none"> 28. DOGGE FLAT 29. LAKE THANE BASIN 30. PLEASANT V. 31. PYRAMID LAKE V. 32. TRUCKEE SPRINGS V. 33. SAN VALLEY 34. TRUCKEE CANYON SEGMENT 35. TRUCKEE RIVER BASIN 36. WARM SPRINGS V. 37. WINEMOCCA LAKE V. | <p>CARSON RIVER BASIN</p> <ul style="list-style-type: none"> 38. CARSON DESERT 39. CARSON V. 40. CHURCHILL V. 41. DAYTON V. 42. EARLY V. | <p>WESTERN REGION</p> <ul style="list-style-type: none"> 43. ANTELOPE V. 44. BENELL FLAT 45. COLE SPRING V. 46. DRY CREEK V. 47. EAST WALKER AREA 48. MADON V. 49. SMITH VALLEY 50. WALKER LAKE V. 51. WALKER LAKE V. 52. WALKER LAKE V. 53. WALKER LAKE V. 54. WALKER LAKE V. 55. WALKER LAKE V. 56. WALKER LAKE V. 57. WALKER LAKE V. 58. WALKER LAKE V. 59. WALKER LAKE V. 60. WALKER LAKE V. 61. WALKER LAKE V. 62. WALKER LAKE V. 63. WALKER LAKE V. 64. WALKER LAKE V. 65. WALKER LAKE V. 66. WALKER LAKE V. 67. WALKER LAKE V. 68. WALKER LAKE V. 69. WALKER LAKE V. 70. WALKER LAKE V. | <p>WALKER RIVER BASIN</p> <ul style="list-style-type: none"> 71. ANTELOPE V. 72. EAST WALKER AREA 73. MADON V. 74. SMITH VALLEY 75. WALKER LAKE V. 76. WALKER LAKE V. 77. WALKER LAKE V. 78. WALKER LAKE V. 79. WALKER LAKE V. 80. WALKER LAKE V. 81. WALKER LAKE V. 82. WALKER LAKE V. 83. WALKER LAKE V. 84. WALKER LAKE V. 85. WALKER LAKE V. 86. WALKER LAKE V. 87. WALKER LAKE V. 88. WALKER LAKE V. 89. WALKER LAKE V. 90. WALKER LAKE V. | <p>CENTRAL REGION</p> <ul style="list-style-type: none"> 91. COLUMBUS SALT MARSH V. 92. CONWAY V. 93. DELMAR V. 94. DANFORD V. 95. DOCK VALLEY 96. DOCK VALLEY 97. DOCK VALLEY 98. DOCK VALLEY 99. DOCK VALLEY 100. DOCK VALLEY 101. DOCK VALLEY 102. DOCK VALLEY 103. DOCK VALLEY 104. DOCK VALLEY 105. DOCK VALLEY 106. DOCK VALLEY 107. DOCK VALLEY 108. DOCK VALLEY 109. DOCK VALLEY 110. DOCK VALLEY 111. DOCK VALLEY 112. DOCK VALLEY 113. DOCK VALLEY 114. DOCK VALLEY 115. DOCK VALLEY 116. DOCK VALLEY 117. DOCK VALLEY 118. DOCK VALLEY 119. DOCK VALLEY 120. DOCK VALLEY 121. DOCK VALLEY 122. DOCK VALLEY 123. DOCK VALLEY 124. DOCK VALLEY 125. DOCK VALLEY 126. DOCK VALLEY 127. DOCK VALLEY 128. DOCK VALLEY 129. DOCK VALLEY 130. DOCK VALLEY 131. DOCK VALLEY 132. DOCK VALLEY 133. DOCK VALLEY 134. DOCK VALLEY 135. DOCK VALLEY 136. DOCK VALLEY 137. DOCK VALLEY 138. DOCK VALLEY 139. DOCK VALLEY 140. DOCK VALLEY |
|---|--|--|---|--|---|--|--|

EXPLANATION

This map shows estimated annual surface and groundwater flows between hydrographic areas and across state lines. Also shown for each area where data are available are estimates of annual runoff, perennial yield and water stored in the upper 100 feet of the groundwater reservoir. The term runoff refers to the estimated annual amount of surface water which flows from the mountains to the alluvial plain where the two meet. Perennial yield is the amount of groundwater which can be removed from a hydrographic area each year without depleting the groundwater reservoir. Where perennial yield was not available system yield is shown followed by the letter Y. System yield is the maximum amount of surface and groundwater which can be removed from a hydrographic area each year for an indefinite period of time.

- LEGEND**
(All figures are in thousands of acre feet)
- 2 → Estimated annual surface water flow.
 - Estimated annual groundwater flow.
 - Man made diversion. Surface water source - estimated annual amount.
 - Man made diversion. Groundwater source - estimated annual amount.
 - Man made diversion. Spring source - estimated annual amount.
 - 2 Estimated annual runoff.
 - 12 Estimated perennial yield.
 - 1,200 Estimated storage.
 - M M Minor quantity. An amount which is either less than 500 acre feet per year, or small in comparison to other quantities in the particular hydrographic area.
 - S S Significant size. However, sufficient information is not currently available to make an estimate.
 - Quantity uncertain. Figure shown represents annual discharge of springs which this flow is known to supply, at least in part.

LEGEND

- BOUNDARY OF HYDROGRAPHIC REGIONS
- BOUNDARY OF HYDROGRAPHIC AREAS (DOTTED WHERE ARBITRARY)

Compiled by:
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B. R. SCOTT
A. S. VAN DENBURGH
B. J. WASEY

Drawn by:
L. M. ROACH JR.

SEPTEMBER 1971

STATE OF NEVADA
WATER RESOURCES
AND
INTER-BASIN FLOWS

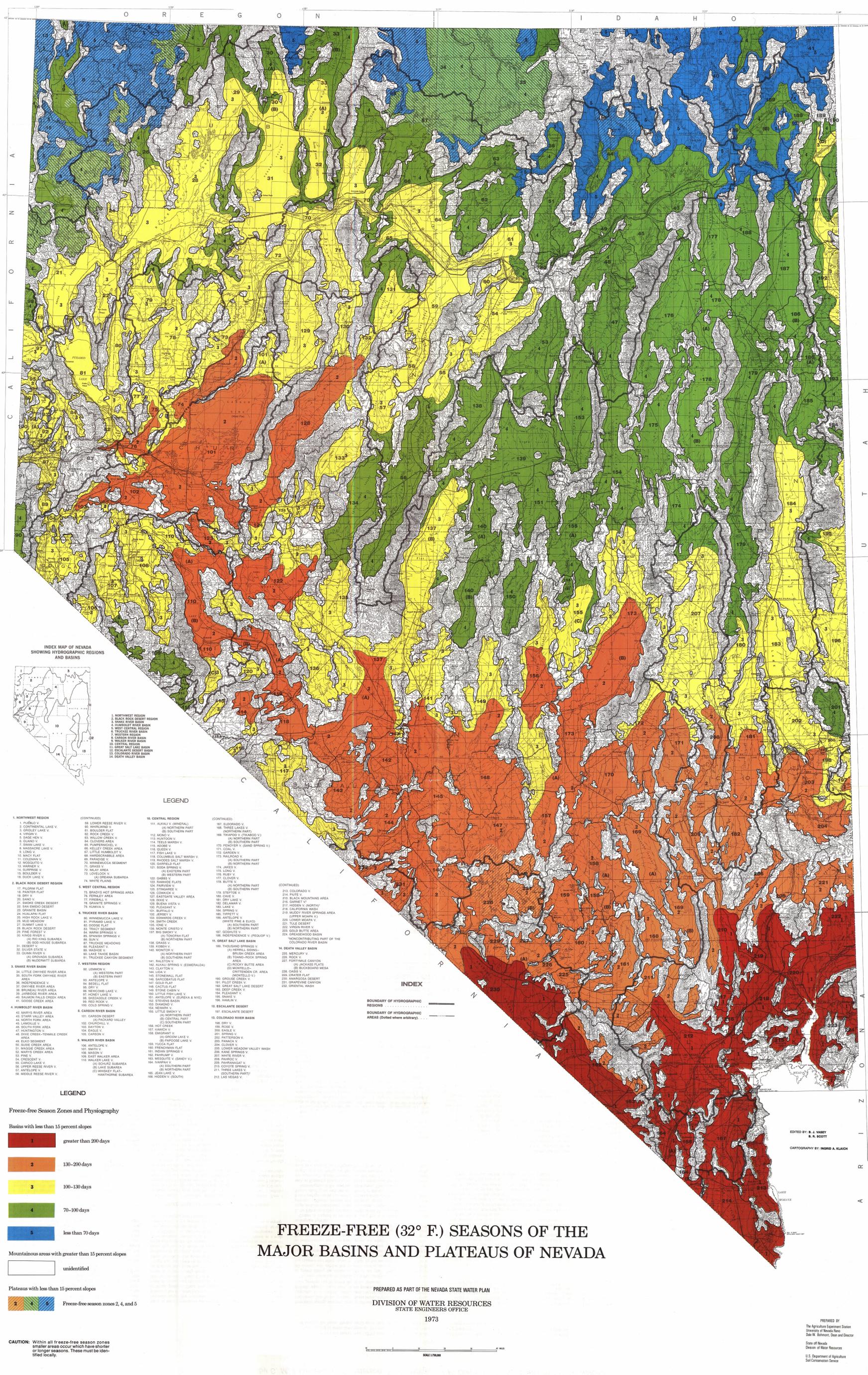
PREPARED BY
DIVISION OF WATER RESOURCES
STATE ENGINEERS OFFICE

SCALE 1:750,000

0 10 20 30 40 50 MILES
0 10 20 30 40 50 KILOMETERS

NOTE

Surface water flows are based on varying periods of record. Man made diversions are estimated 1970 figures except those on the Truckee Canal which are estimates based on historic diversion. Southern Nevada Project is shown as projected first stage level. Inflow from Arizona computed from gaged flow at Hoover Dam (Period of record 1934 to 1969). Outflow from Nevada based on Davis Dam gage (Period of record 1949-1969). Include Diversion estimated for 1971.



- LEGEND**
- | | | | | | | | | | |
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| <p>1. NORTHWEST REGION</p> <p>1. CONTINENTAL LAKE V. AREA
2. SPINNEY V. AREA
3. SPINNEY V. AREA
4. SPINNEY V. AREA
5. SPINNEY V. AREA
6. SPINNEY V. AREA
7. SPINNEY V. AREA
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11. SPINNEY V. AREA
12. SPINNEY V. AREA
13. SPINNEY V. AREA
14. SPINNEY V. AREA
15. SPINNEY V. AREA</p> | <p>2. BLACK ROCK DESERT REGION</p> <p>16. PANTER FLAT
17. PANTER FLAT
18. PANTER FLAT
19. PANTER FLAT
20. PANTER FLAT
21. PANTER FLAT
22. PANTER FLAT
23. PANTER FLAT
24. PANTER FLAT
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26. PANTER FLAT
27. PANTER FLAT
28. PANTER FLAT
29. PANTER FLAT
30. PANTER FLAT</p> | <p>3. WEST CENTRAL REGION</p> <p>31. BRADY'S HOT SPRINGS AREA
32. BRADY'S HOT SPRINGS AREA
33. BRADY'S HOT SPRINGS AREA
34. BRADY'S HOT SPRINGS AREA
35. BRADY'S HOT SPRINGS AREA
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43. BRADY'S HOT SPRINGS AREA
44. BRADY'S HOT SPRINGS AREA
45. BRADY'S HOT SPRINGS AREA</p> | <p>4. TRUCKEE RIVER BASIN</p> <p>46. TRUCKEE RIVER BASIN
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59. TRUCKEE RIVER BASIN
60. TRUCKEE RIVER BASIN</p> | <p>5. CARSON RIVER BASIN</p> <p>61. CARSON RIVER BASIN
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65. CARSON RIVER BASIN
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72. CARSON RIVER BASIN
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74. CARSON RIVER BASIN
75. CARSON RIVER BASIN</p> | <p>6. HUMBOLDT RIVER BASIN</p> <p>76. HUMBOLDT RIVER BASIN
77. HUMBOLDT RIVER BASIN
78. HUMBOLDT RIVER BASIN
79. HUMBOLDT RIVER BASIN
80. HUMBOLDT RIVER BASIN
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87. HUMBOLDT RIVER BASIN
88. HUMBOLDT RIVER BASIN
89. HUMBOLDT RIVER BASIN
90. HUMBOLDT RIVER BASIN</p> | <p>7. WESTERN REGION</p> <p>91. WESTERN REGION
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105. WESTERN REGION</p> | <p>8. CENTRAL REGION</p> <p>106. CENTRAL REGION
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117. CENTRAL REGION
118. CENTRAL REGION
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120. CENTRAL REGION</p> | <p>9. EASTERN REGION</p> <p>121. EASTERN REGION
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124. EASTERN REGION
125. EASTERN REGION
126. EASTERN REGION
127. EASTERN REGION
128. EASTERN REGION
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131. EASTERN REGION
132. EASTERN REGION
133. EASTERN REGION
134. EASTERN REGION
135. EASTERN REGION</p> | <p>10. DEATH VALLEY BASIN</p> <p>136. DEATH VALLEY BASIN
137. DEATH VALLEY BASIN
138. DEATH VALLEY BASIN
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140. DEATH VALLEY BASIN
141. DEATH VALLEY BASIN
142. DEATH VALLEY BASIN
143. DEATH VALLEY BASIN
144. DEATH VALLEY BASIN
145. DEATH VALLEY BASIN
146. DEATH VALLEY BASIN
147. DEATH VALLEY BASIN
148. DEATH VALLEY BASIN
149. DEATH VALLEY BASIN
150. DEATH VALLEY BASIN</p> |
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Freeze-free Season Zones and Physiography

Basins with less than 15 percent slopes

	1	greater than 200 days
	2	130-200 days
	3	100-130 days
	4	70-100 days
	5	less than 70 days

Mountainous areas with greater than 15 percent slopes

	undifferentiated
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Plateaus with less than 15 percent slopes

					Freeze-free season zones 2, 4, and 5
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CAUTION: Within all freeze-free season zones smaller areas occur which have shorter or longer seasons. These must be identified locally.

FREEZE-FREE (32° F.) SEASONS OF THE MAJOR BASINS AND PLATEAUS OF NEVADA

PREPARED AS PART OF THE NEVADA STATE WATER PLAN
DIVISION OF WATER RESOURCES
STATE ENGINEERS OFFICE
1973



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B. R. SCOTT
CARTOGRAPHY BY: INGRID A. KLACH

PREPARED BY
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Dean W. Bakstrom, Dean and Director
State of Nevada
Division of Water Resources
U.S. Department of Agriculture
Soil Conservation Service

BASE: REDUCTION OF U.S.G.S. TOPOGRAPHIC MAP OF NEVADA, 1:500,000, EDITION OF 1955



State of Nevada
WATER PLANNING
REPORT